

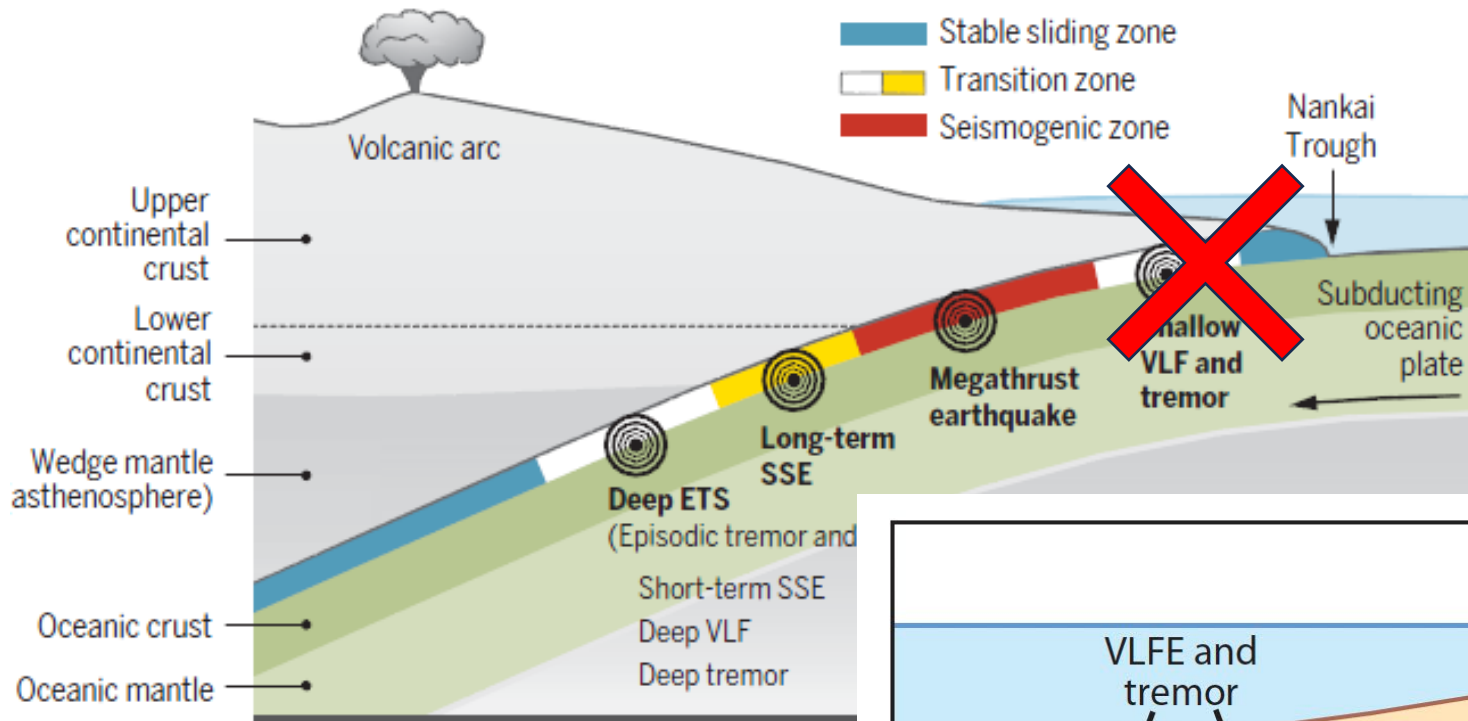
Evidence for megathrust locking and shallow coseismic slip in the Cascadia accretionary wedge



Harold Tobin, Anna Ledeczi, Madeleine
Lucas

University of Washington

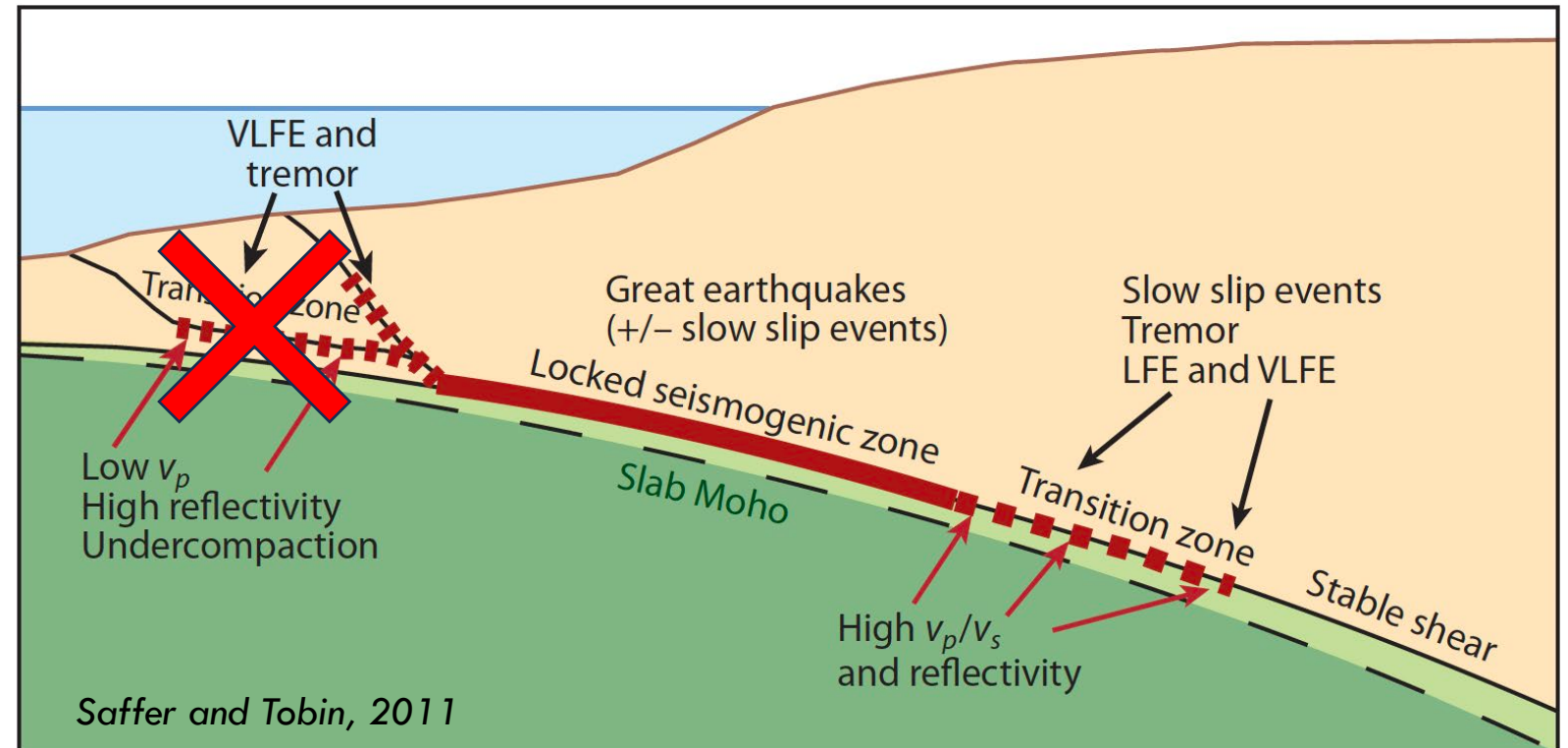




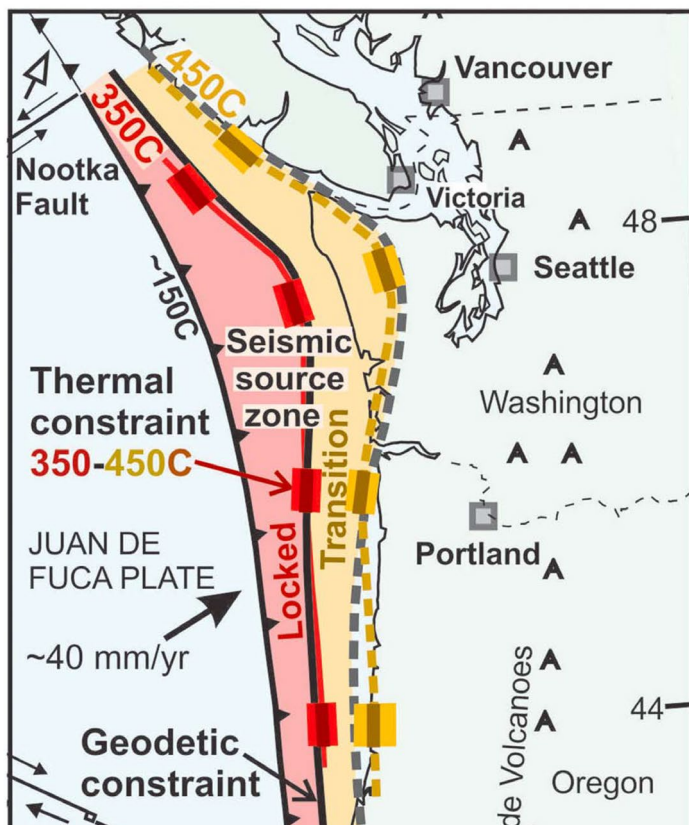
Obara and Kato, 2016

Generic paradigm:
Locked seismogenic zone is
flanked by aseismic or
conditionally stable regions

Cascadia is not typical!



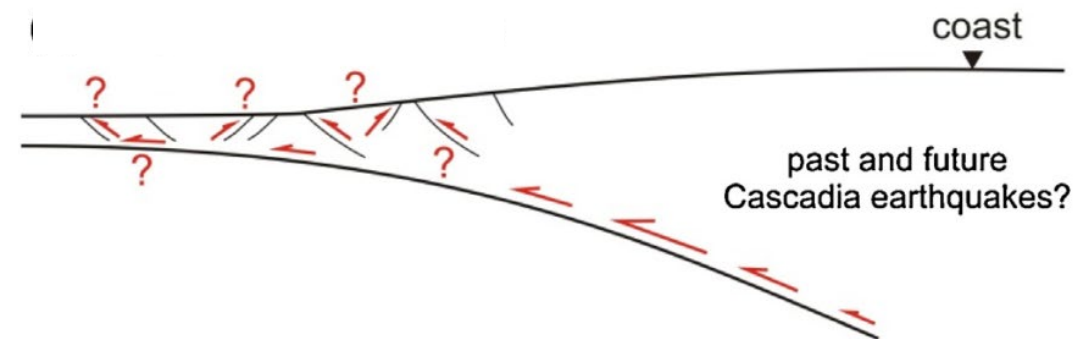
Saffer and Tobin, 2011



Hyndman et al., 1995, 2015

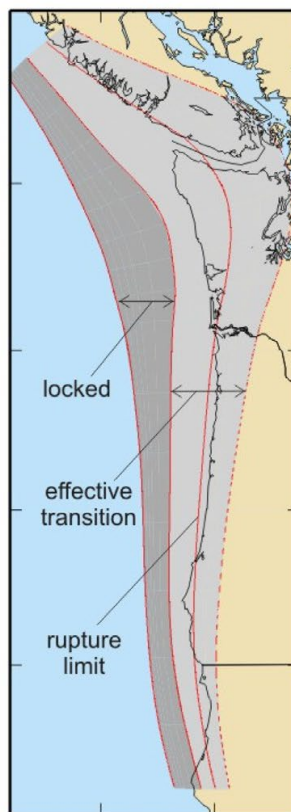
Thermal
controls

From the earliest work on the up-dip limit (Hyndman 1993, etc.), Cascadia was recognized as being $\geq 150^\circ$ at the deformation front



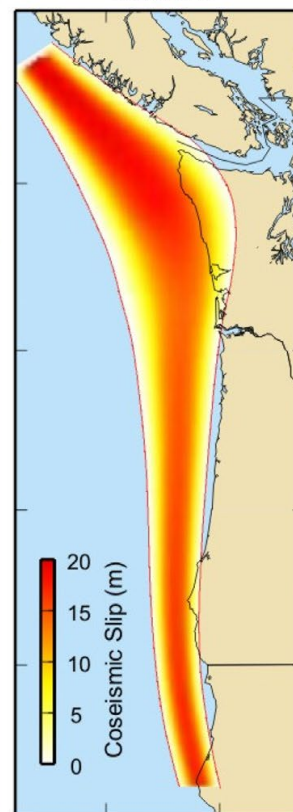
Wang and Trehu [2016]

(b) Wang et al. (2003)



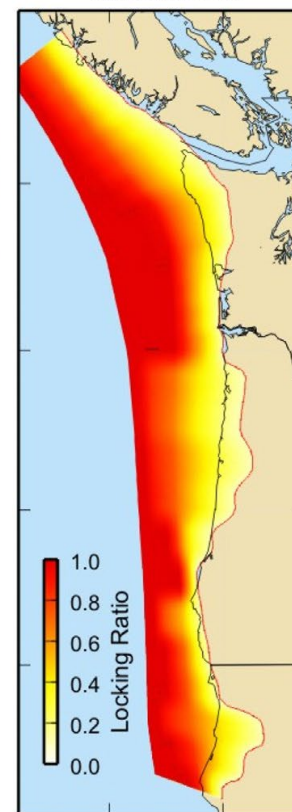
126°W 124°W

(c) Preist et al. (2010)
simplified



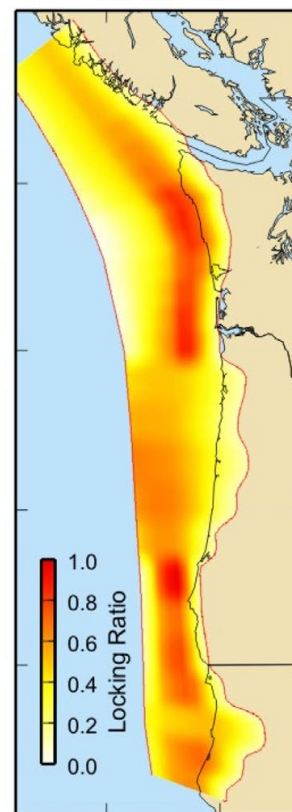
126°W 124°W

(d) Schmalzle et al. (2014)
Gamma model



126°W 124°W

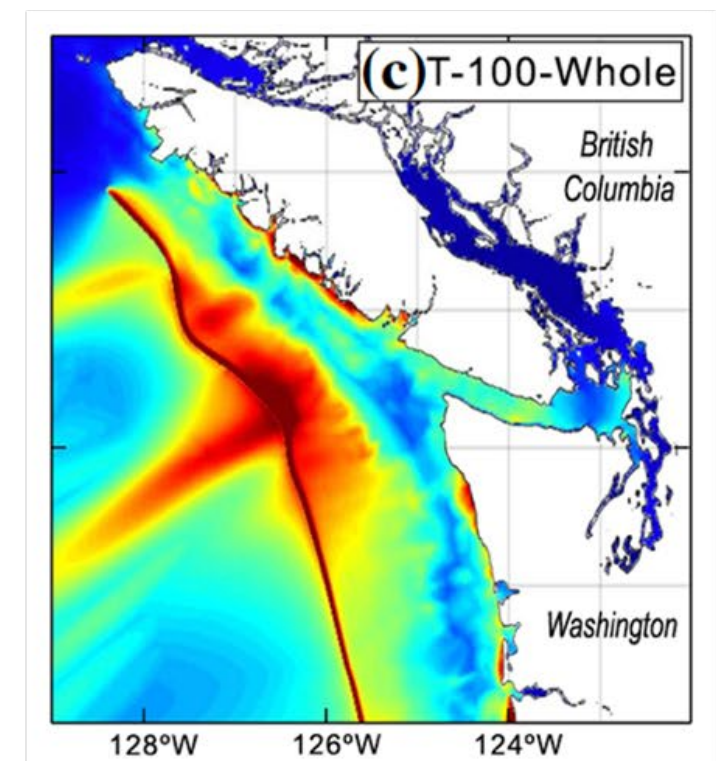
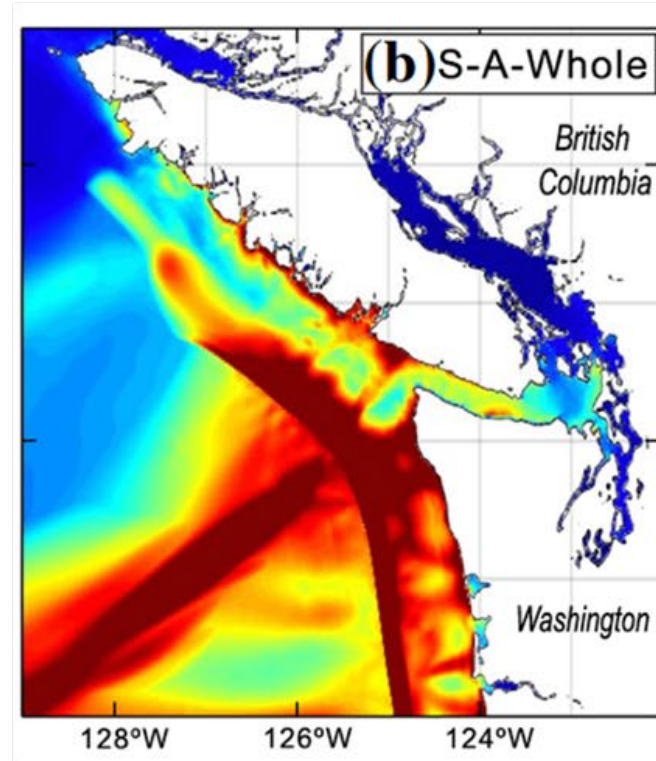
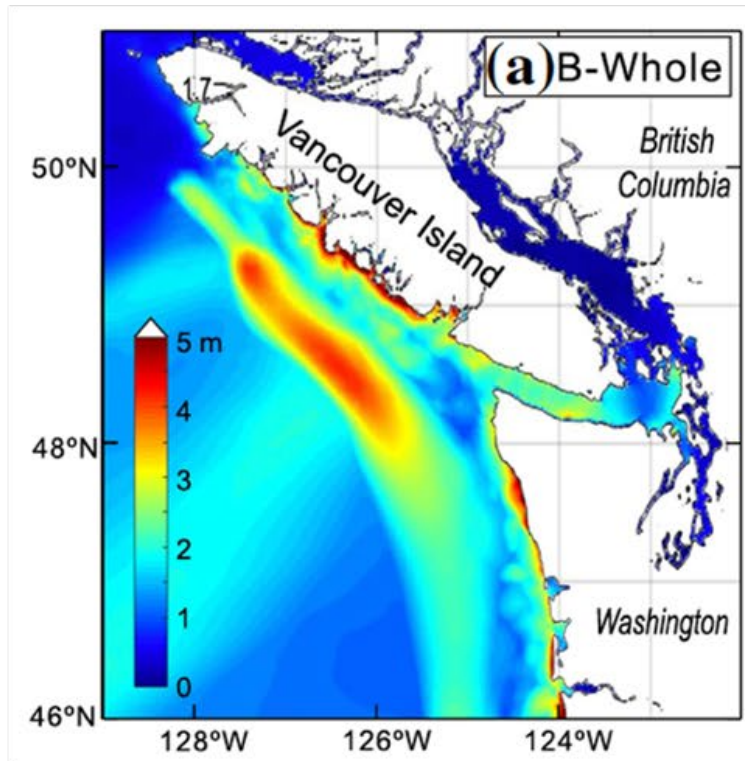
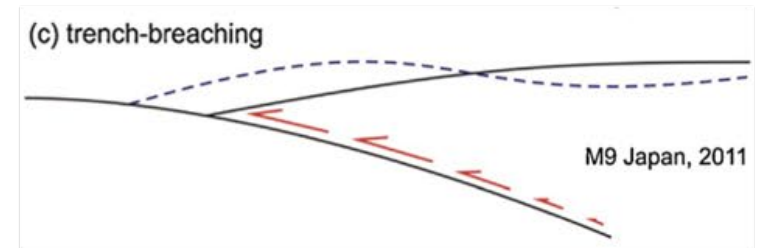
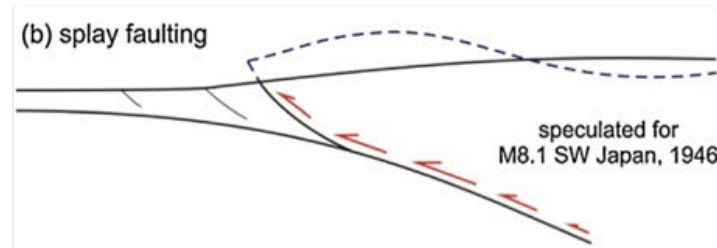
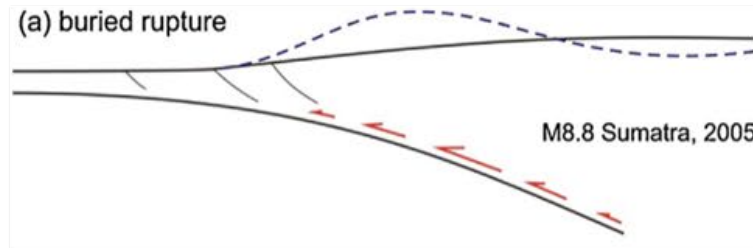
(e) Schmalzle et al. (2014)
Gaussian model



126°W 124°W

What is the shallow locking state in Cascadia? Where will rupture go?

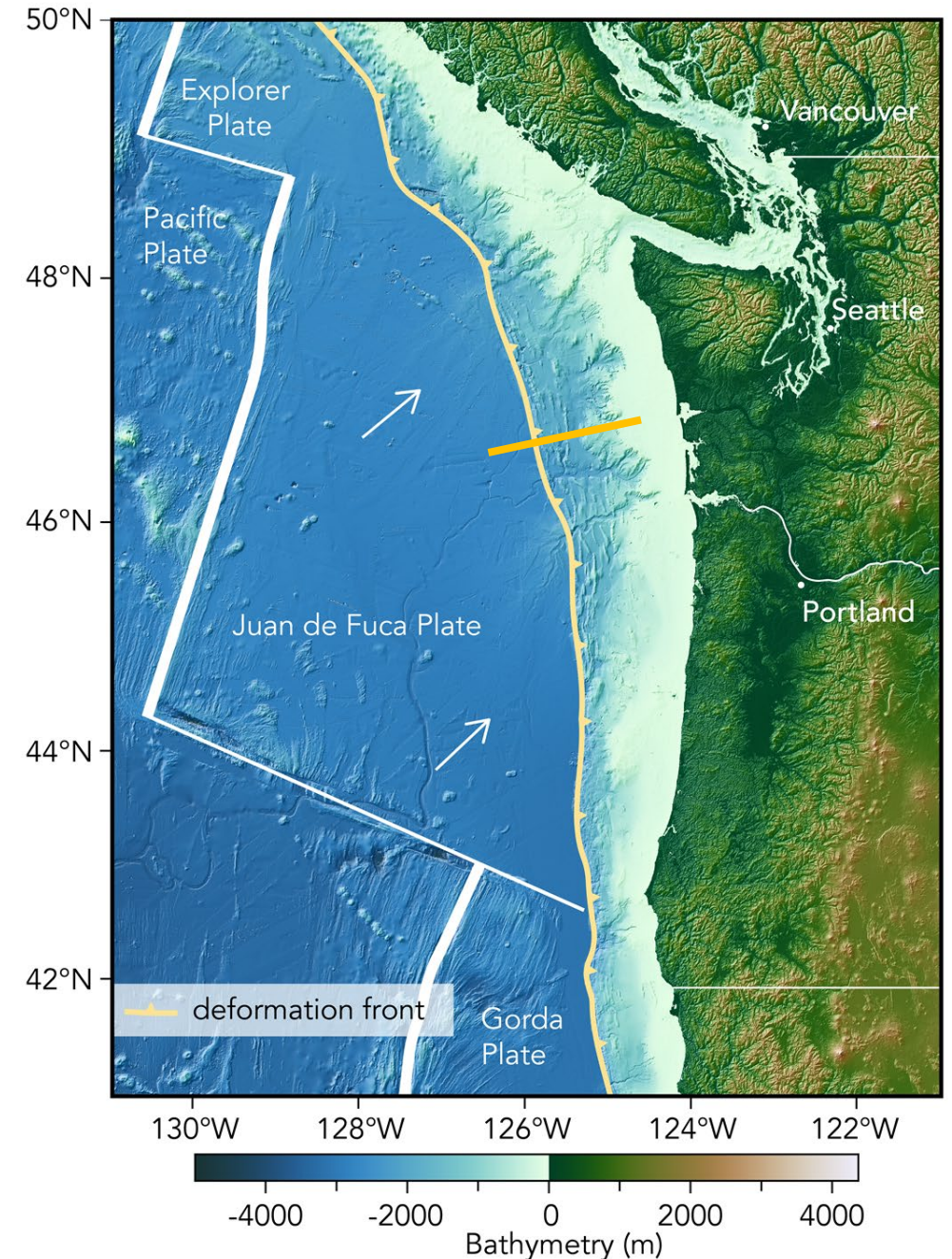
Shallow rupture and megasplay rupture both impact tsunami hazard



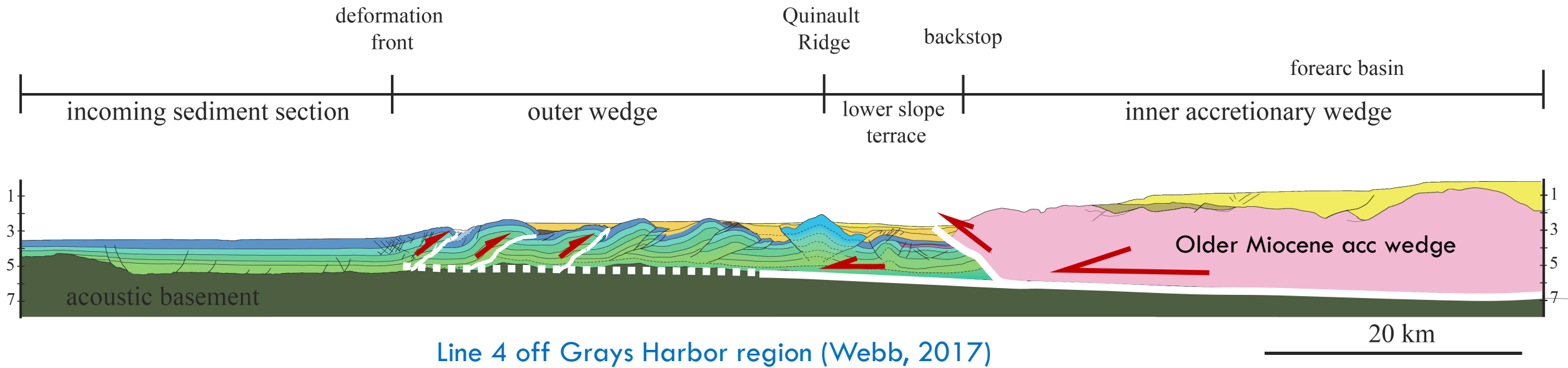
top: Wang & Trehu (2016) bottom: Gao et al. (2018)

Cascadia structure offshore

- Young, hot plate subducts ~ 40 mm/yr
- Potential for M9 class event:
 - Recurrence time ~ 450 -500 years
 - Most recent was 325 years ago
- Wide & high-volume accretionary wedge of deformed sediments above the shallowly-dipping megathrust fault
- Widest and flattest off Washington state where it is also now accumulating the greatest elastic strain (*slip deficit*)

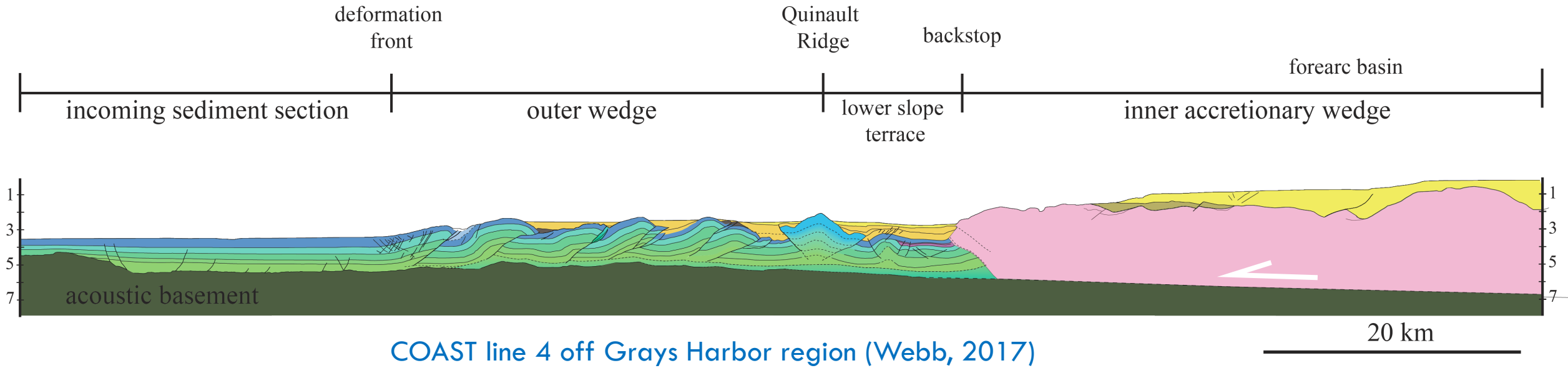


Central Cascadia: Evidence of outer wedge rupture and/or mega-splay fault slip



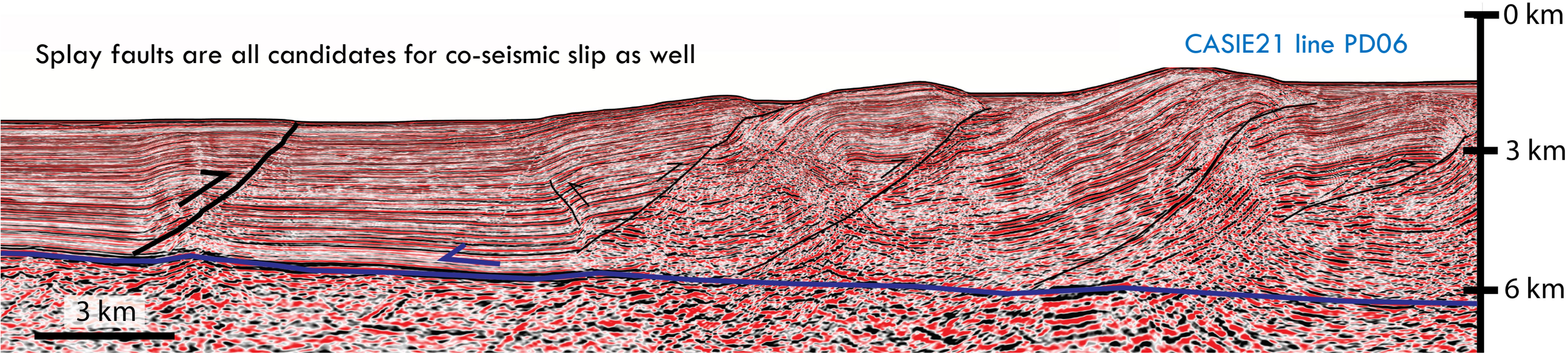
- *How far out to the deformation front does fast slip go on the megathrust décollement?*
- *What about shallow splay faults that reach the surface?*
- *Is there evidence of a mega-splay fault?*

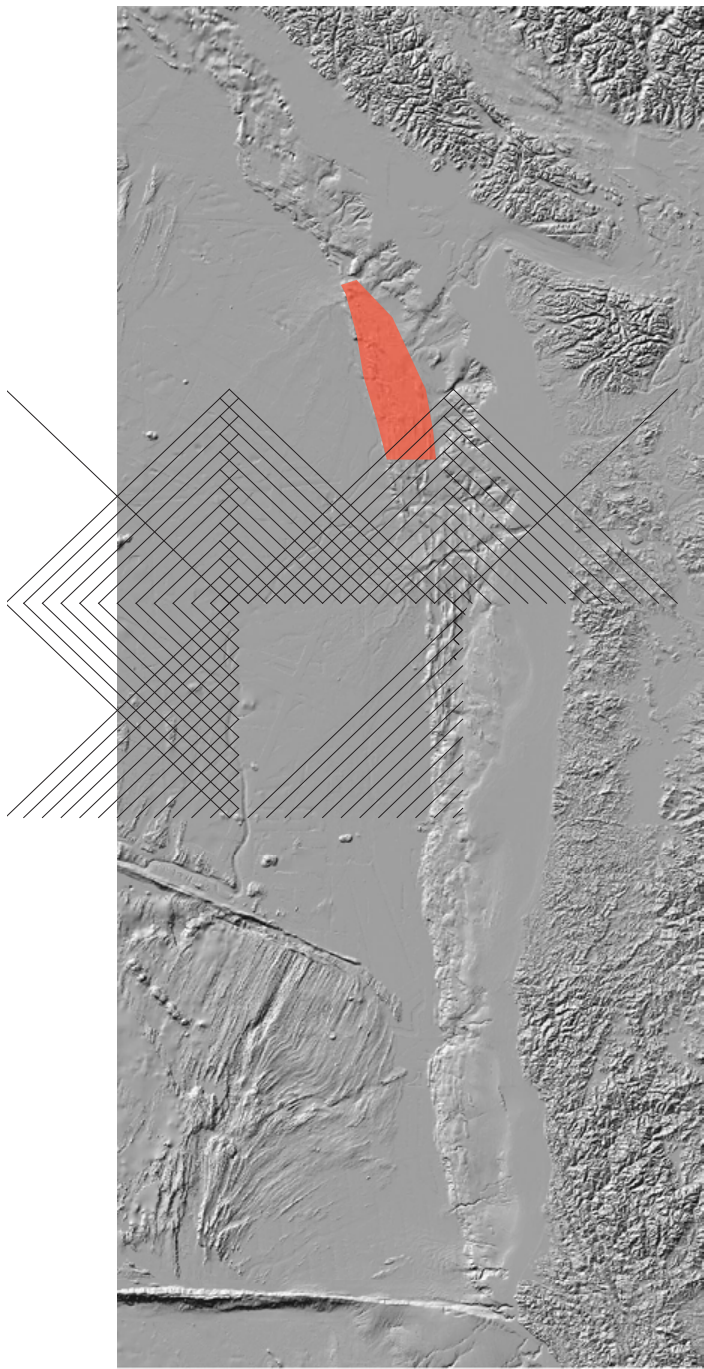
Cascadia's *landward-vergent* fold and thrust wedge



Decollement is the plate boundary, and lies **~3 km** below the surface right at the deformation front

Splay faults are all candidates for co-seismic slip as well



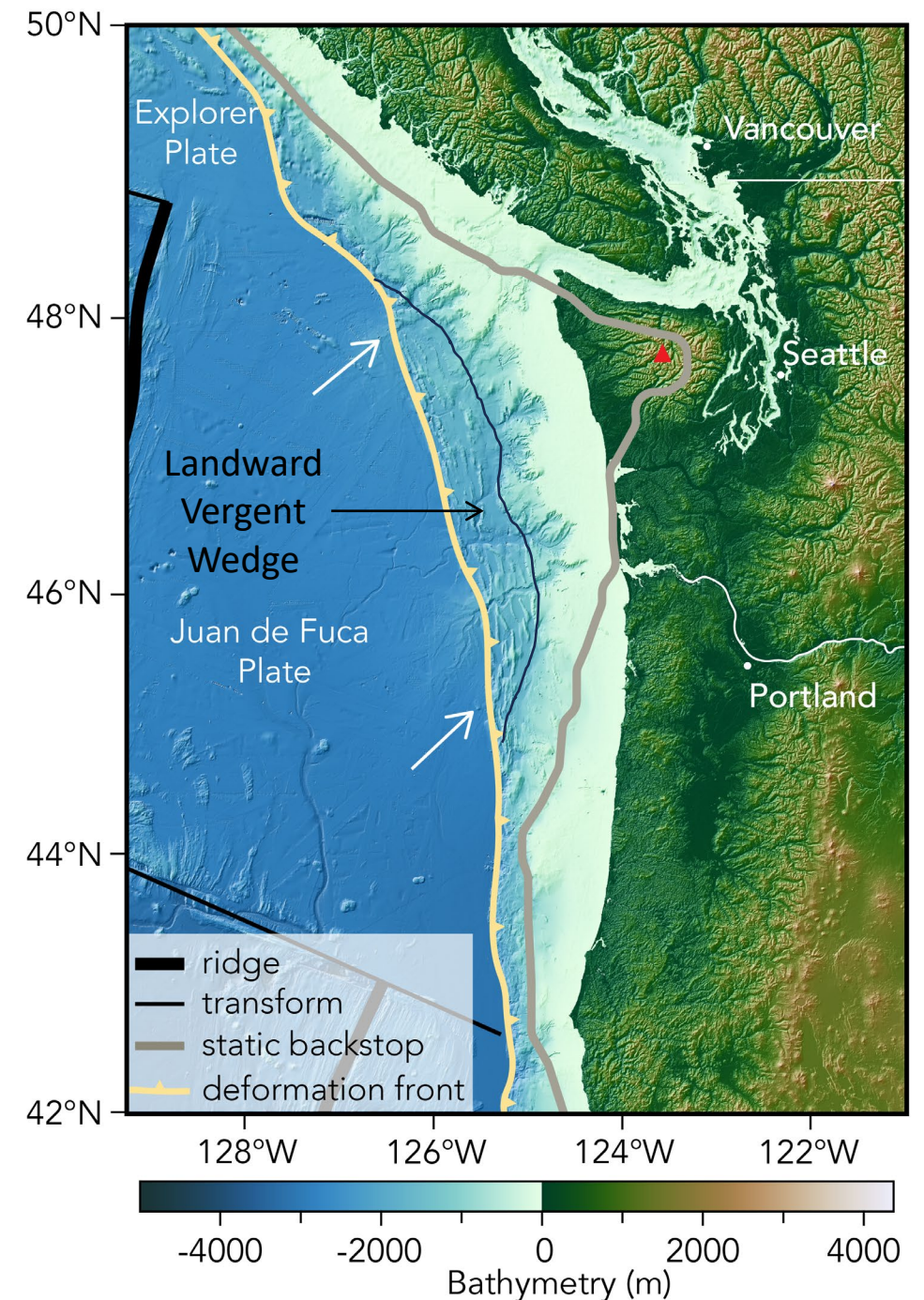


Landward Vergence Zone(s)

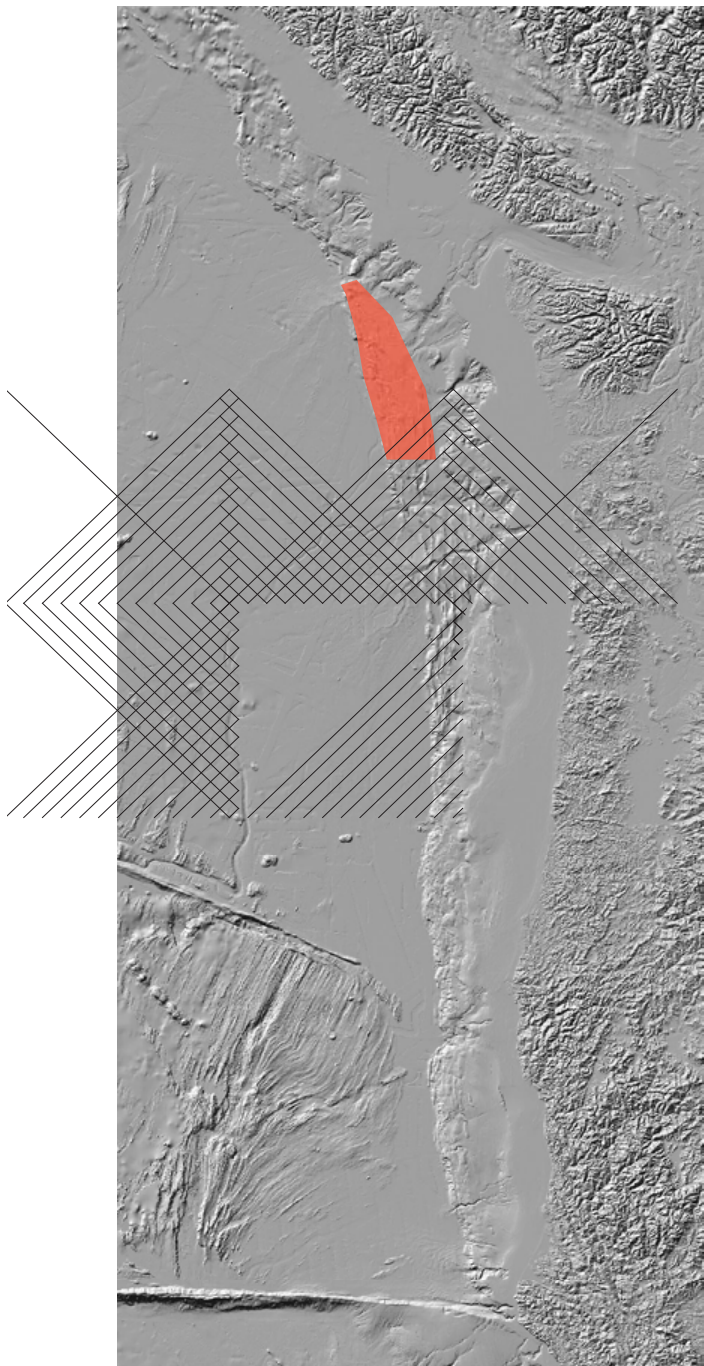
Landward vergence zone has been recognized since Seely, 1977 -- (also MacKay et al., 1995; Gulick et al., 1998; Adam et al., 2004, etc.)

Systematic LV is globally very rare – most extensively developed LVZ is Cascadia (also Sumatra 2004 M9)

Northern and southern patches at Cascadia



Watt & Brothers, 2021

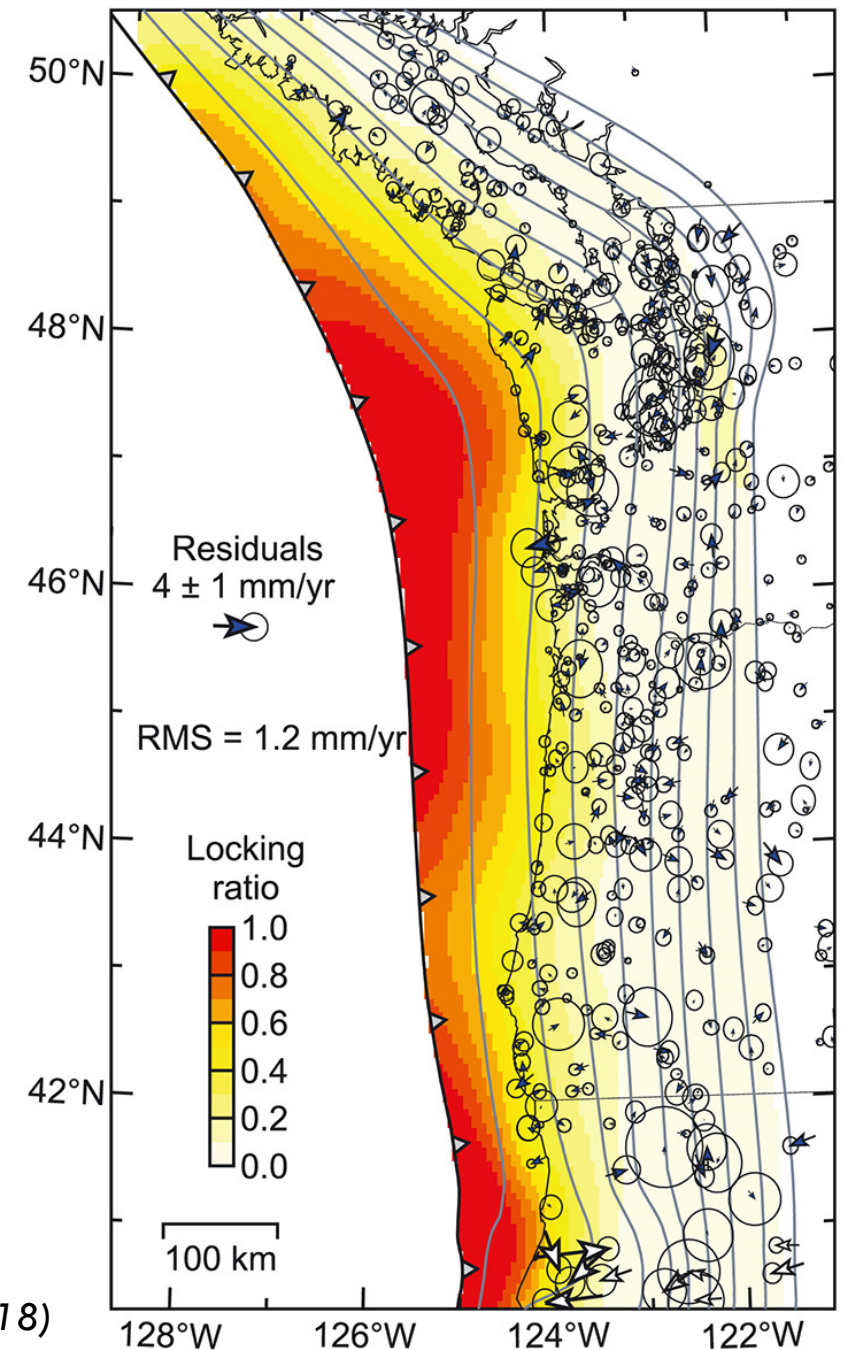


Landward Vergence and Geodetic Locking: Coincidence?

Geodetic locking models for
Cascadia suggest it is locked to the
trench for most of the length

The extent of the LV dominated
zone corresponds to this
geophysical evidence of present-
day locking and slip deficit of the
plate boundary

Watt & Brothers (2021)



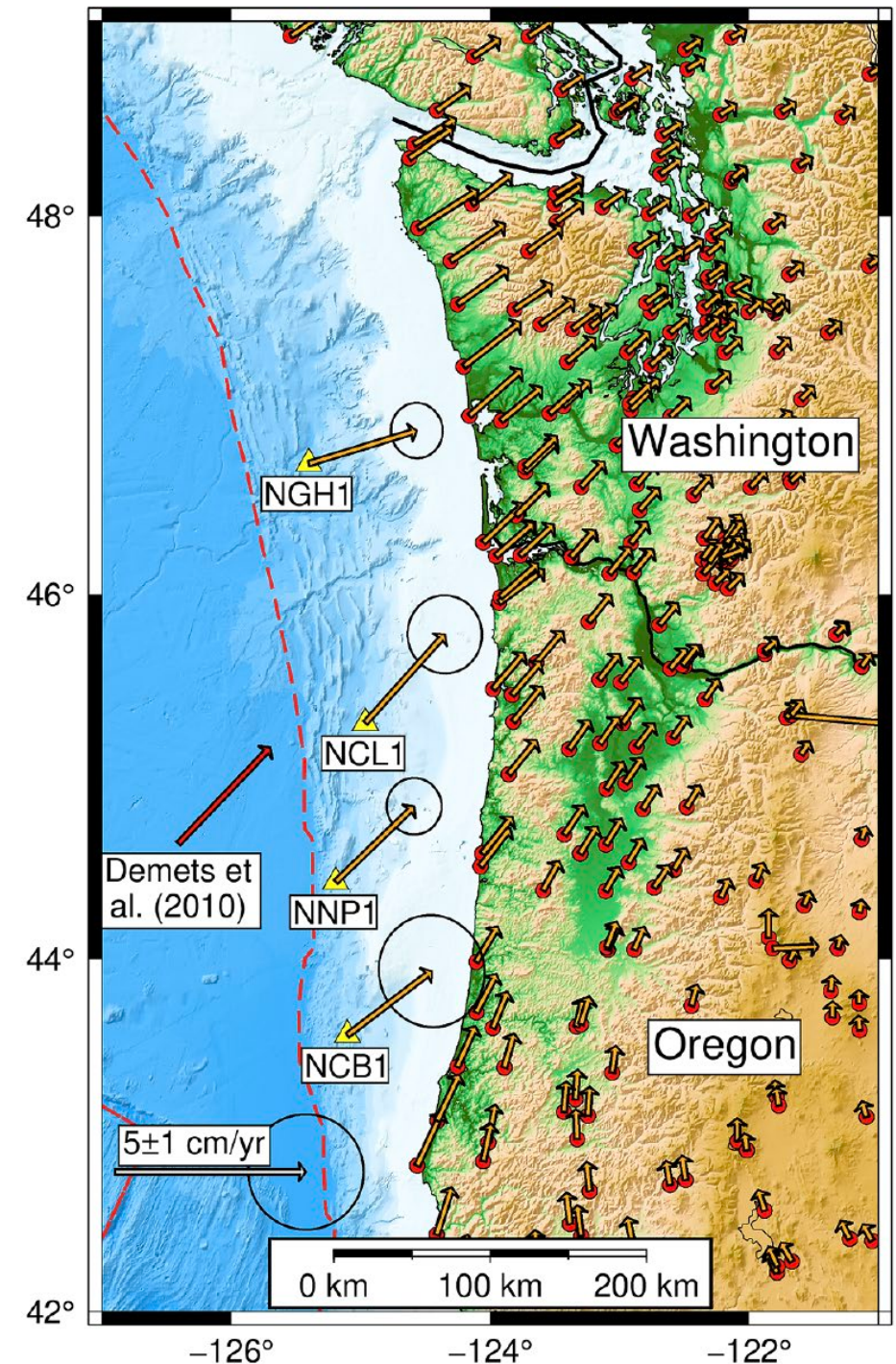
Li et al., (2018)

New offshore geodesy from GNSS-A

Consistent with models showing
strong up-dip locking

*Are conditions in the plate boundary
consistent with shallow locking?*

DeSanto et al., 2025 EPSL



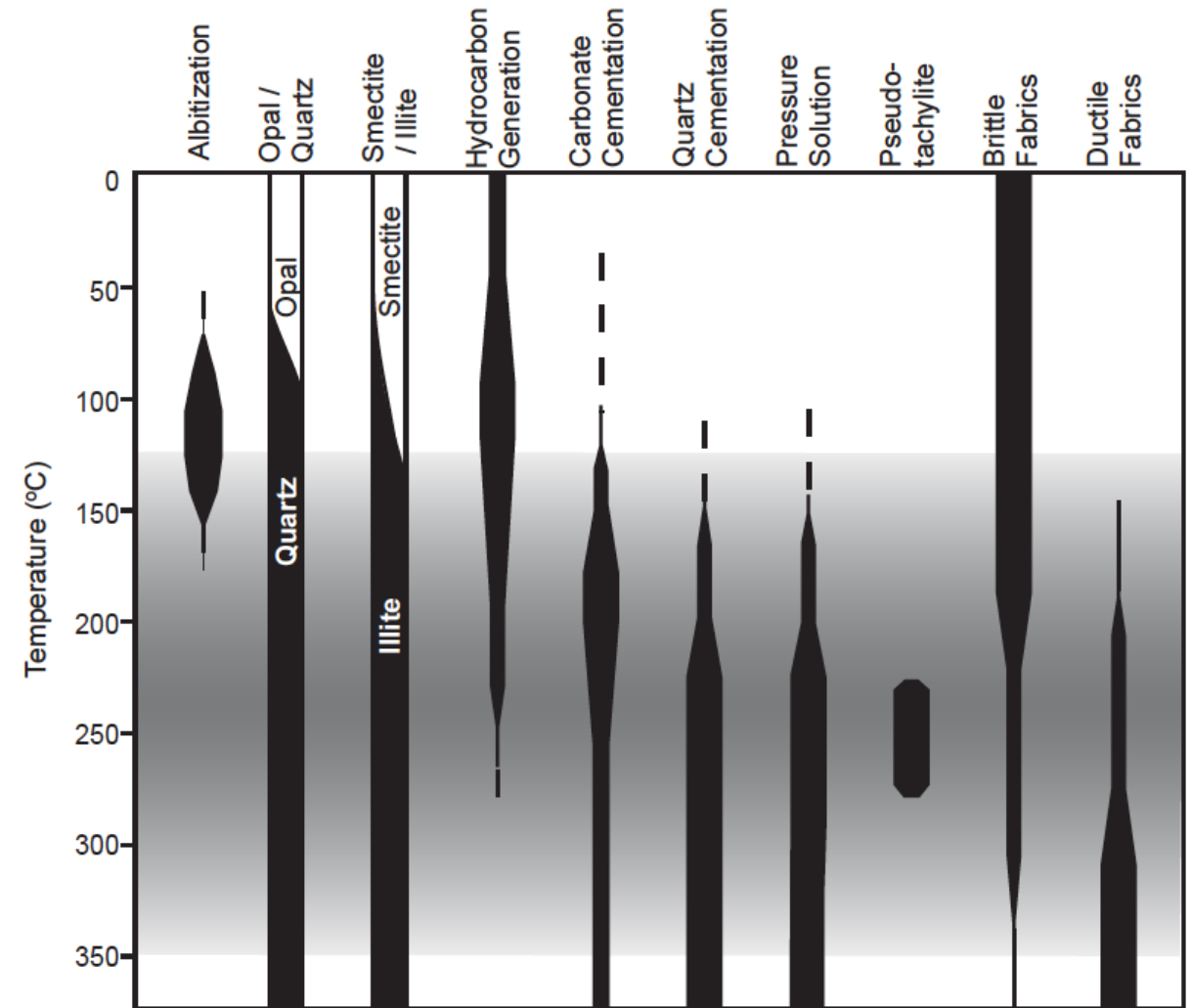
What controls frictional locking behavior of the shallow subduction zone decollement?

- **Fault-normal stress** magnitude (burial)
- **Strength** (rigidity) at fault depth: lithification and cementation
- **Pore fluid pressure**
- **Wall-rock strength** (compliance) especially the upper plate

Rate-state framework defining velocity-strengthening (stable creep) vs. velocity weakening (unstable & seismic)

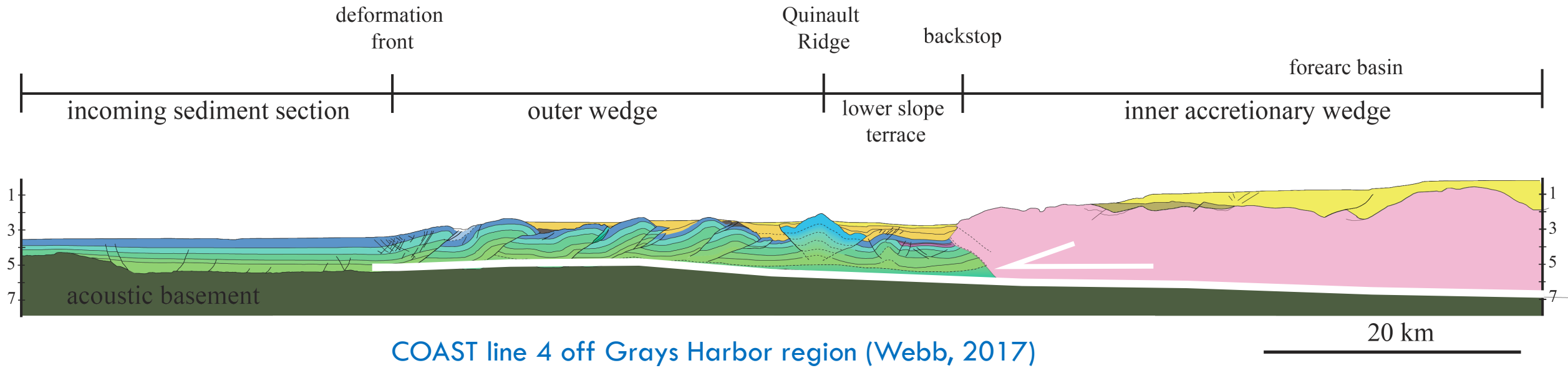
Mineralization and fluid sources change with temperature in typical s.z. sediments

→ *~100 - 150° C is a transition for many processes*

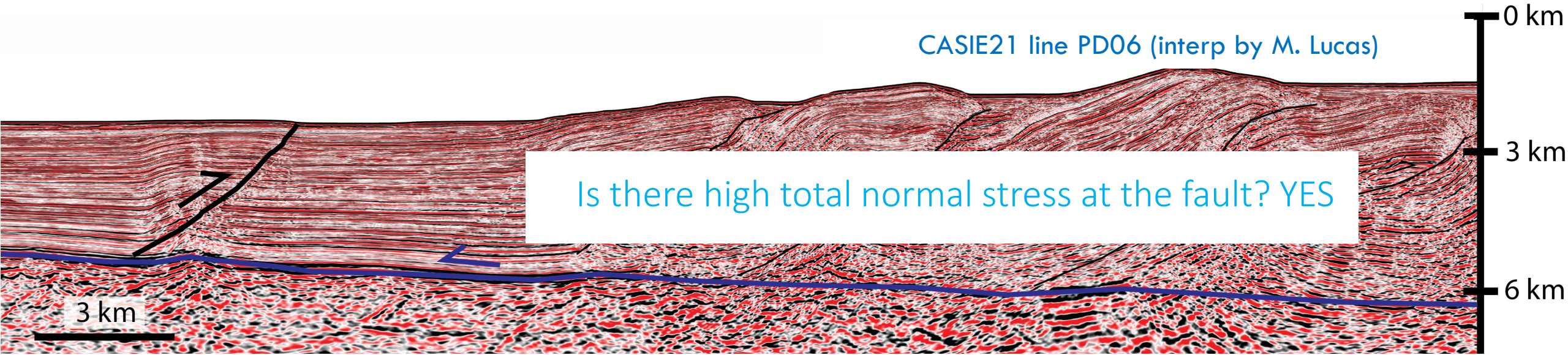


Moore, 2007

“Slip to the trench” = slip to the toe at décollement depth



Decollement is the plate boundary, and lies ~3 km below the surface right at the deformation front



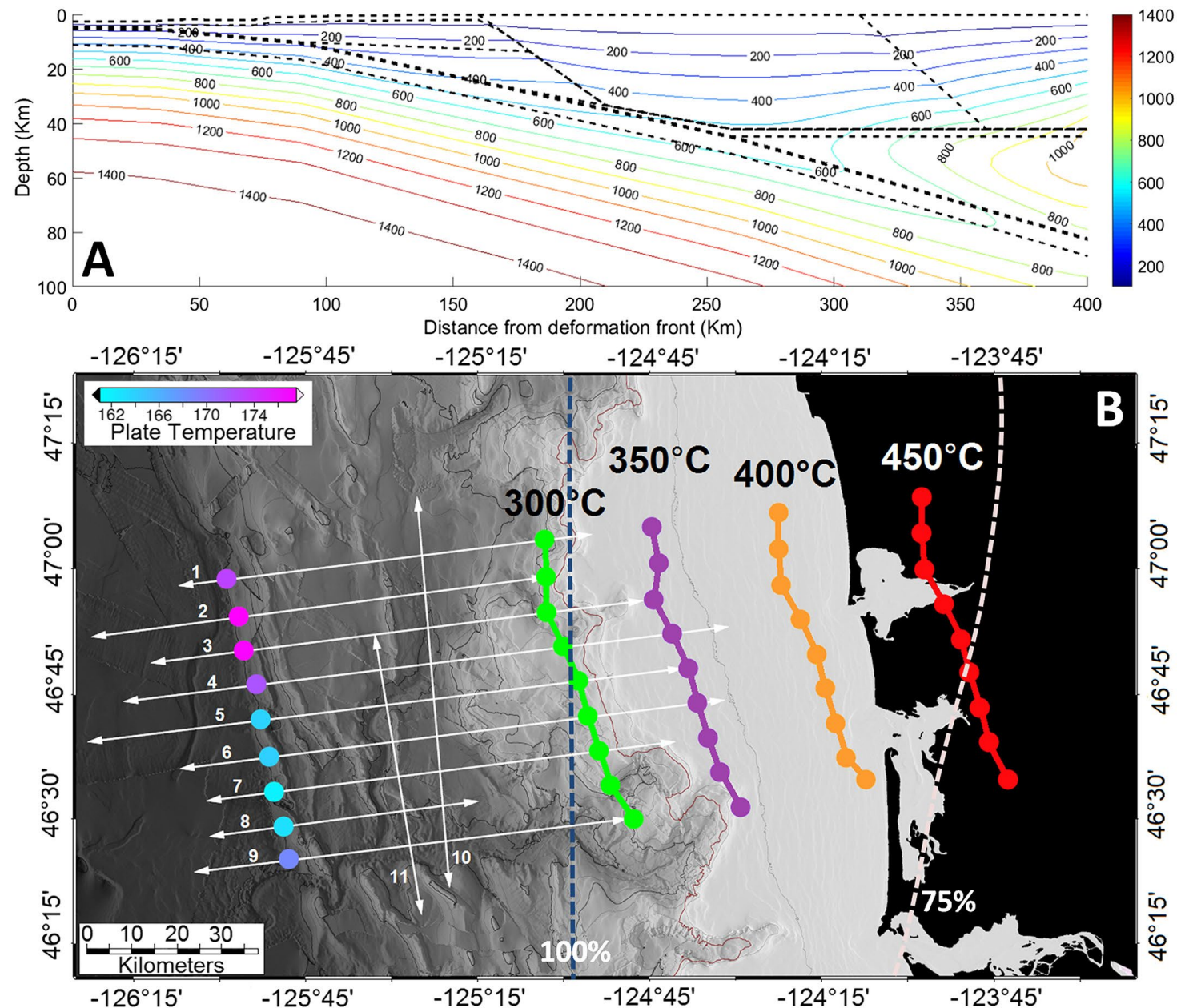
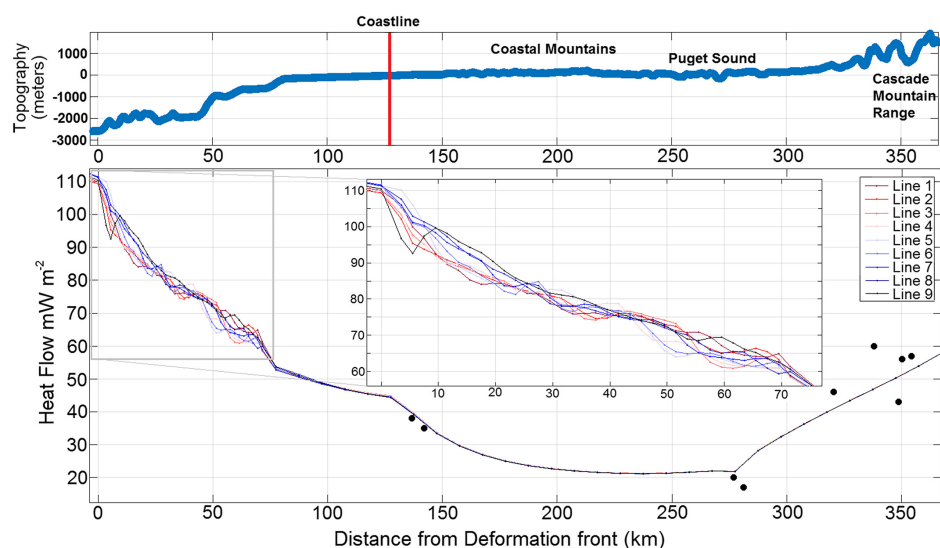
Is it hot enough? *YES*

Cascadia thermal models agree that the temperature at the base of the sediment section at the front exceeds $\sim 150^{\circ}\text{C}$

Salmi et al. (2017) heat flow data and BSR derived temperature gradient estimates to constrain a thermal model

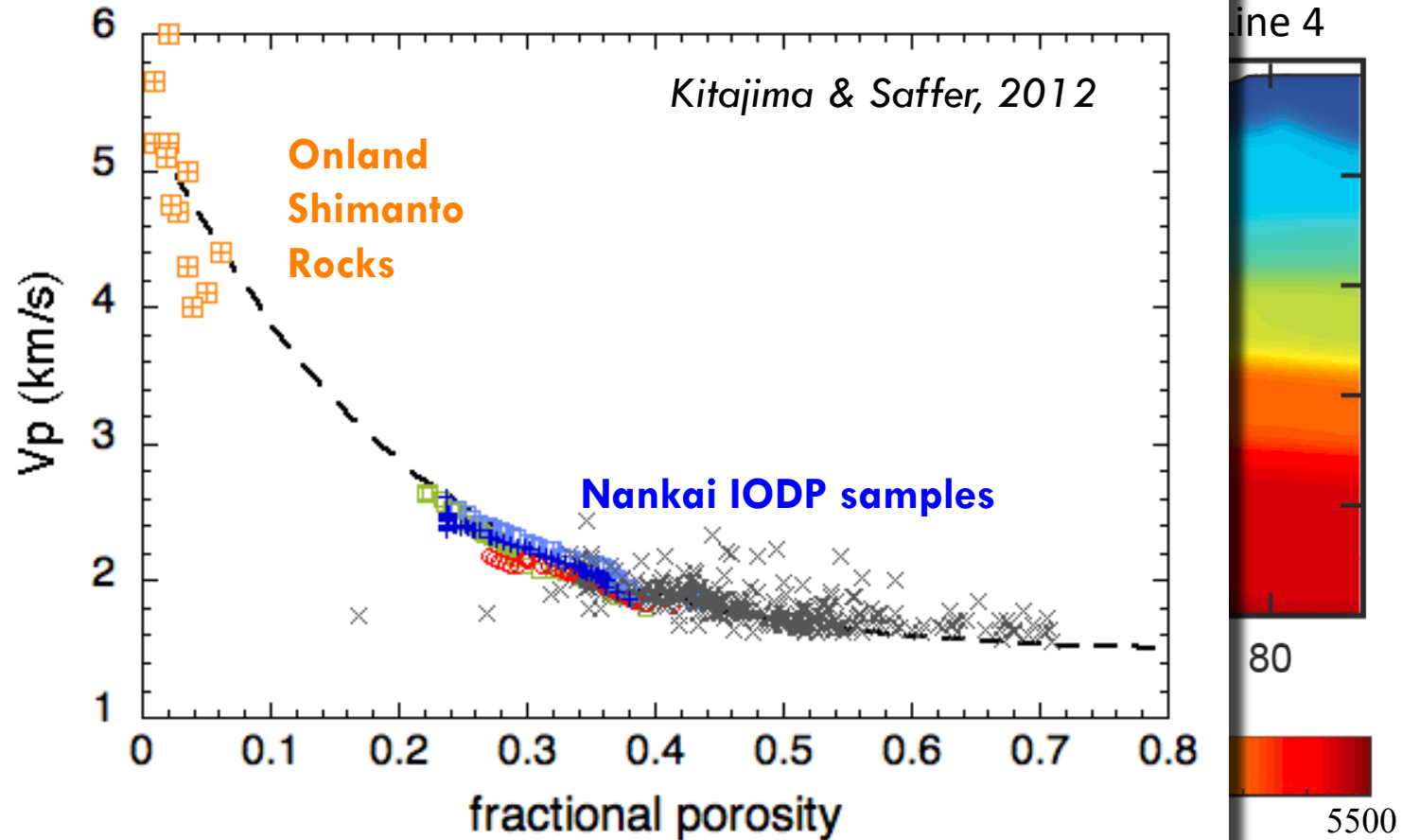
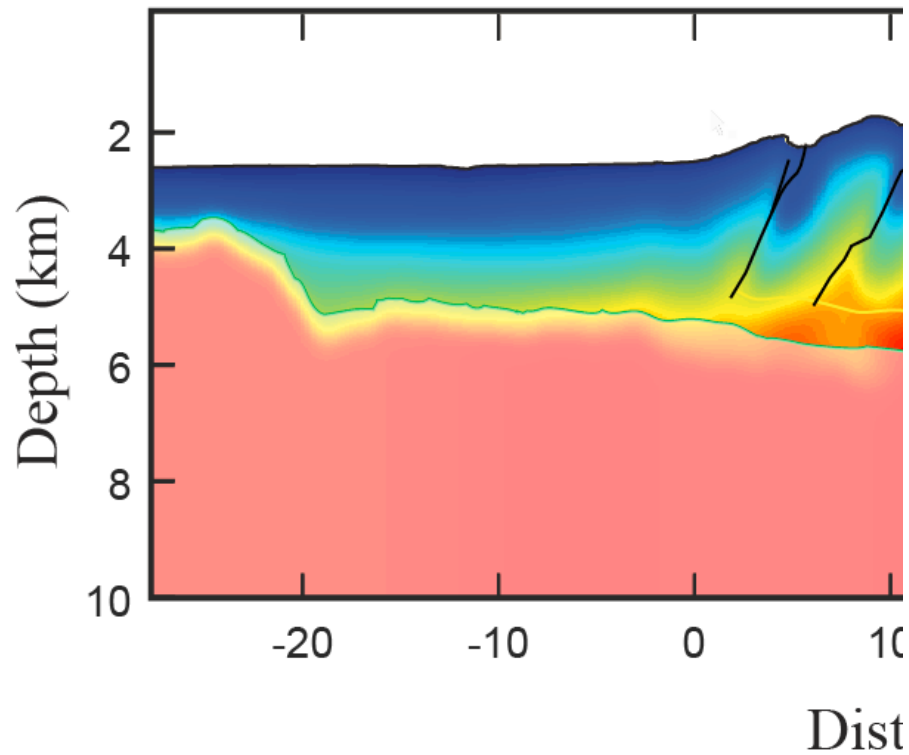
→ At the def. front, $T = \sim 170^{\circ}\text{C}$

→ Heat flow = 110 mW/m^2



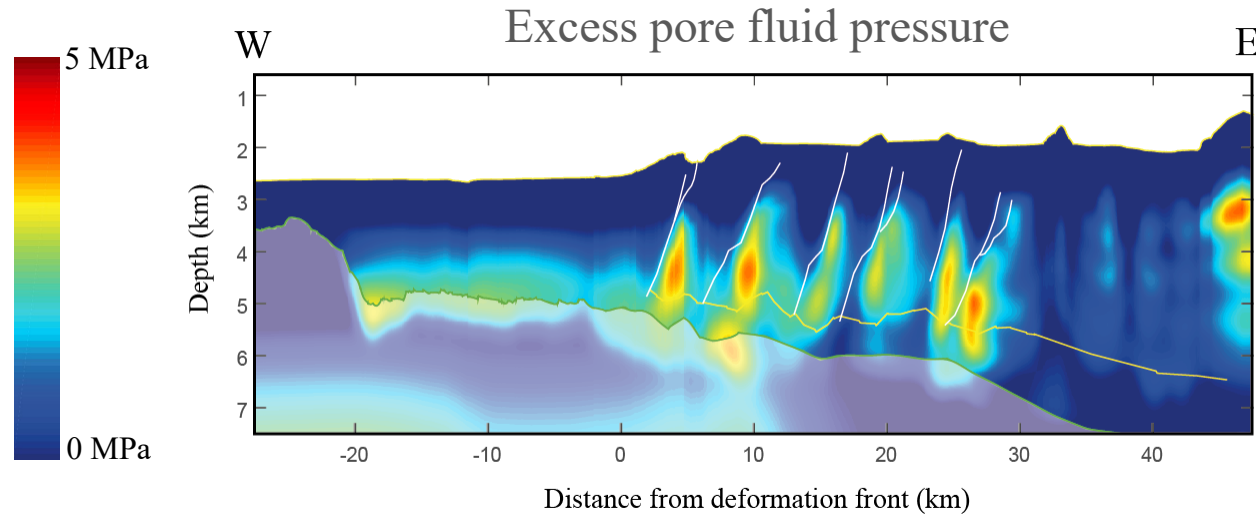
Are the rocks lithified?

V_p interval velocity from horizon-based tomography for Prestack Depth Migration



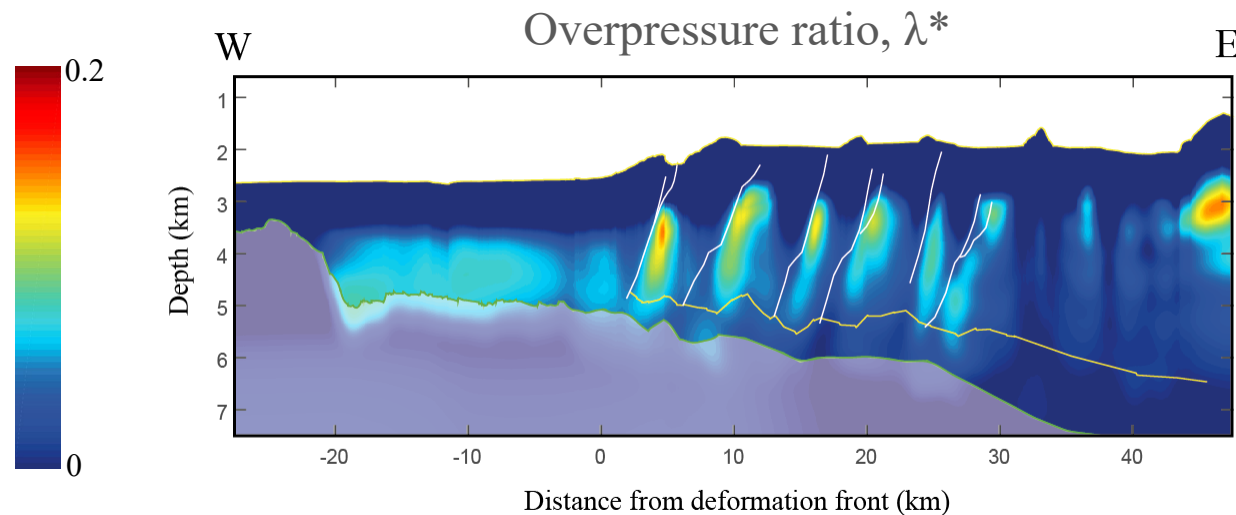
Decollement at the front is high V_p : $\geq 4000 \text{ m/s}$

Is there pore fluid overpressure? NO



Calculated apparent underconsolidation and therefore pore fluid overpressure from seismic interval velocity

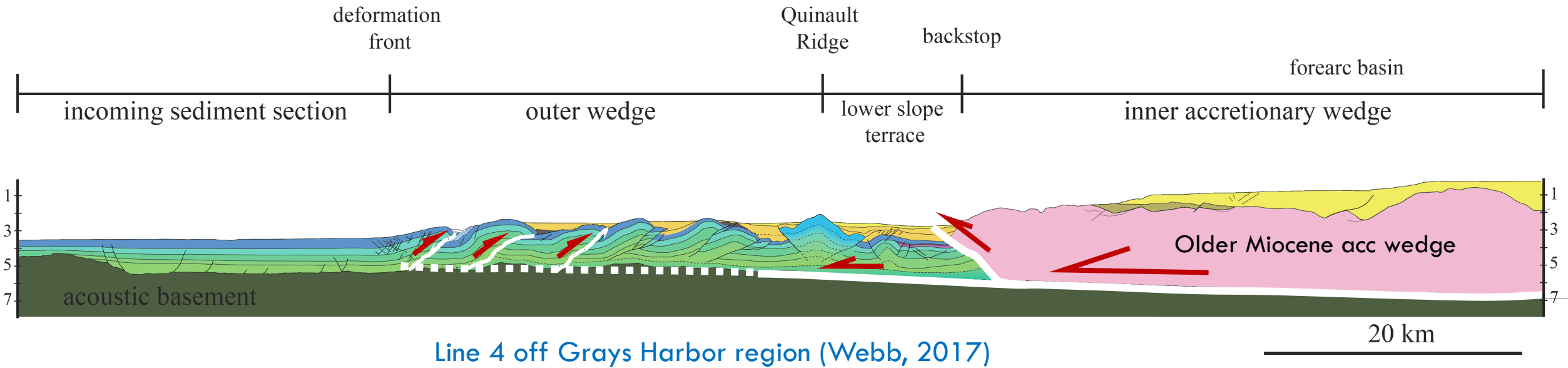
→ Evidence for only very minor overpressure at depth – close to hydrostatic



High seismic velocity, deep burial, normal pore pressure imply a strong wedge environment

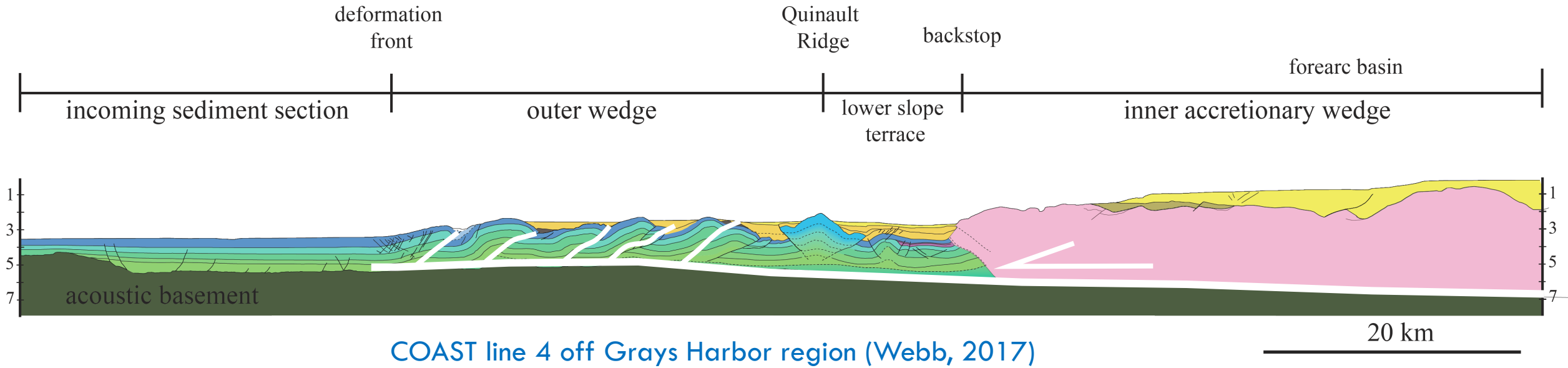
→ Frictional locking and unstable fast slip behavior are *expected*

Ok, so is there any evidence of shallow rupture and/or splay fault slip?



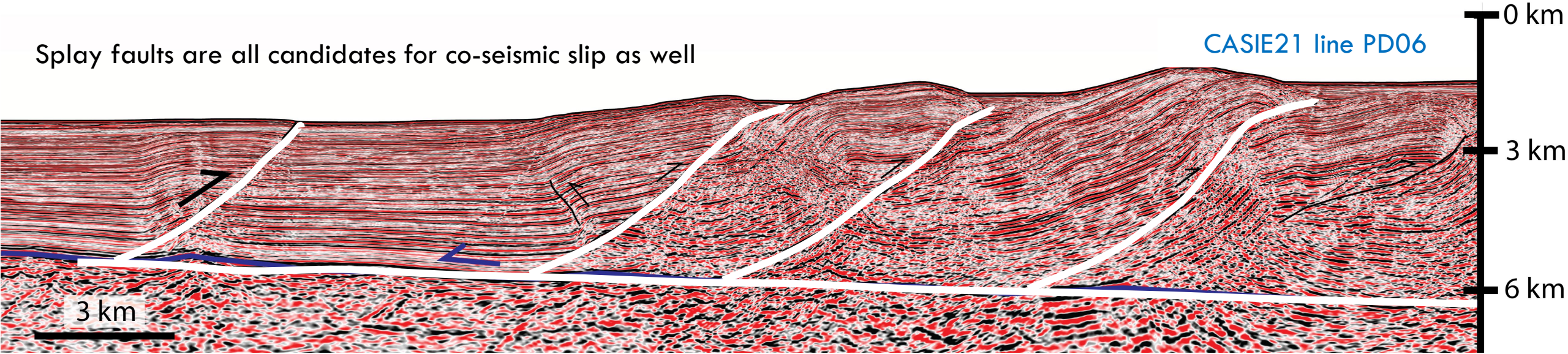
- *How far out to the deformation front does fast slip go on the megathrust décollement?*
- *What about shallow splay faults that reach the surface?*
- *Is there evidence of a mega-splay fault?*

Is there evidence of shallow coseismic slip?



Decollement is the plate boundary, and lies **~3 km** below the surface right at the deformation front

Splay faults are all candidates for co-seismic slip as well

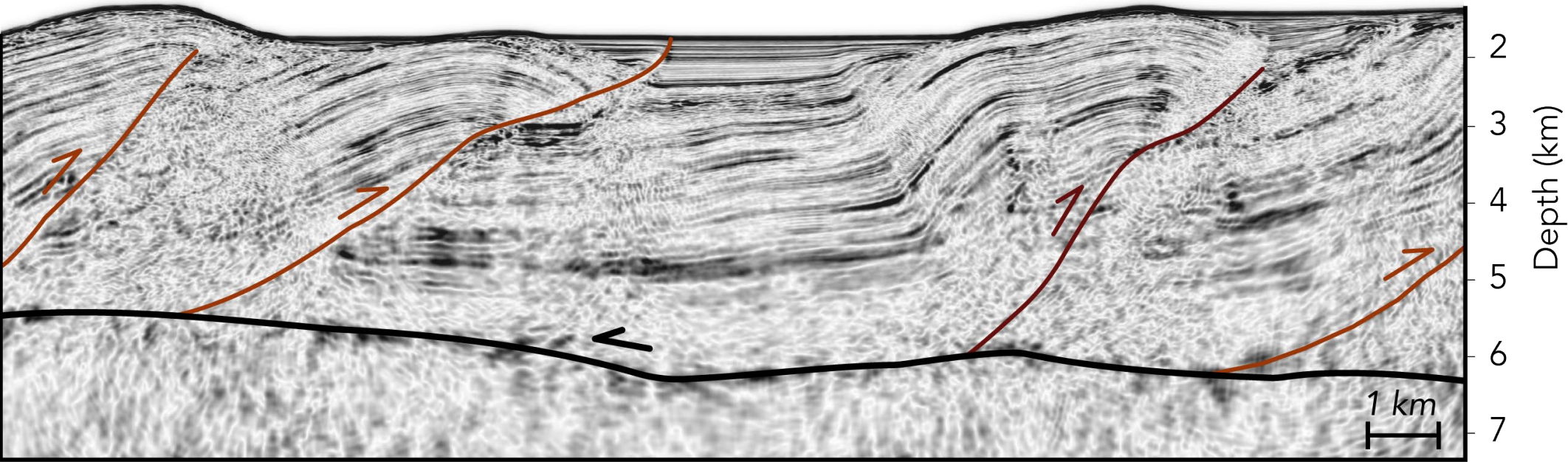


Anna Ledeczi: Joint interpretation of near surface (<1 km) and deeper (<10 km) imaging to map active faults



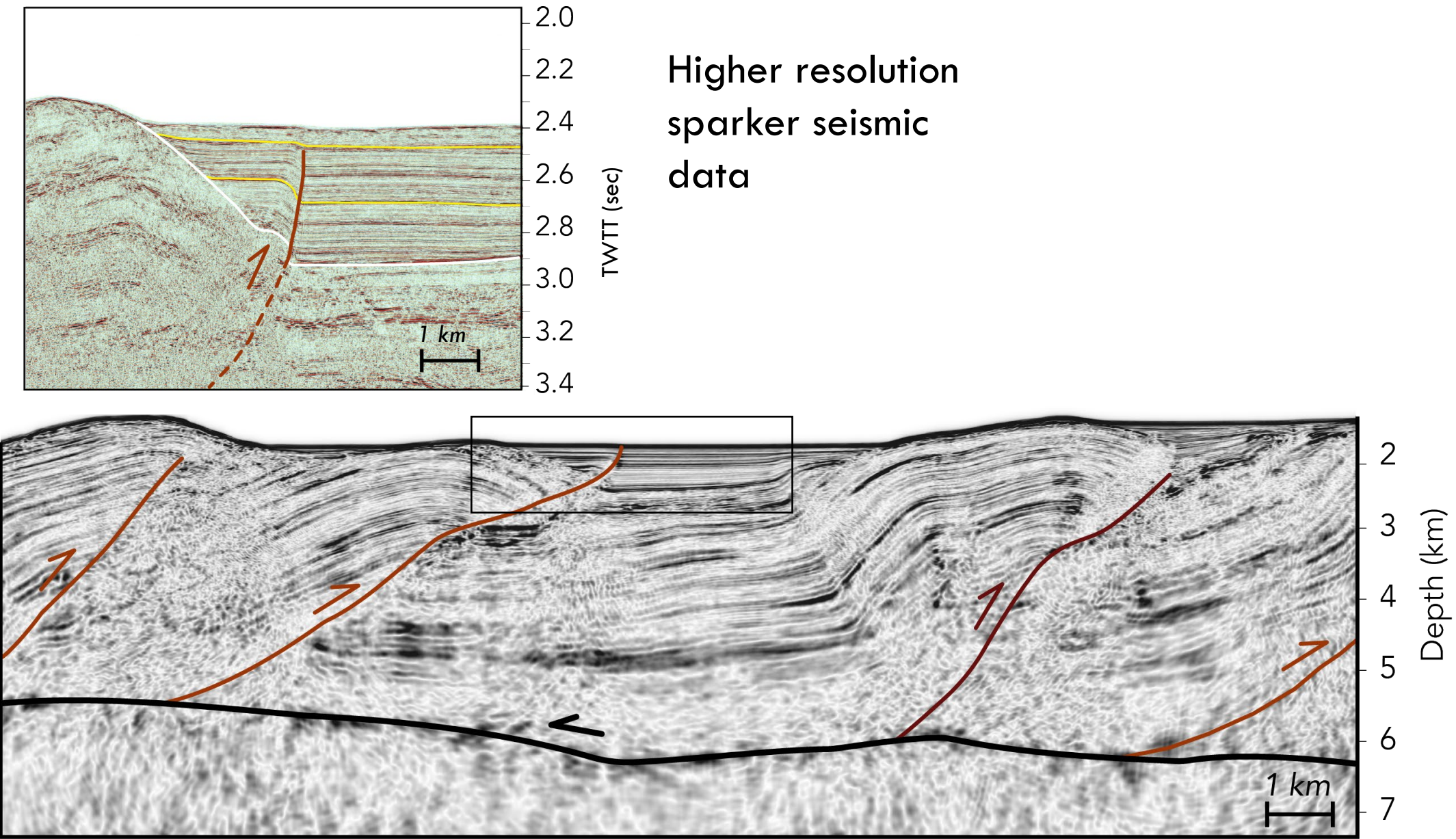
Anna
Ledeczi

CASIE21 PD06B



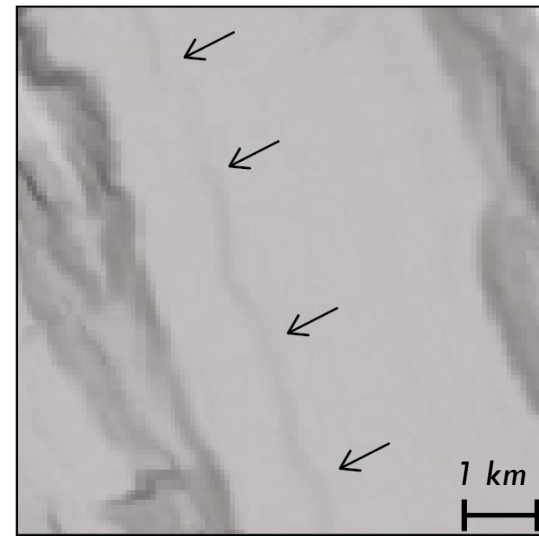
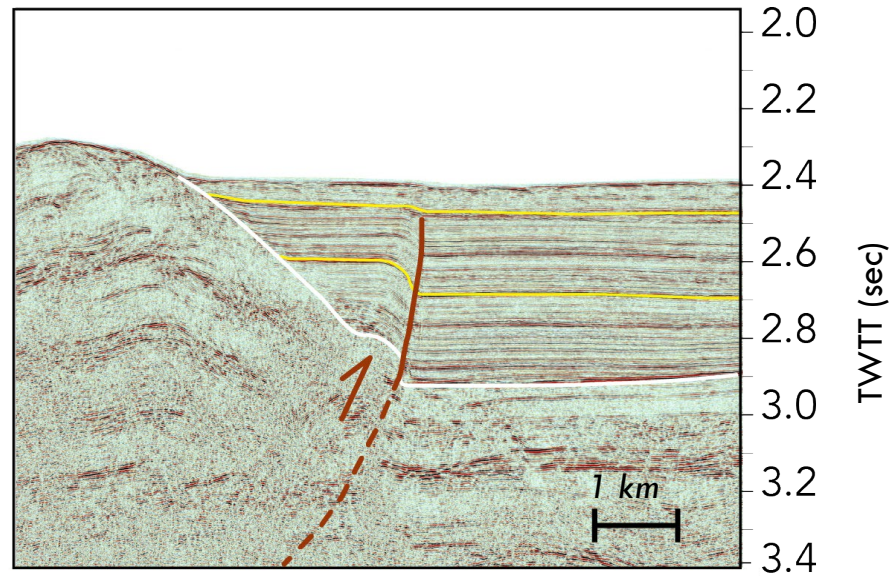
Ledeczi et al.,
2024

Joint interpretation of near surface (<1 km) and deeper (<10 km) seismic reflection imaging

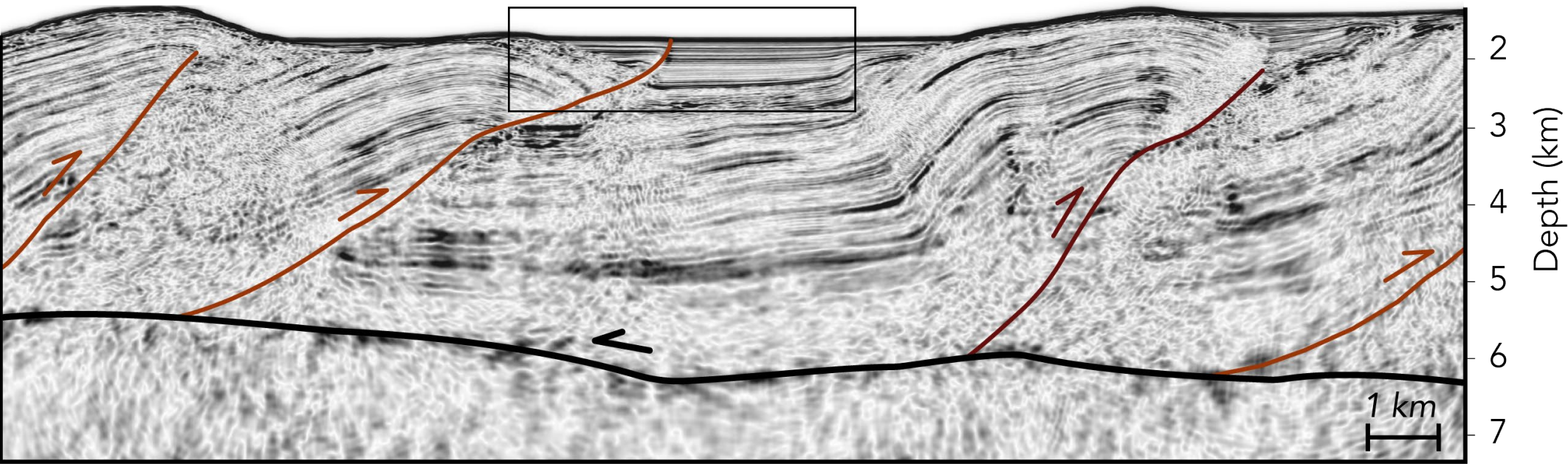


Ledeczi et al.,
2024

Joint interpretation of near surface (<1 km) and deeper (<10 km) seismic reflection imaging

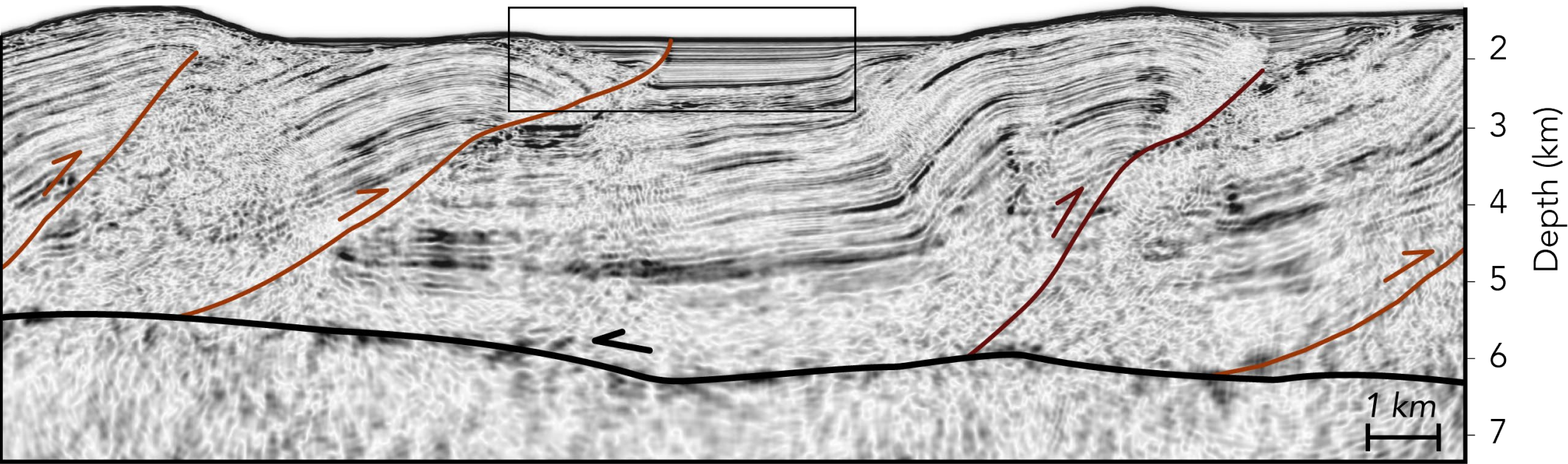
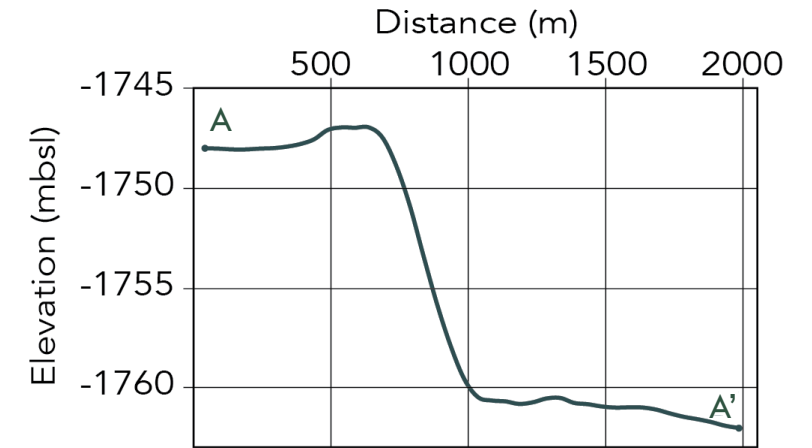
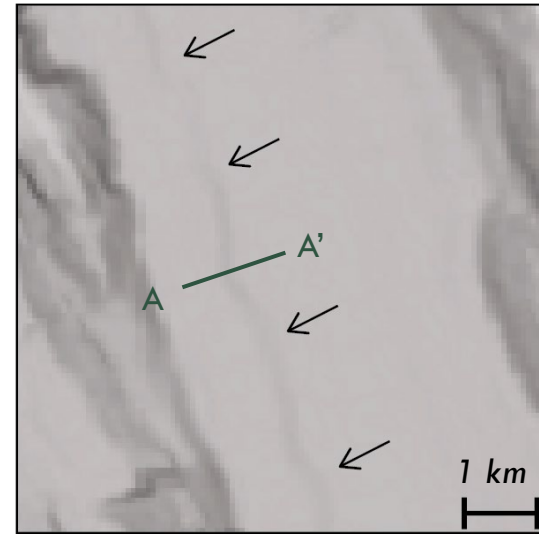
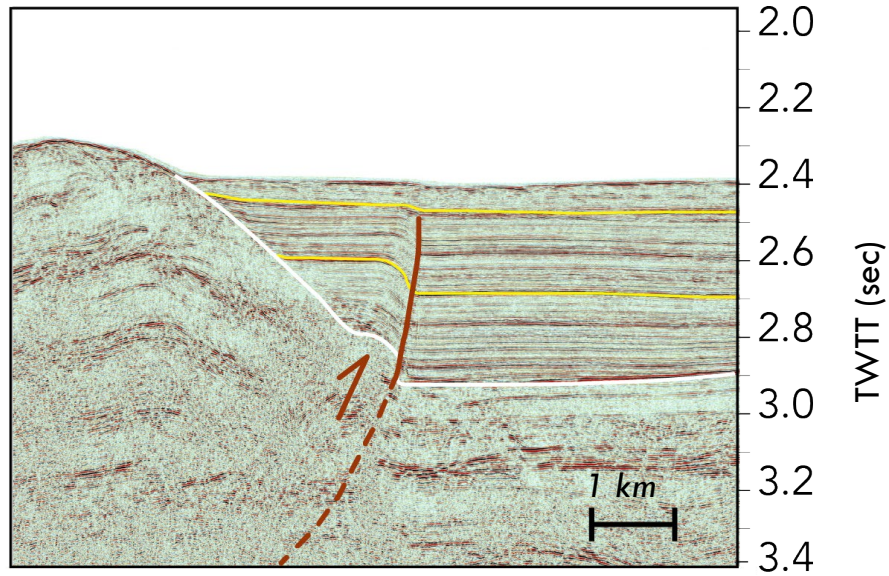


USGS 30-m
bathymetry



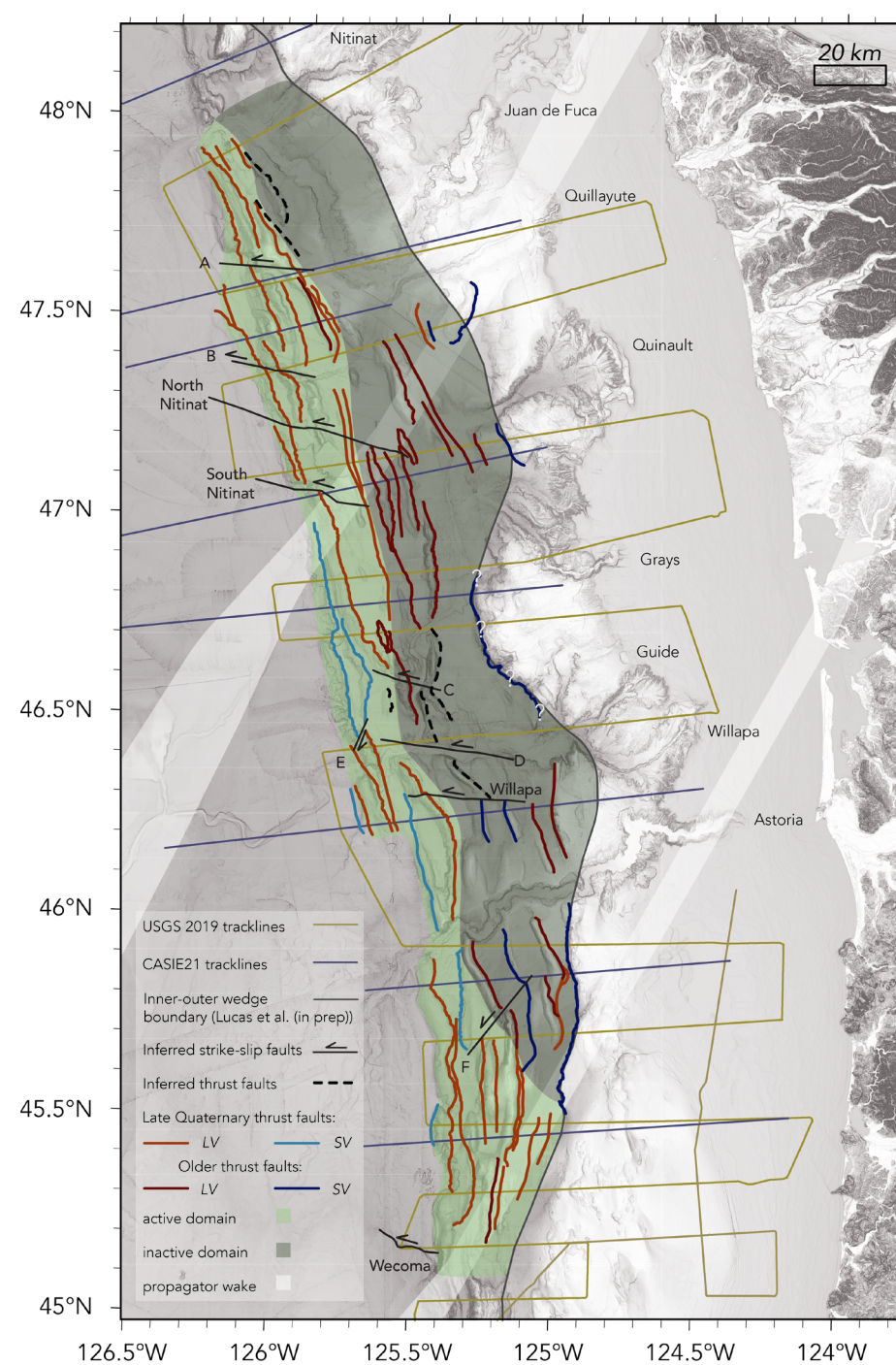
Ledeczi et al.,
2024

Joint interpretation of near surface (<1 km) and deeper (<10 km) seismic reflection imaging



13 meter scarp!

Ledeczi et al.,
2024

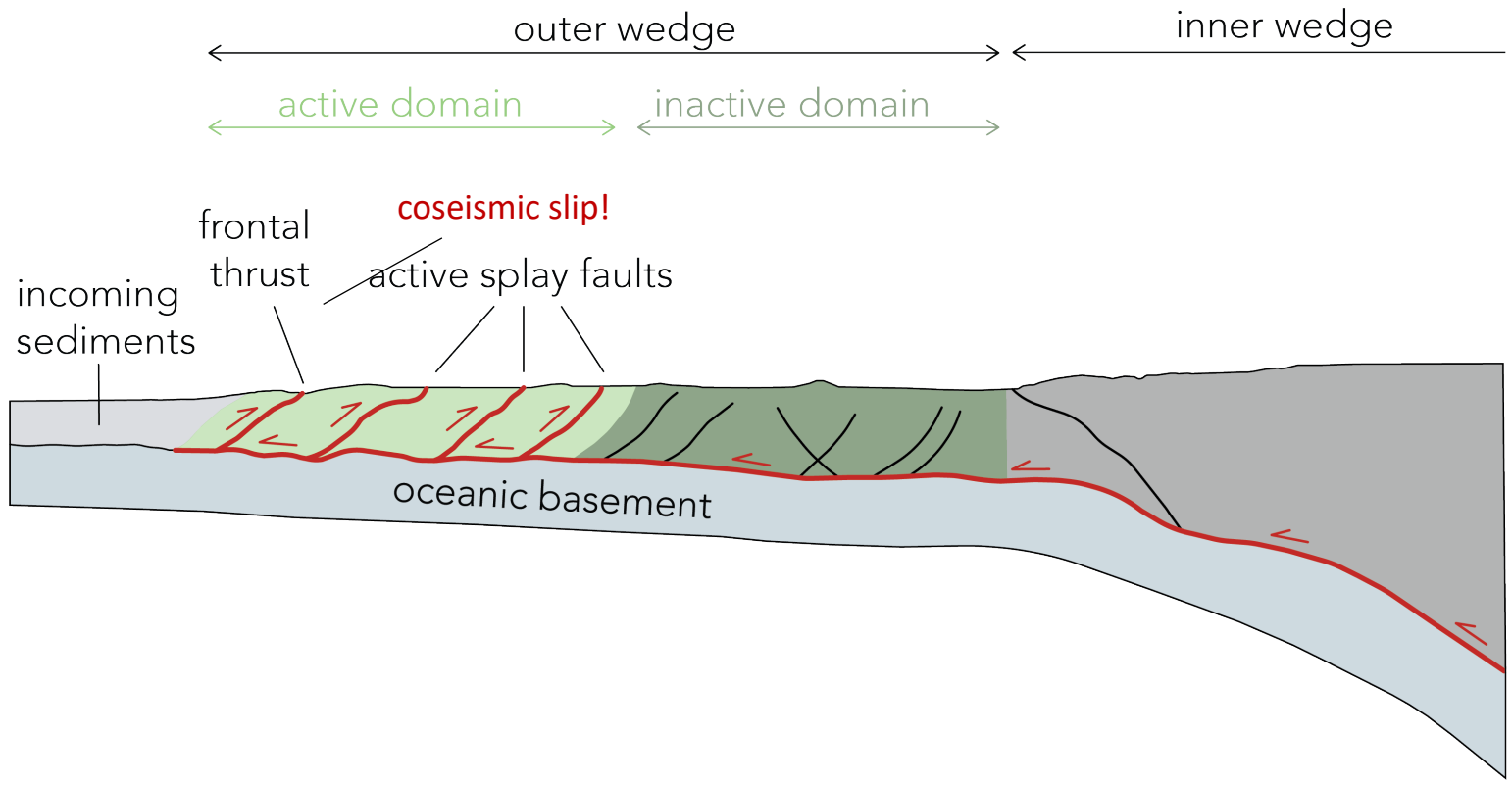
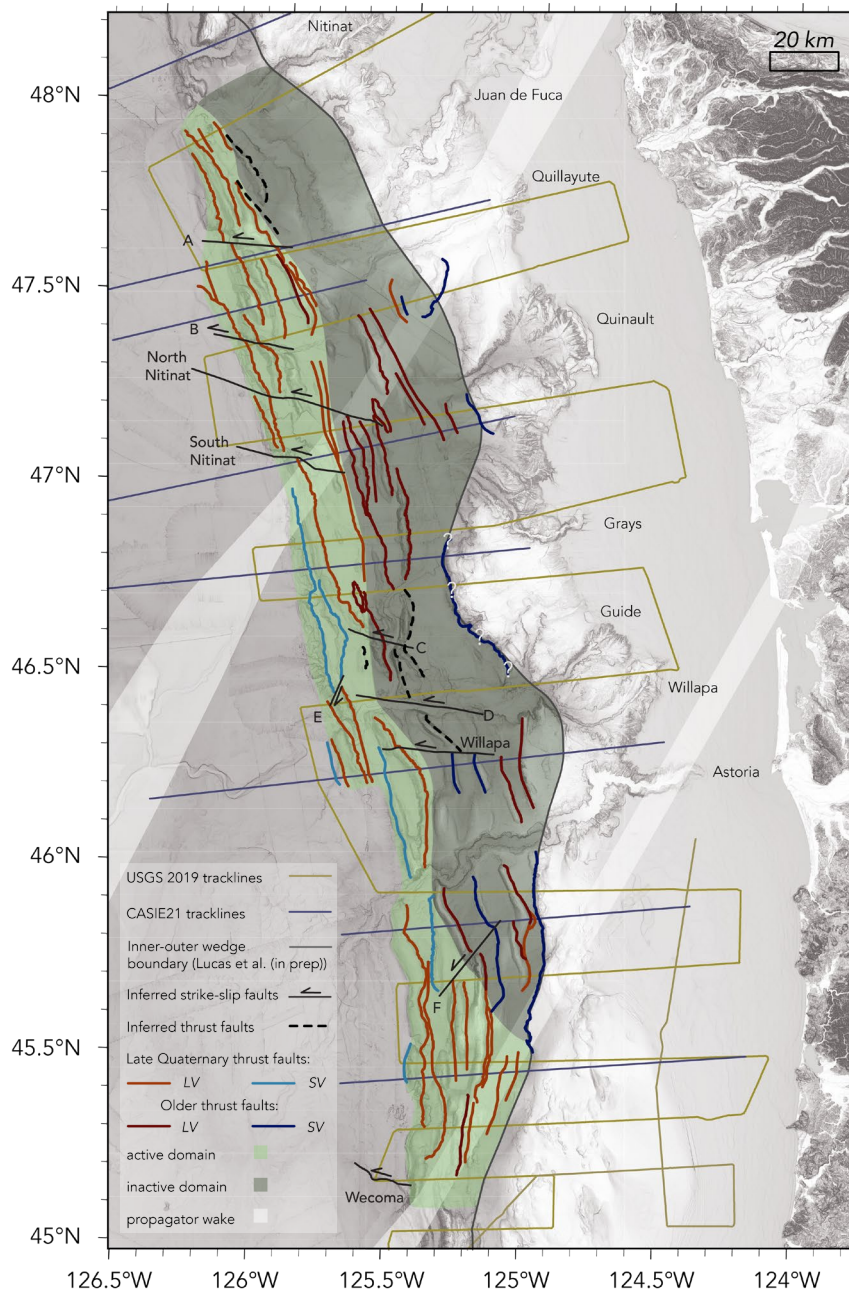


Widespread recent activity of splay faults within ~30 km of the deformation front

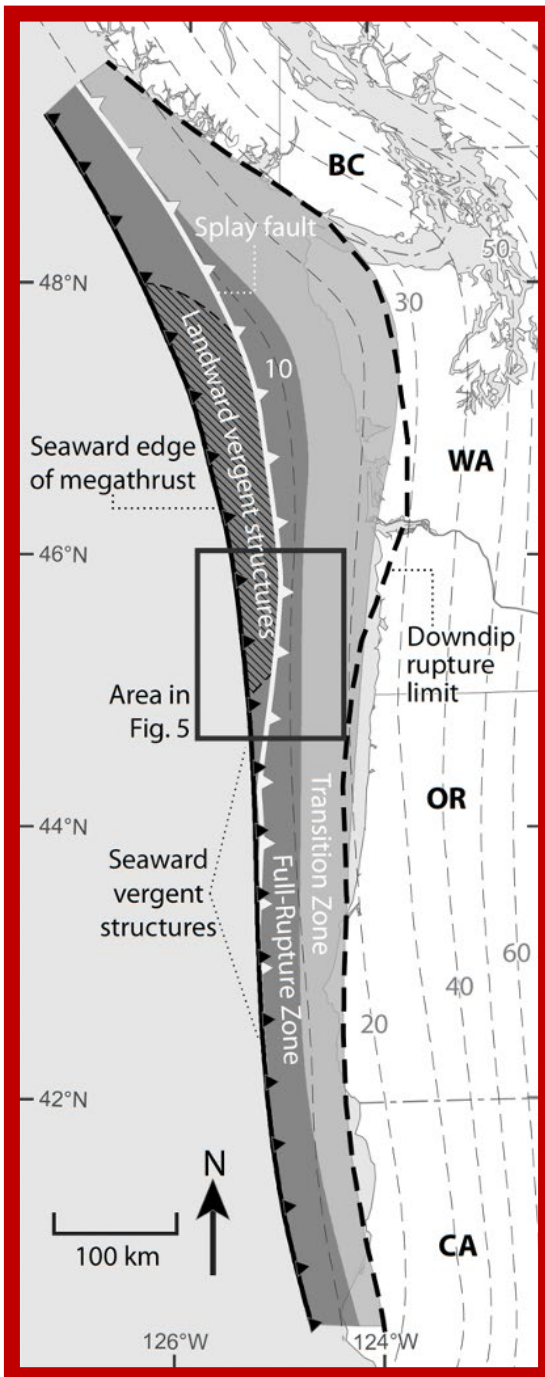
- New data resolution allows us to map out faults which are actually active
 - Previous work has identified similar faults, but all were called active
- Partitioning of recent activity into “active domain” and older activity into “inactive domain”
 - Active domain < 30 km wide
 - Inactive domain 10 to 40 km wide
- **Recently active = potentially active in next M9**

Faults in the active domain are candidates for recent (and therefore future) coseismic slip

- During megathrust events, shallow slip is likely and distributed onto multiple splay faults in the active domain
- The inner wedge and the inactive domain may move as a rigid block



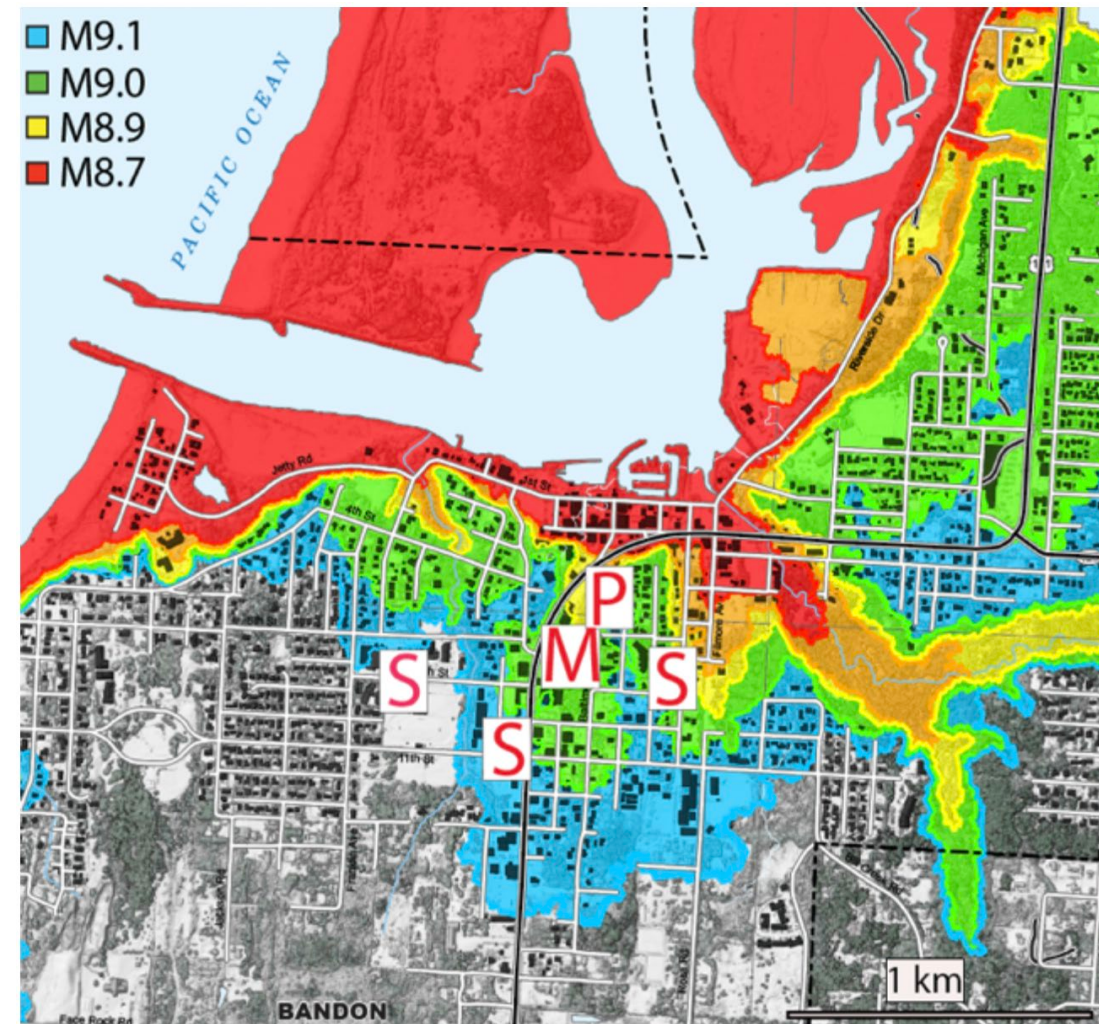
A megasplay fault at the inner wedge / outer wedge boundary has been proposed and is used as a preferred scenario in tsunami hazard models



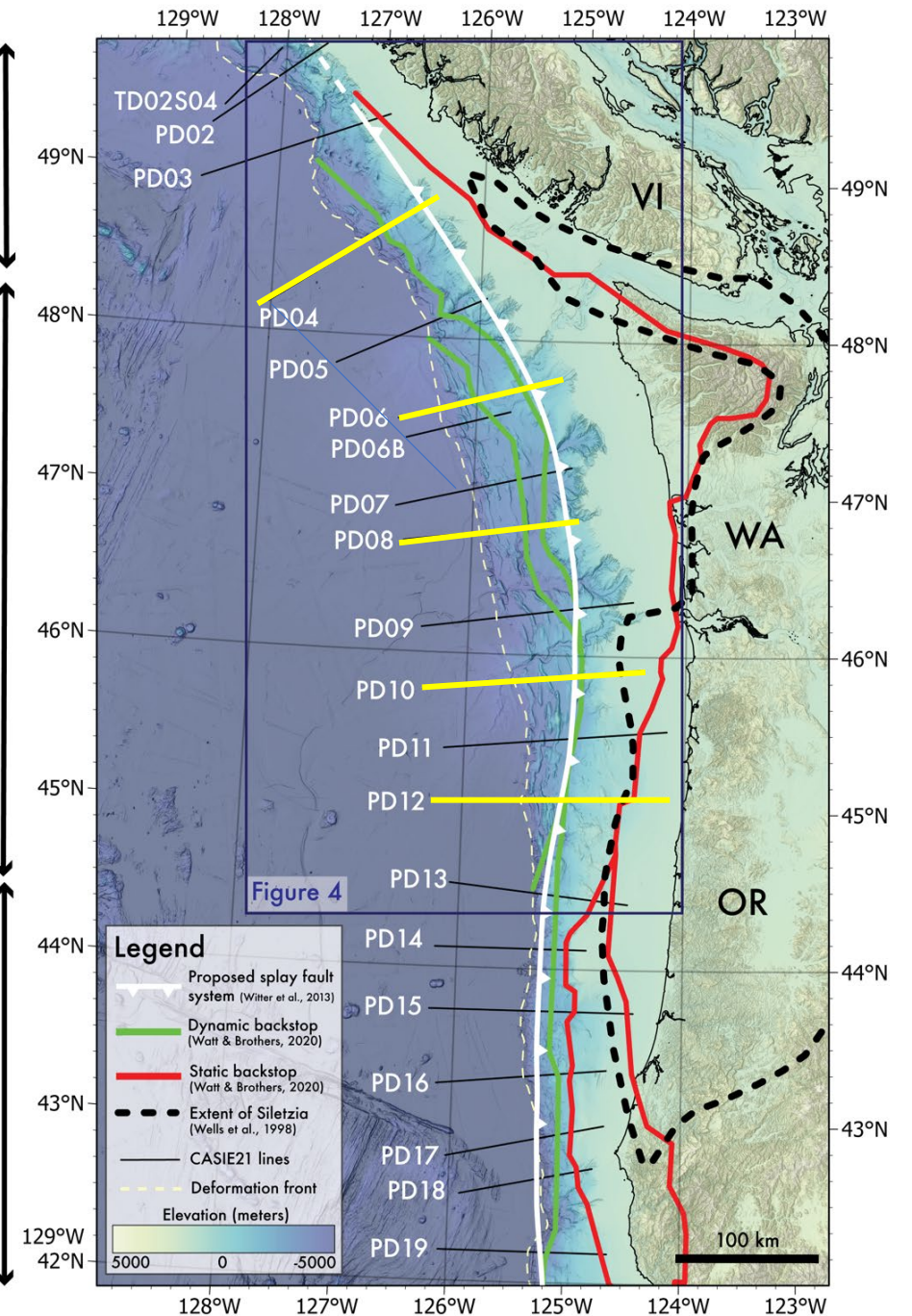
Official tsunami inundation scenarios for Bandon, Oregon ranging from M8.7 – M9.1

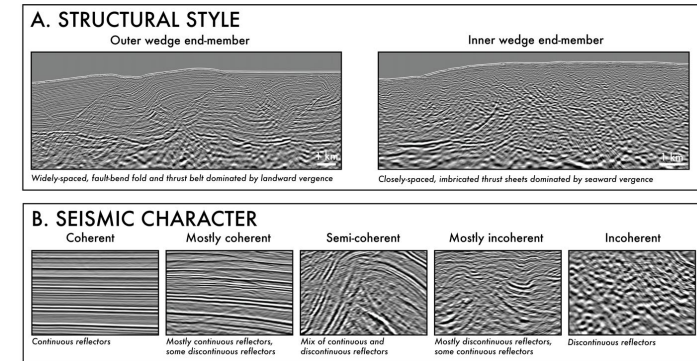
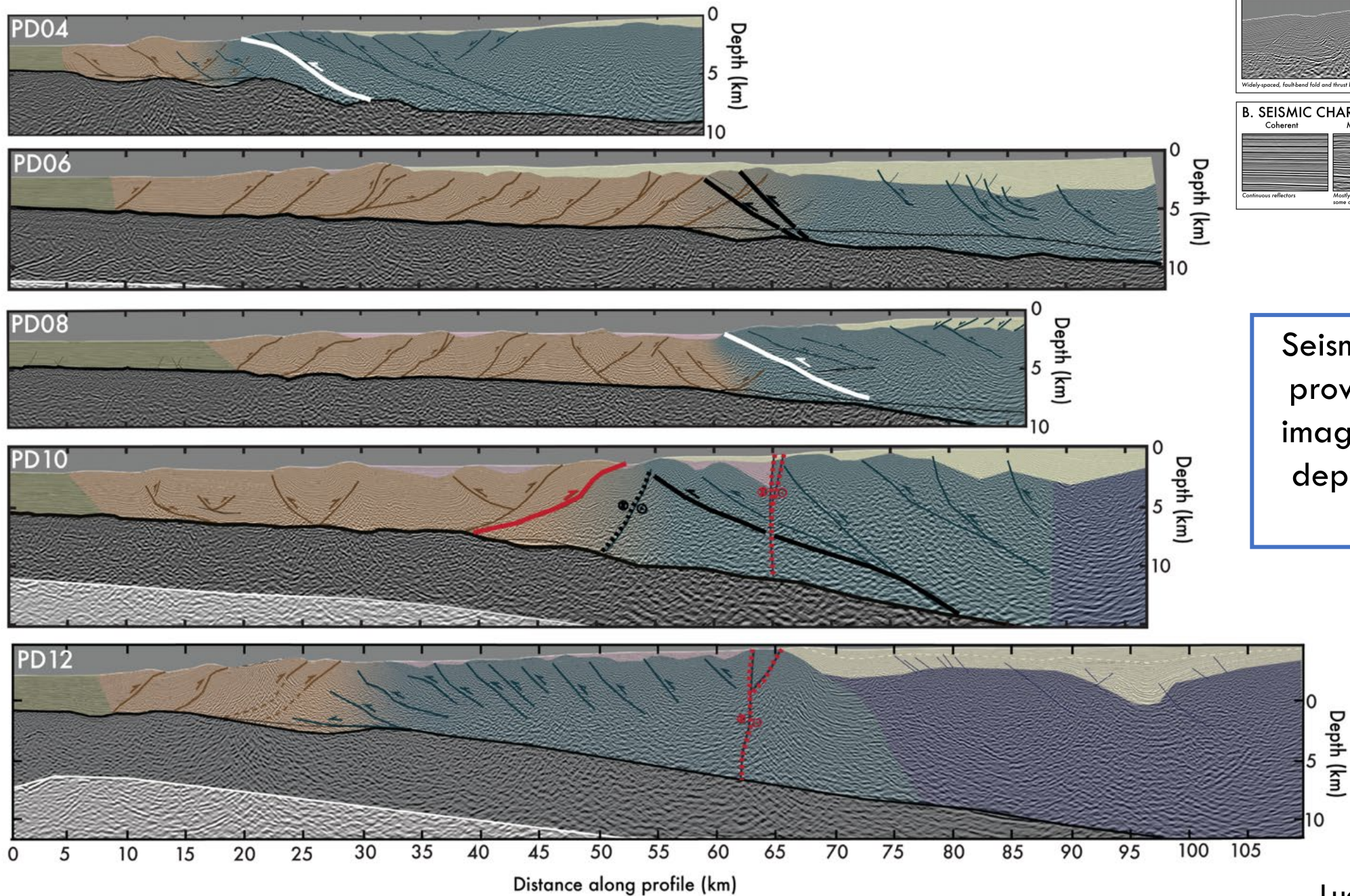
Uses models with megasplay rupture

Witter et al. (2013)



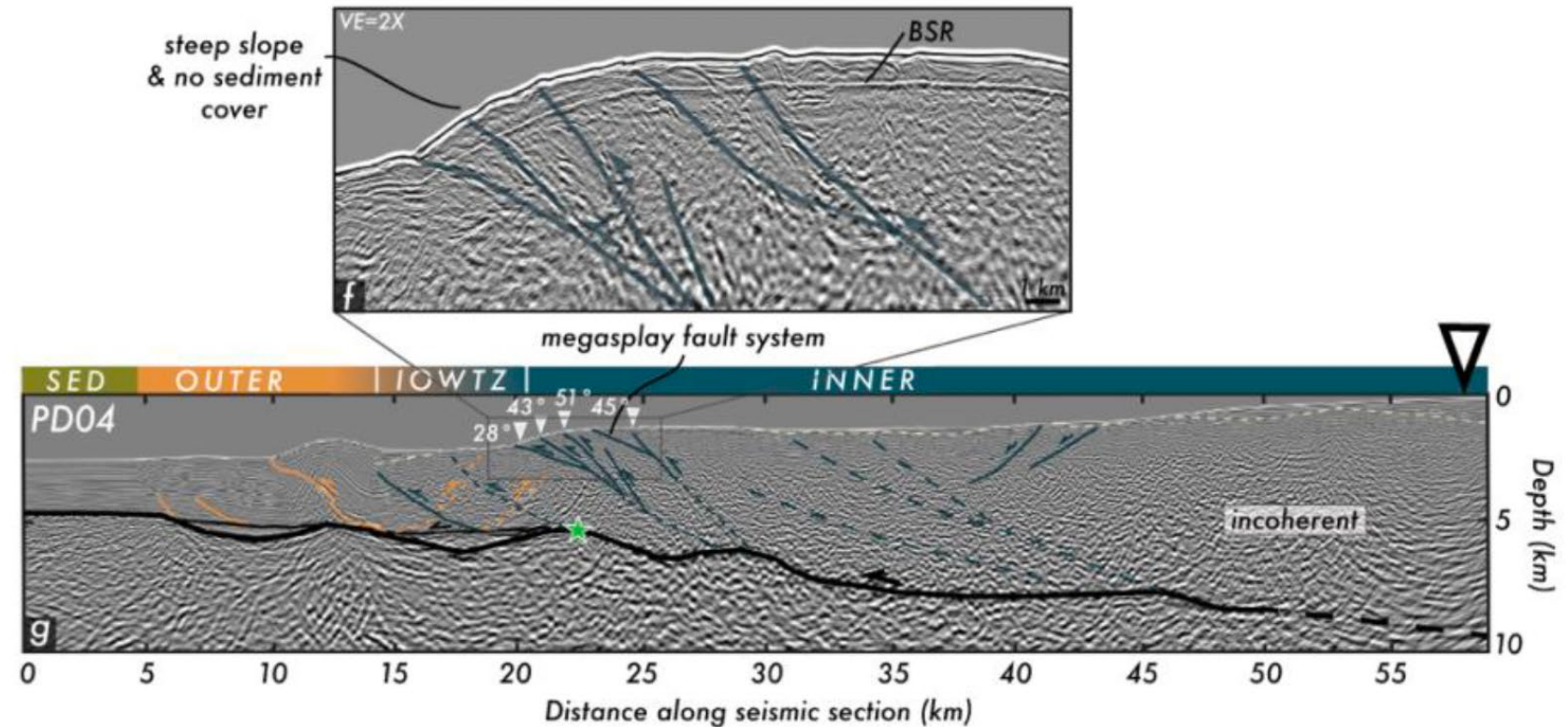
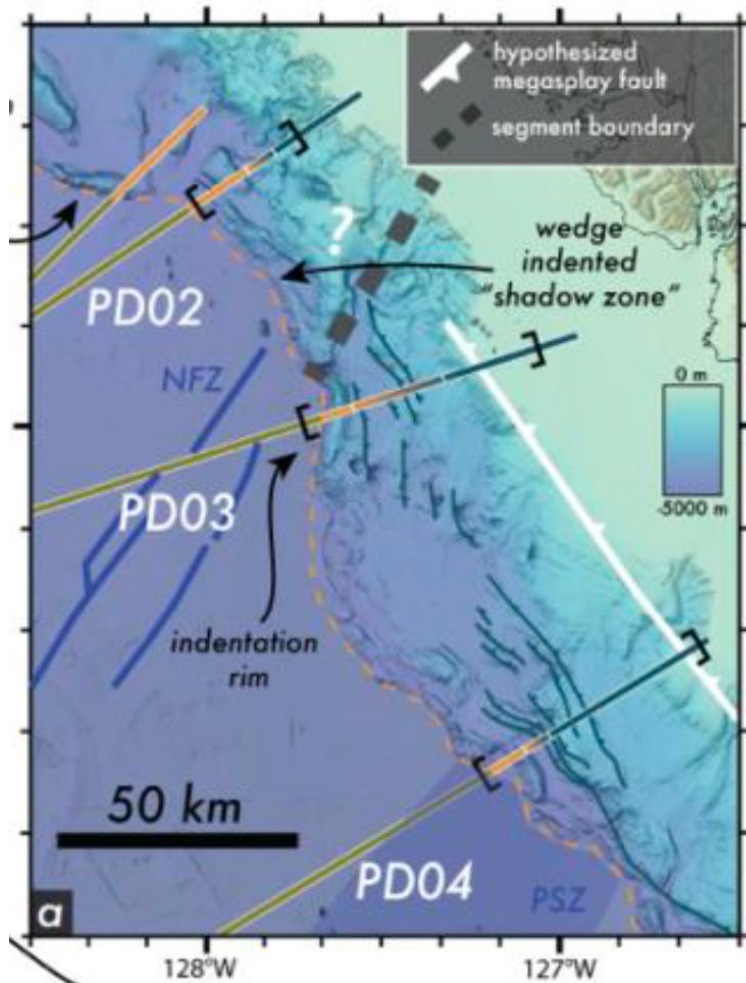
Lucas et al., 2025



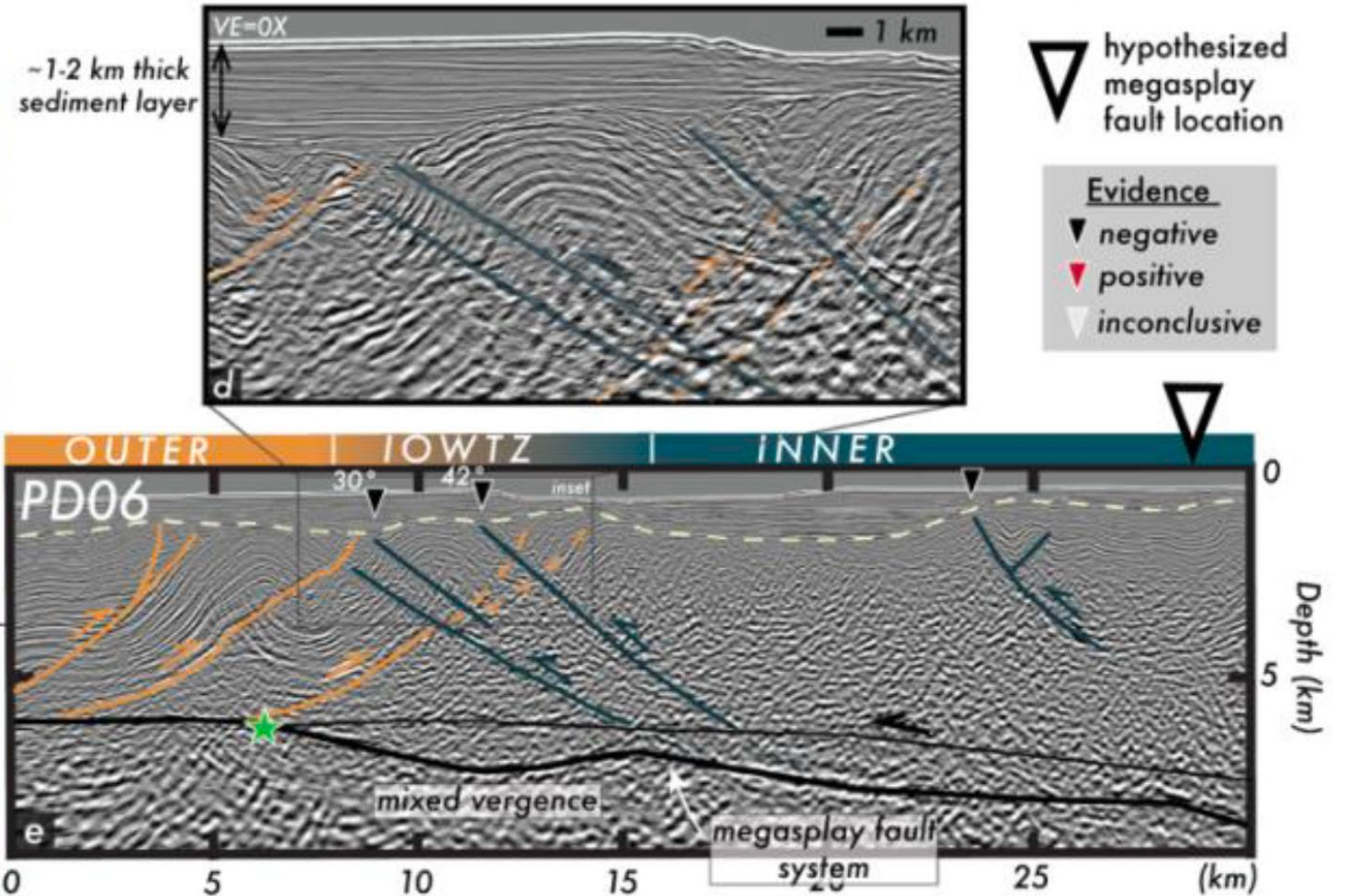
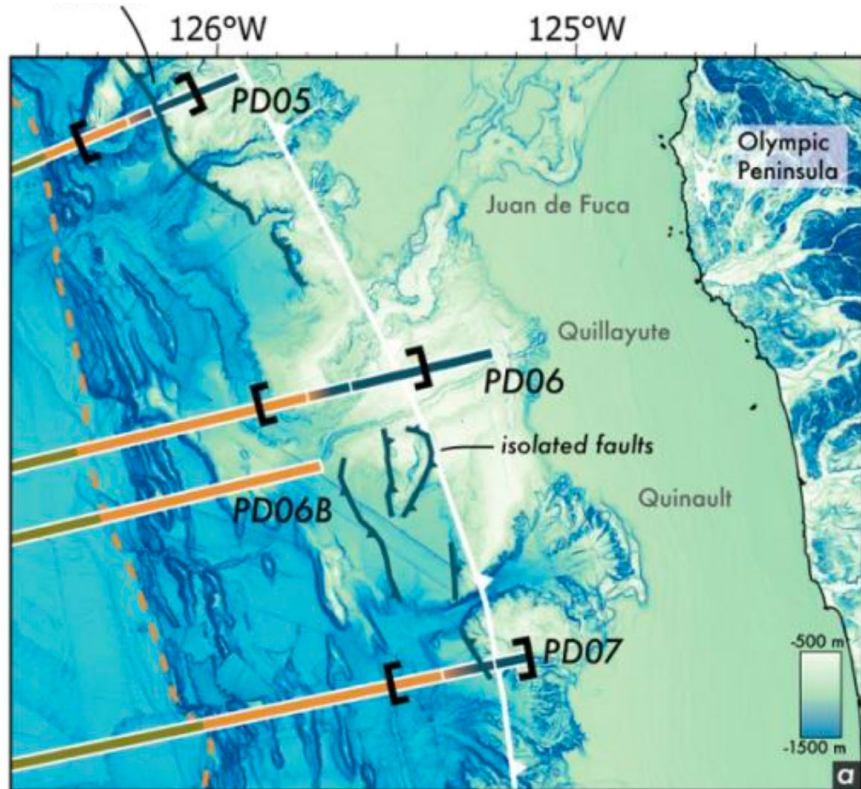


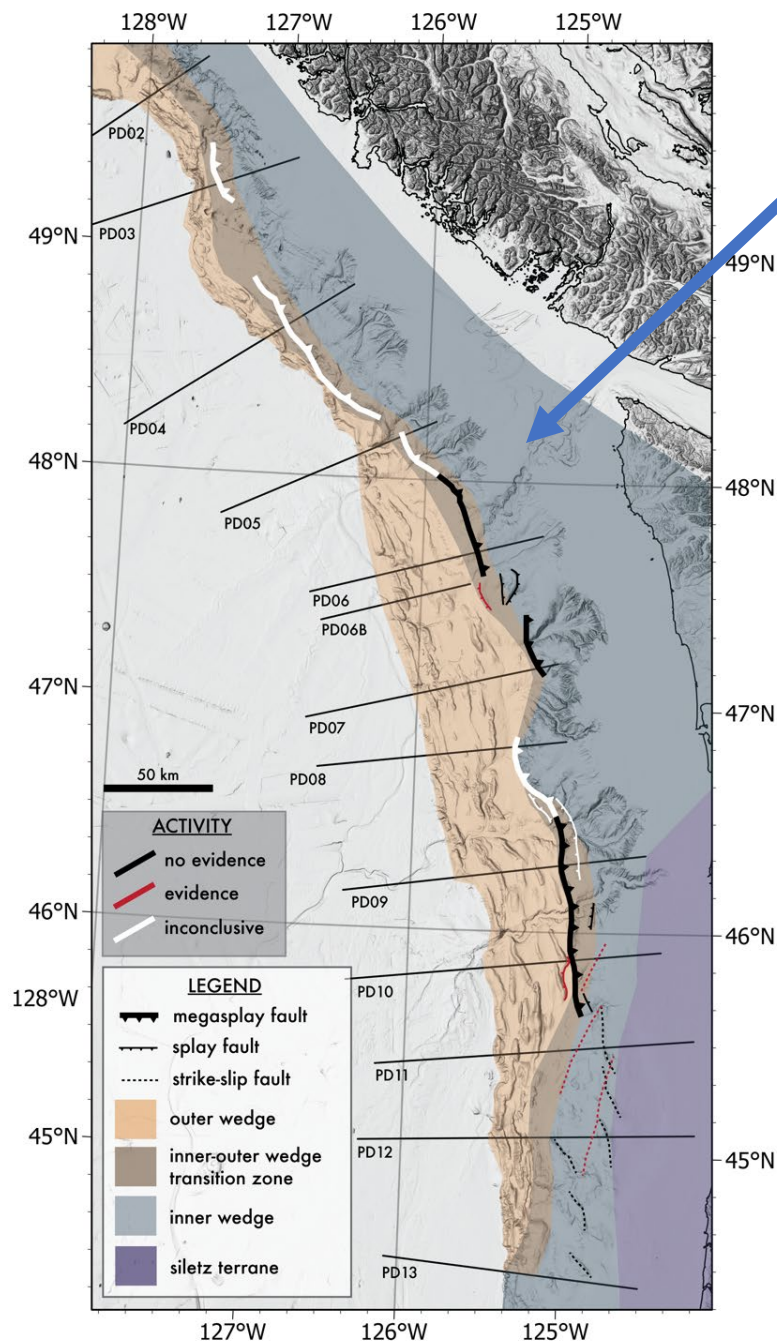
Seismic reflection profiles provide subsurface fault imaging down to ~15 km depth in unprecedented detail

Off Vancouver Island, some areas show potentially active megasplay fault system



Off Washington & N. Oregon, evidence is clear that there is no such active fault in most areas





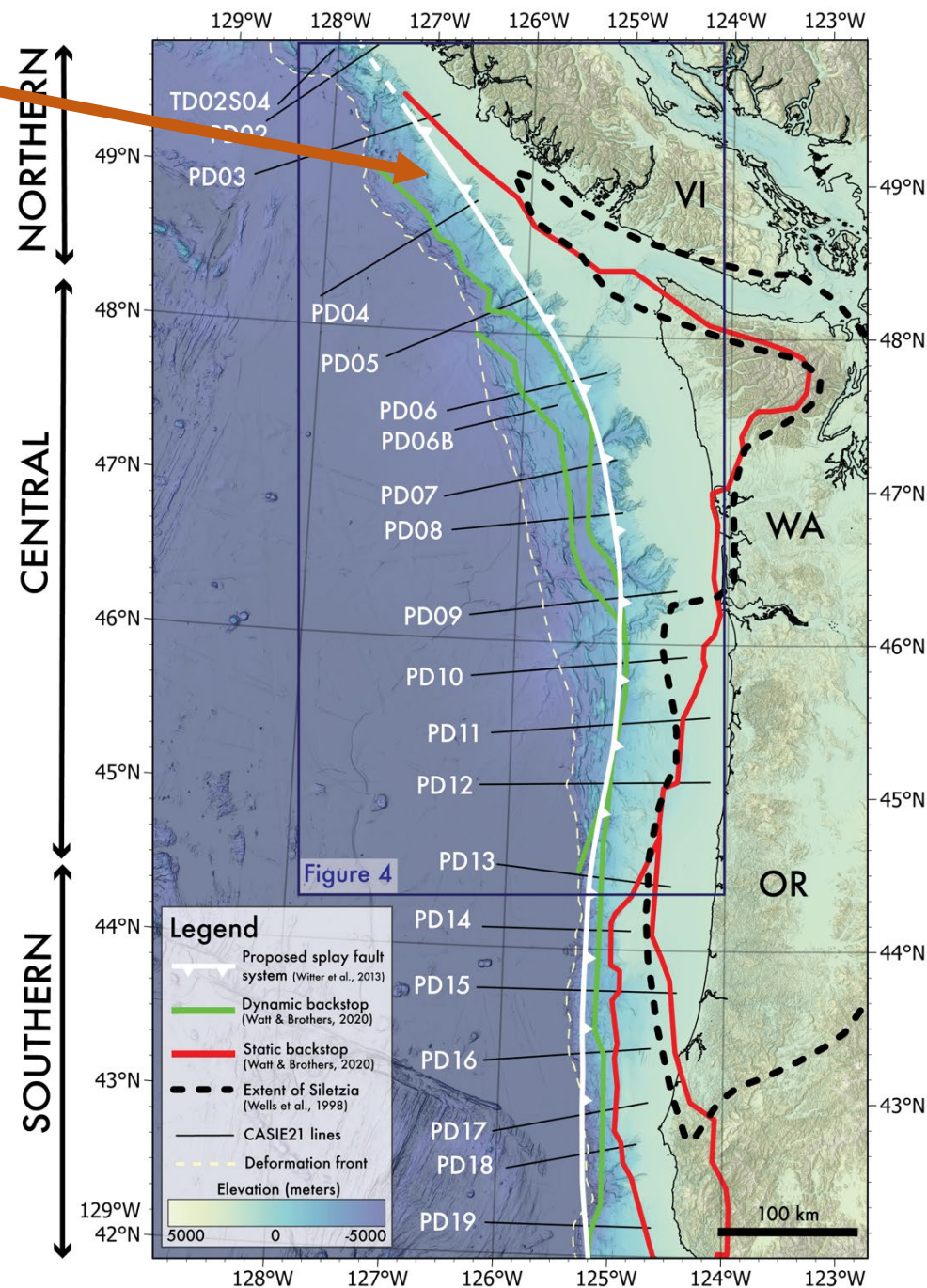
YES

NO

Bottom line:

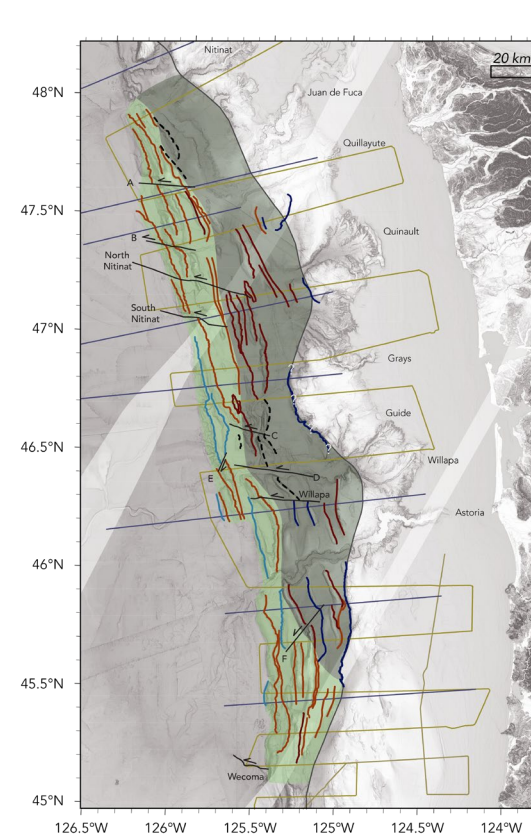
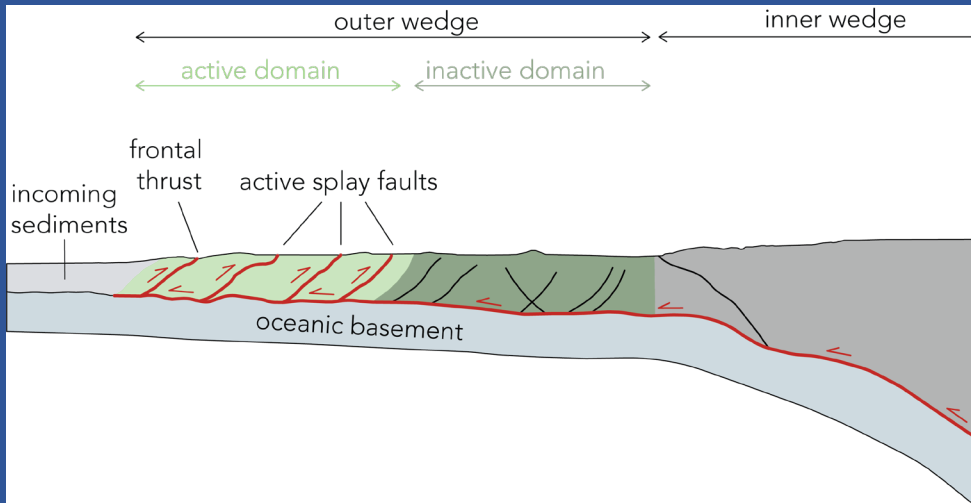
Strong evidence that there is **no through-going megasplay fault** off Washington or Oregon

Lucas et al., 2025



Seismic slip distribution offshore

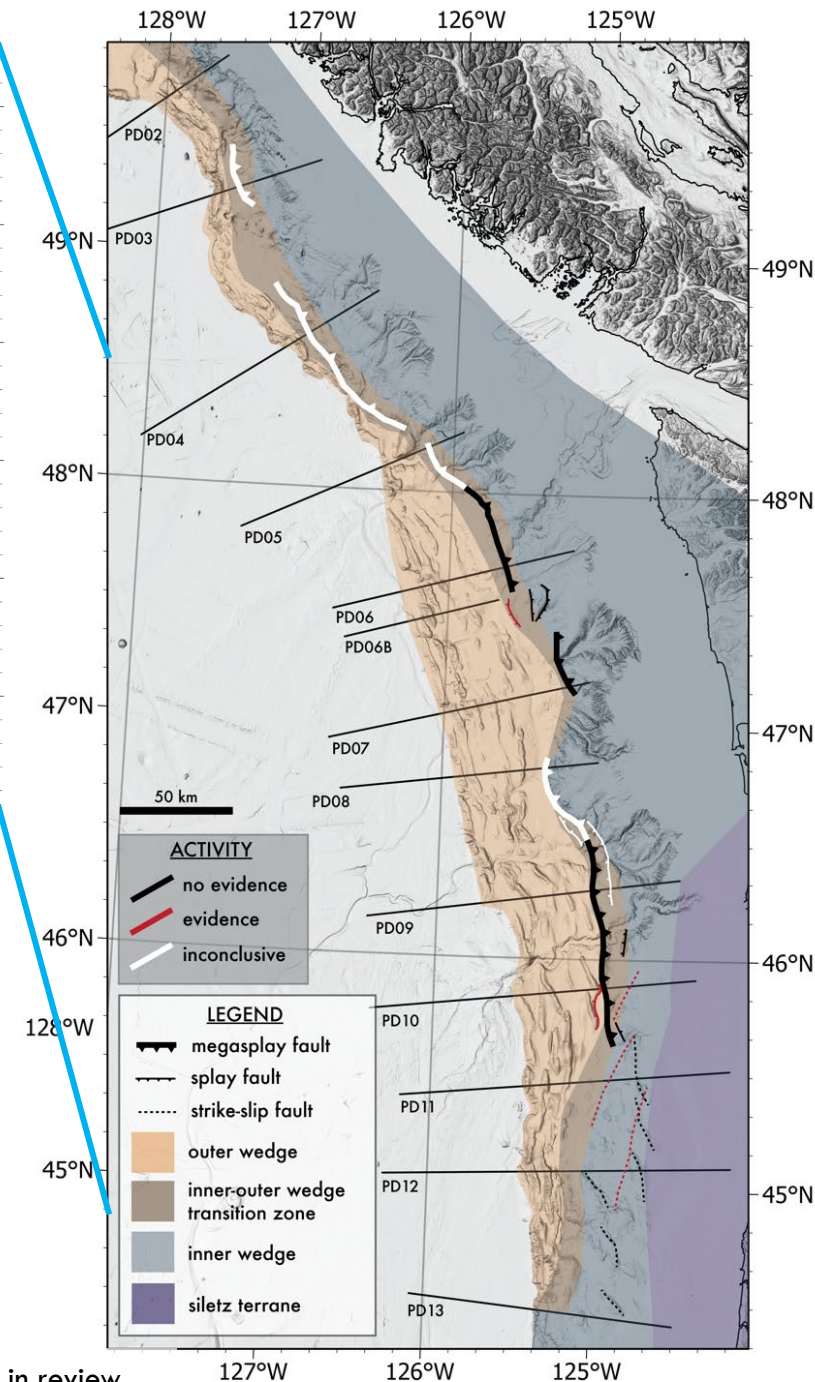
- Slip of the locked megathrust under the entire LVZ to the def. front is *likely*
- There is patchy evidence at best for a major megasplay fault – commonly-used tsunami scenarios need re-evaluation
- Slip *is* likely to occur on multiple LV splay faults in the outermost ~ 30 km, which may add to seafloor displacement



Ledeczi et al., 2024

Many recently-active faults in outermost region (light green)

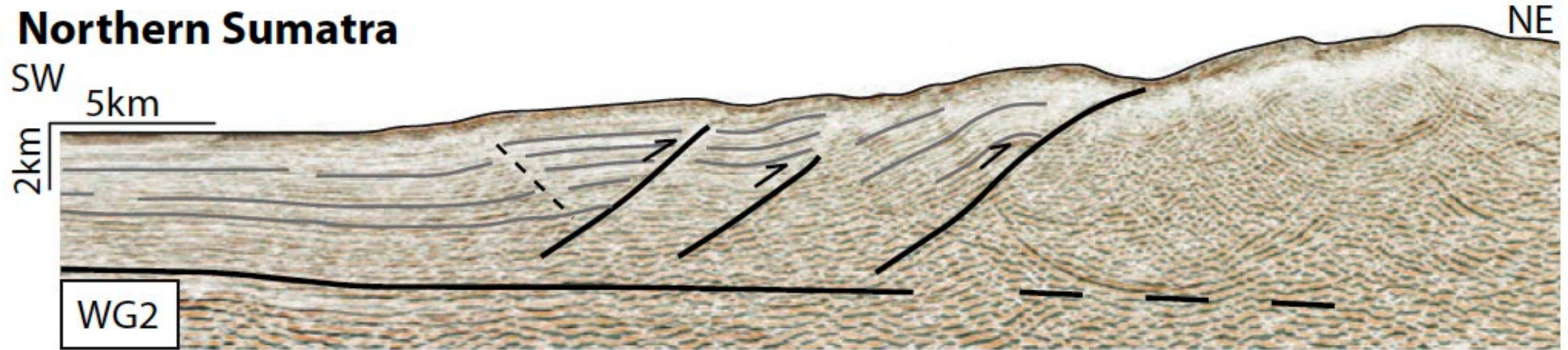
“Megasplay” fault is discontinuous to absent



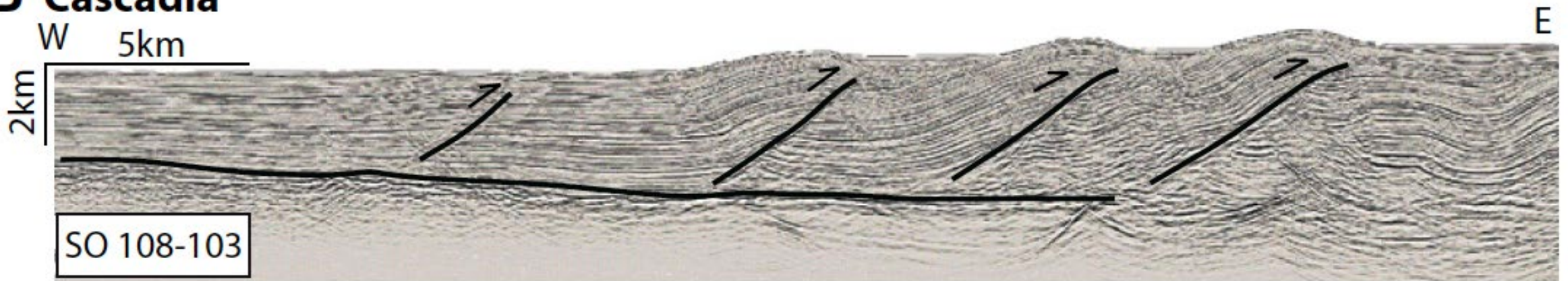
Lucas et al., in review

Is the landward-vergent Cascadia wedge direct evidence of shallow co-seismic slip?

A Northern Sumatra



B Cascadia



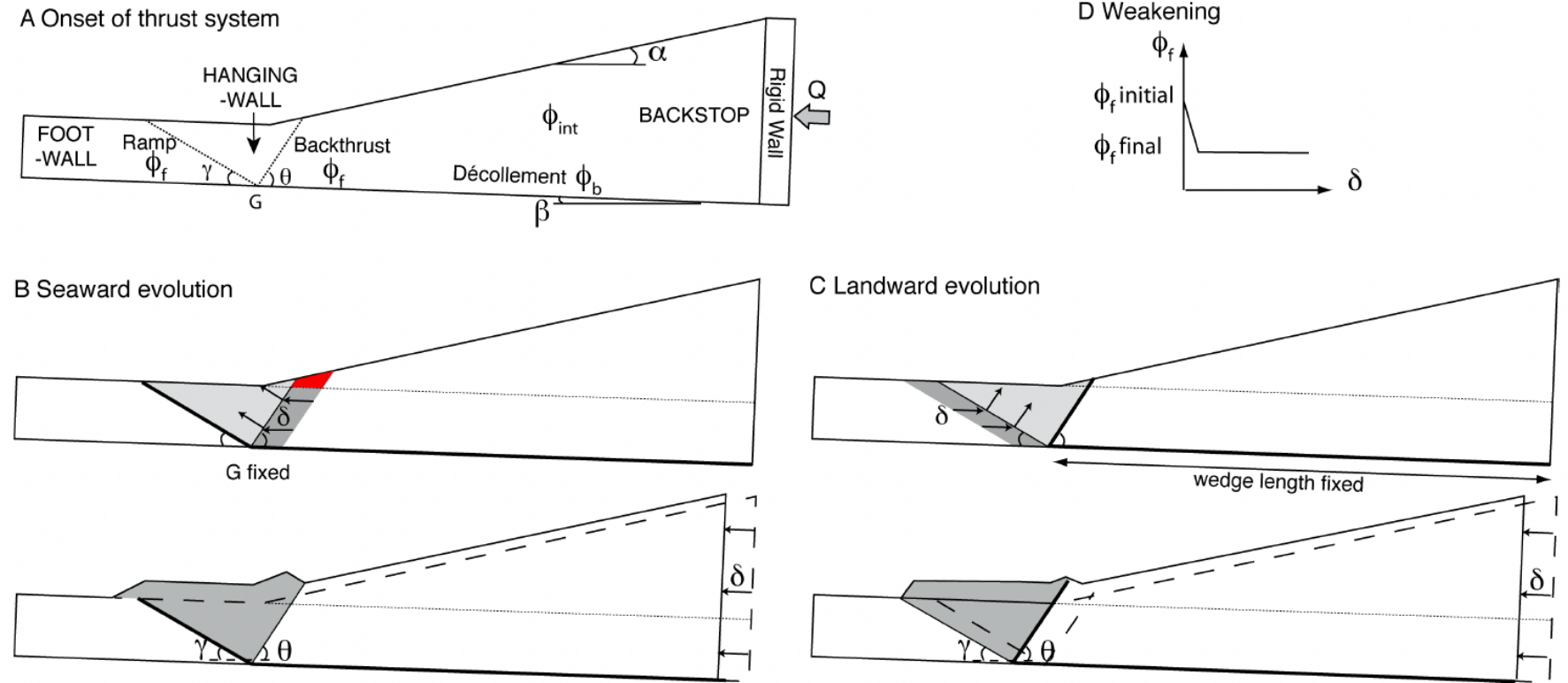
Development of landward vergent thrusts in Coulomb wedge requires extremely low coefficient of friction.

Representative values for the LVZ from CASIE21 lines:

surface slope $\alpha = 1.0^\circ$

megathrust dip $\beta = 2.0^\circ$

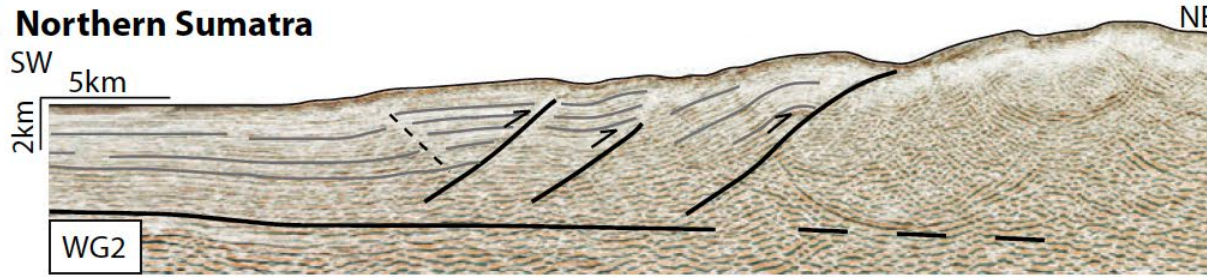
LV fault dip $\theta = 40^\circ$



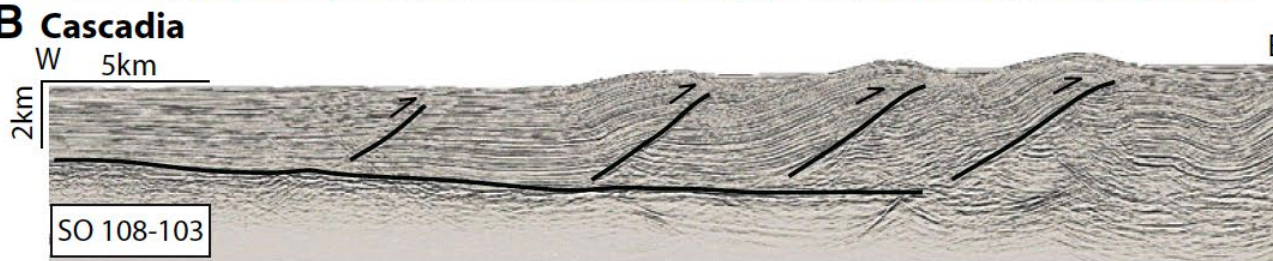
Mechanical analysis by Cubas et al. (2016 in *Geology*)

Cascadia case: $\mu_{\text{int}} = 0.57 \rightarrow$ LV only permitted when when $\mu_b = 0.01$ to 0.03
 For $\alpha = \sim 2^\circ$ and $\beta = 1.1^\circ$

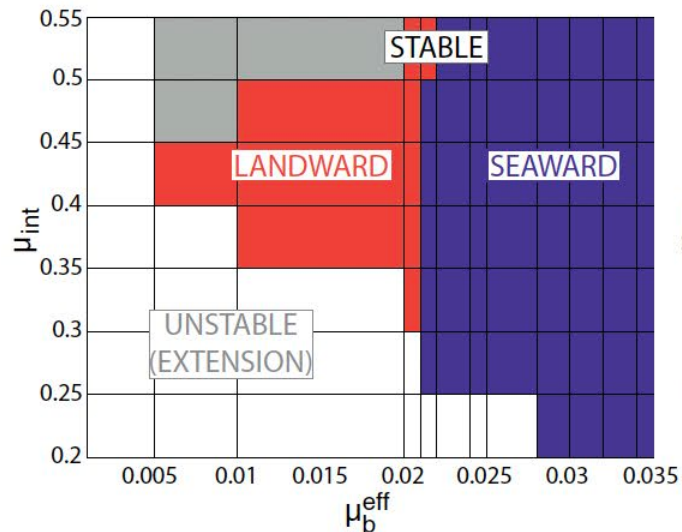
A Northern Sumatra



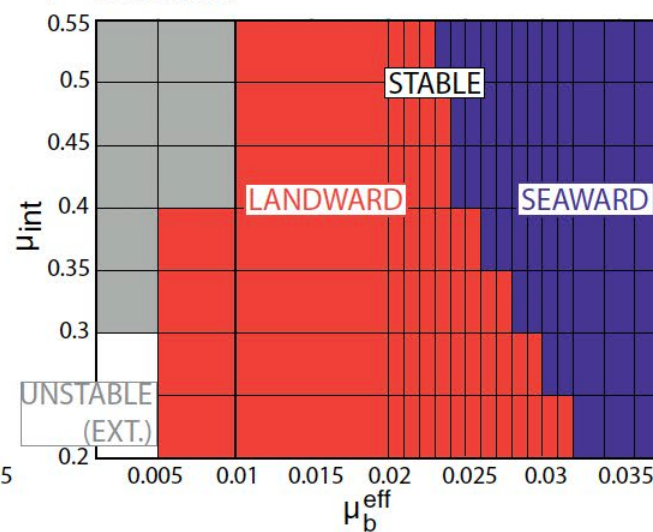
B Cascadia



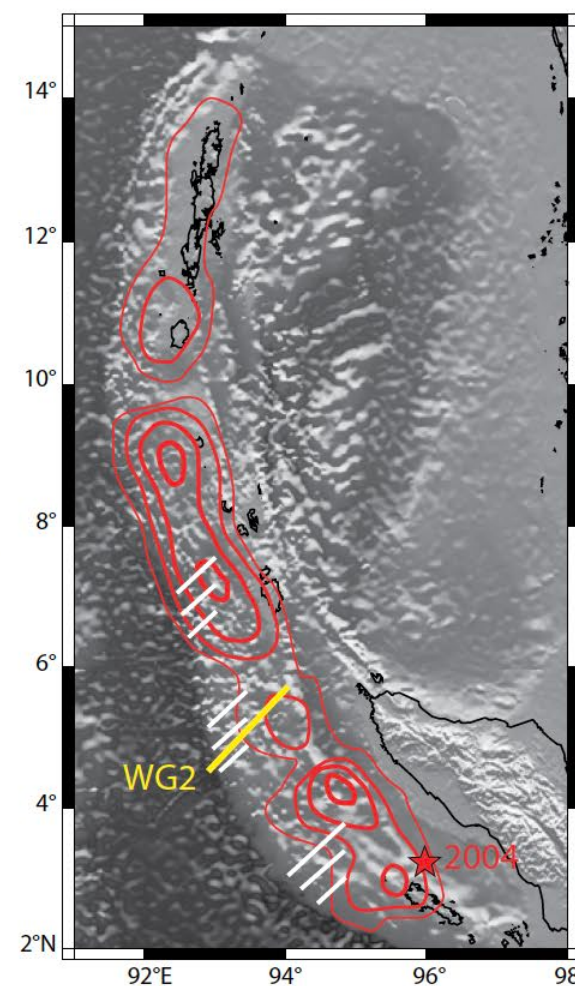
E Northern Sumatra



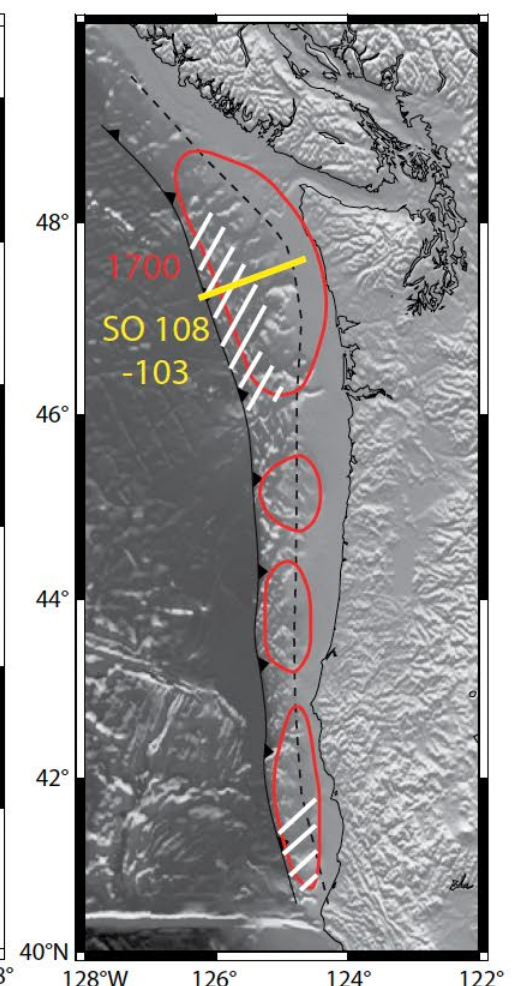
F Cascadia



C Northern Sumatra



D Cascadia

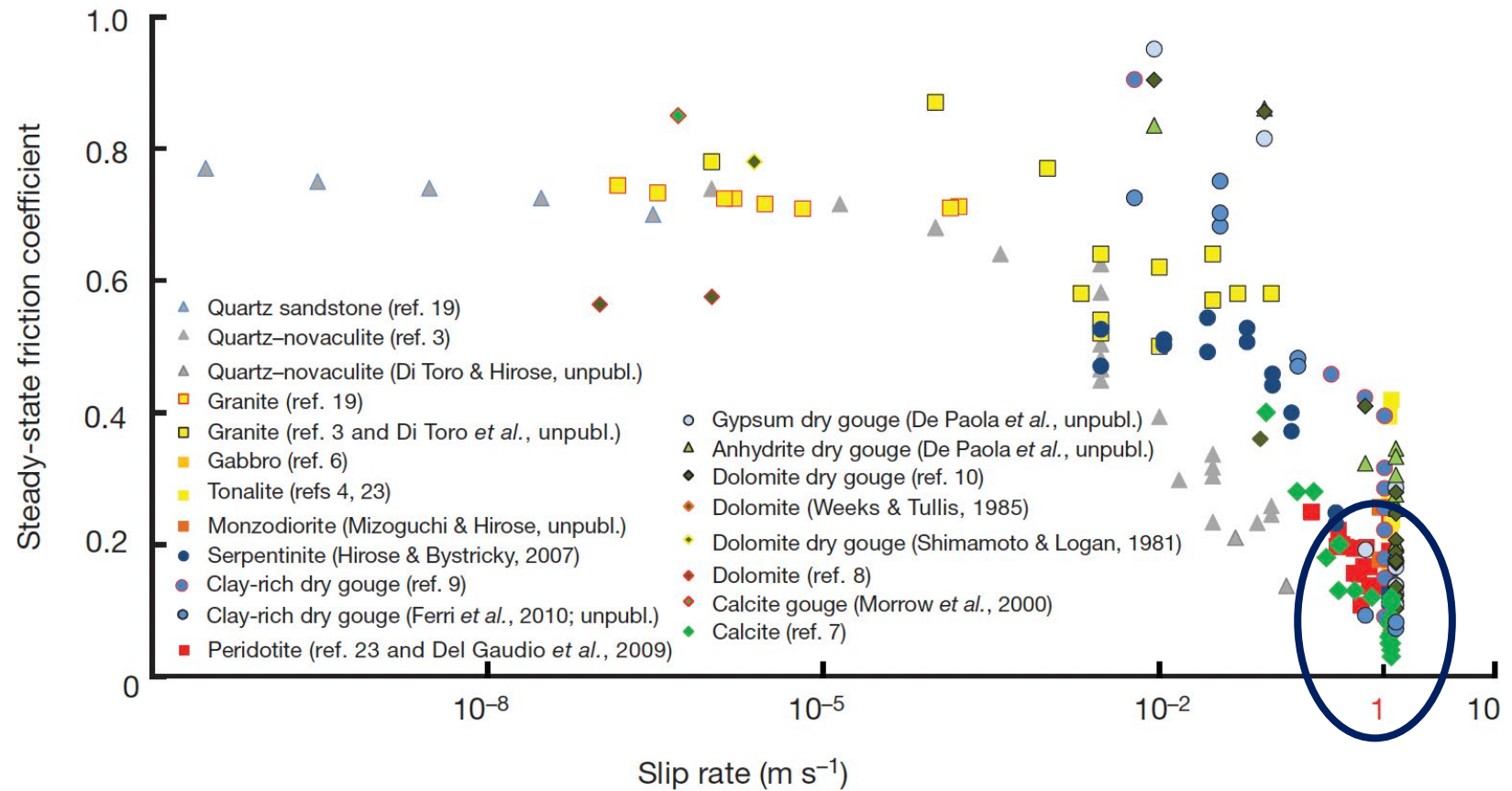


Cubas et al., 2016

Nothing has intrinsic friction that low

- No intrinsically low-friction viable candidate materials
- No evidence for extant high pore pressure ... and that promotes aseismic creep, not locking
- → *Dynamic weakening is the only viable mechanism*

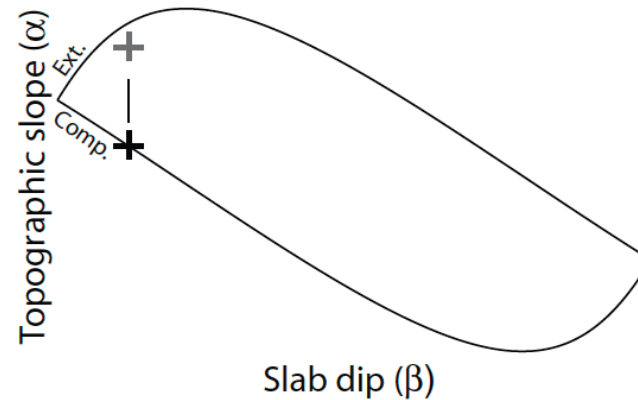
For the landward vergent zone to form, locking and dynamic weakening is not only permitted but required.



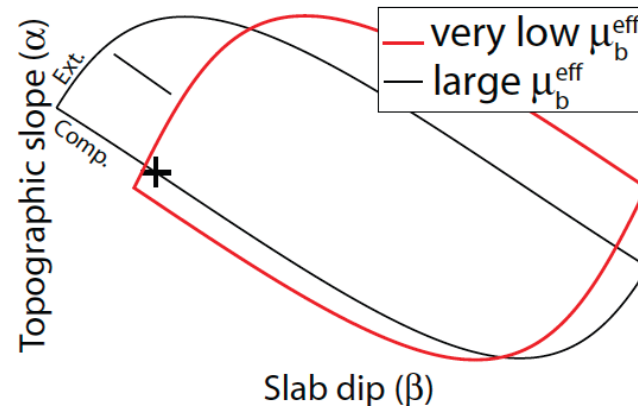
Landward vergent zone is **statically strong** but **dynamically weak**

- ❖ slip on landward vergent thrusts (backthrusts) is favored *only* for basal friction $\mu^{\text{eff}} \leq 0.03$
- ❖ requires a strong wedge internal friction of ~ 0.5
- ❖ likely develops only via dynamic weakening mechanisms because materials with such low μ_b (e.g., talc) are not present

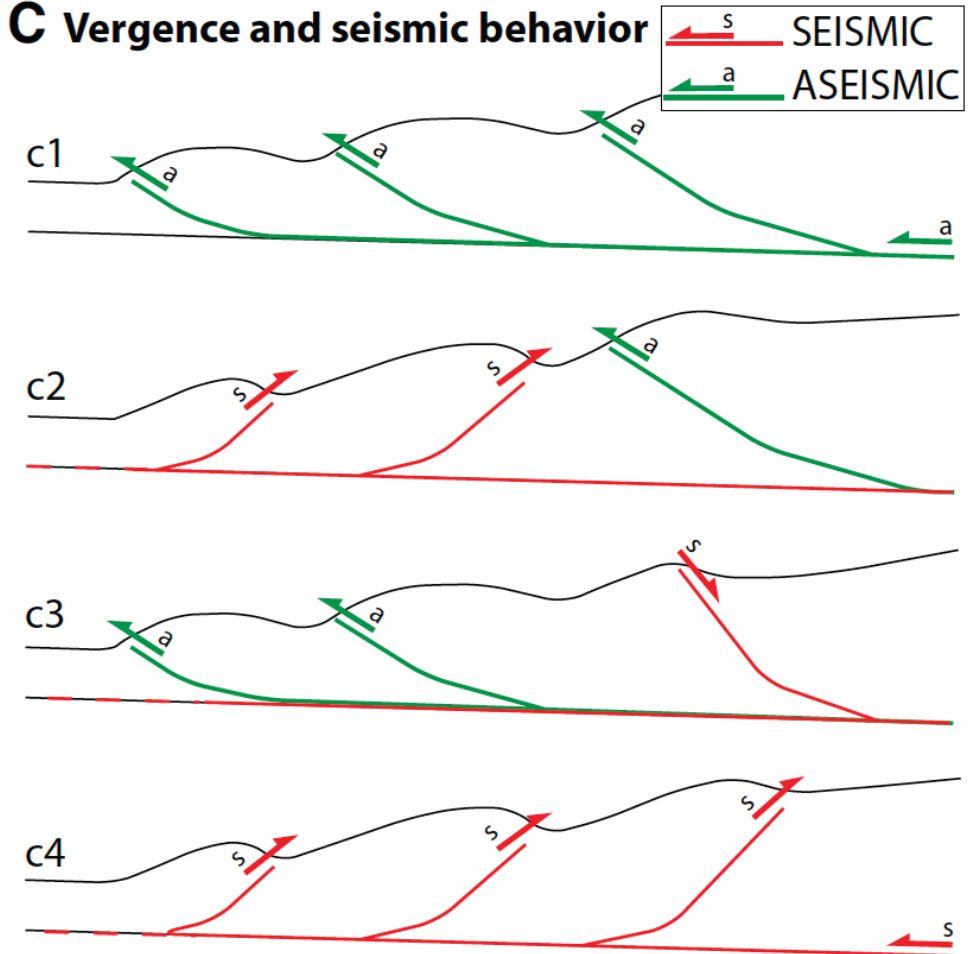
A High sedimentation rate



B Friction drop



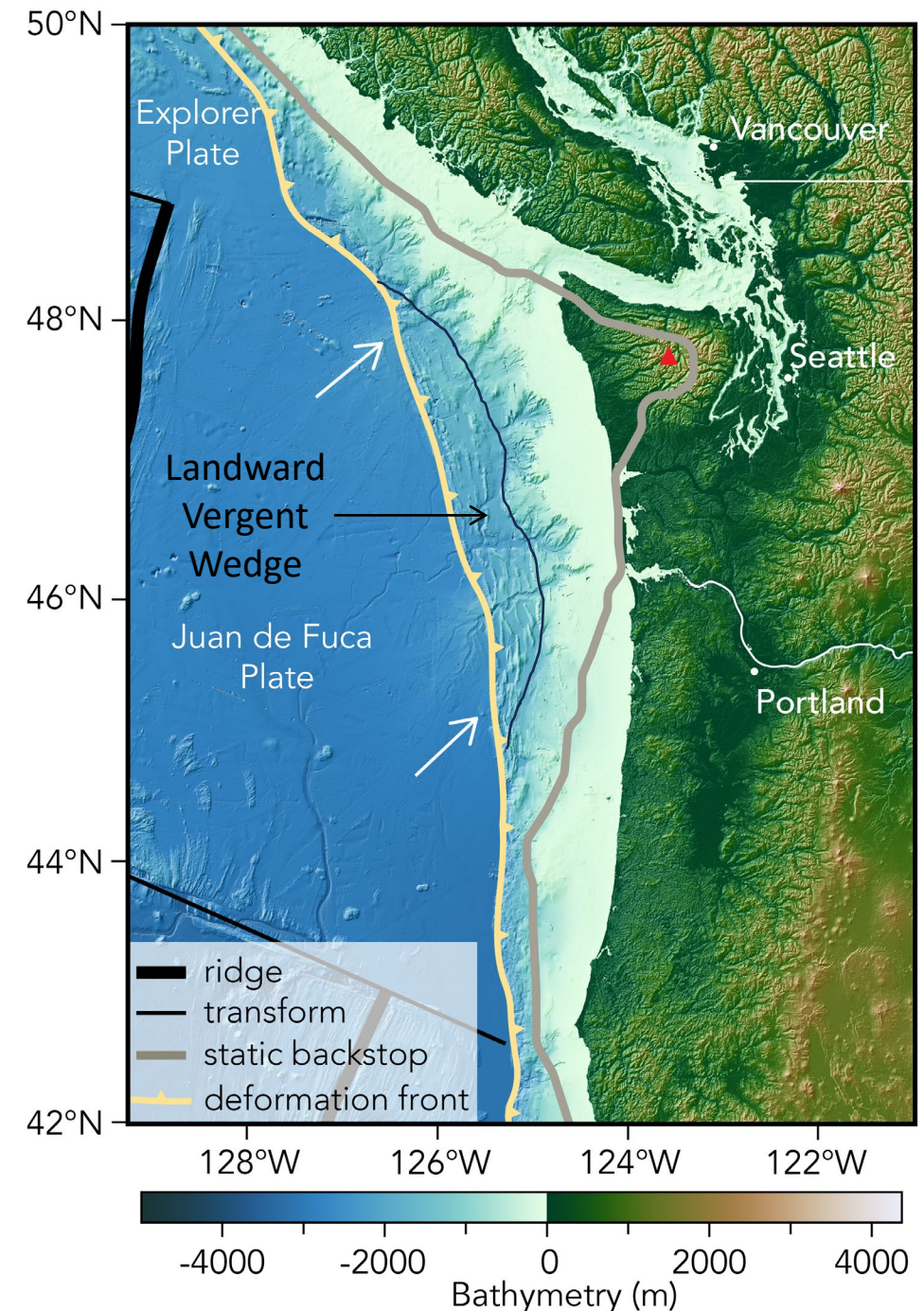
C Vergence and seismic behavior



Macro-scale geologic structure is a seismic indicator: Dynamic friction is a necessary condition for LVZ to form.

The megathrust (and outer splay faults) are capable of locking and contributing to slip deficit to the deformation front.

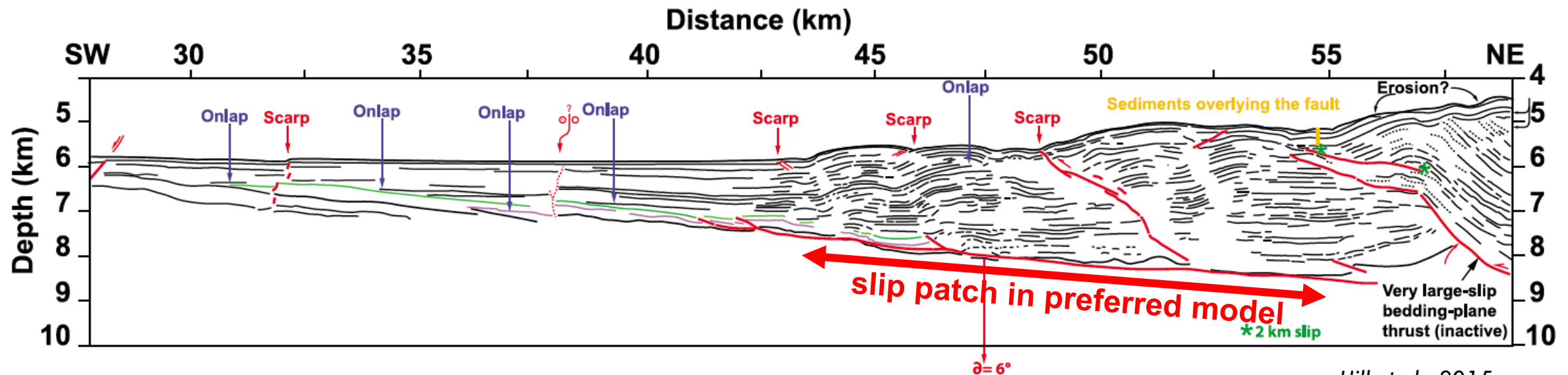
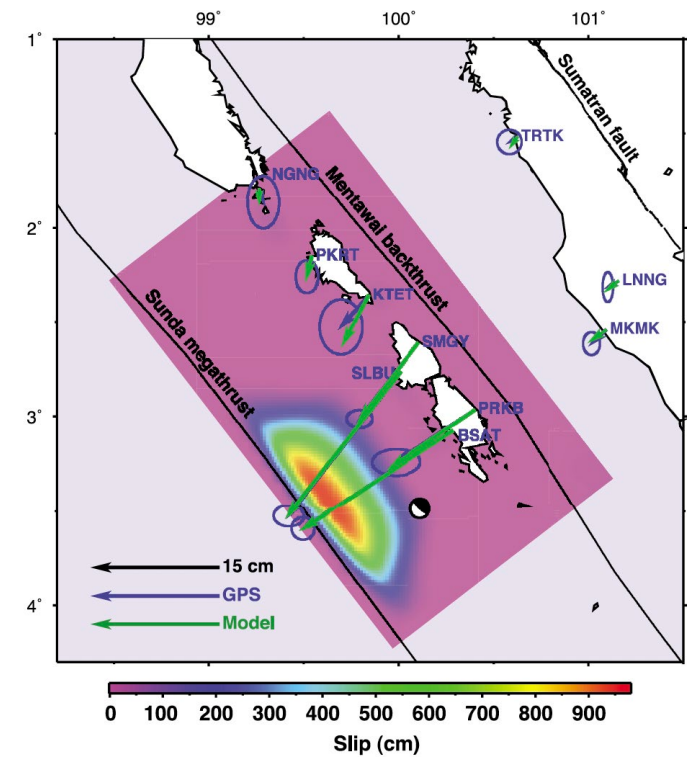
Potentially tsunamigenic surface displacements over a ~30-km wide and 400 km long must be evaluated.



2010 Mentawai M_w 7.8 tsunami earthquake

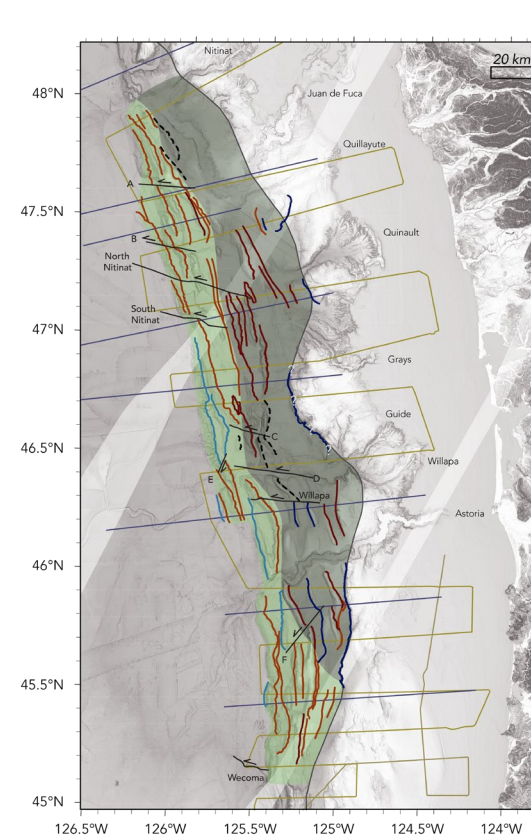
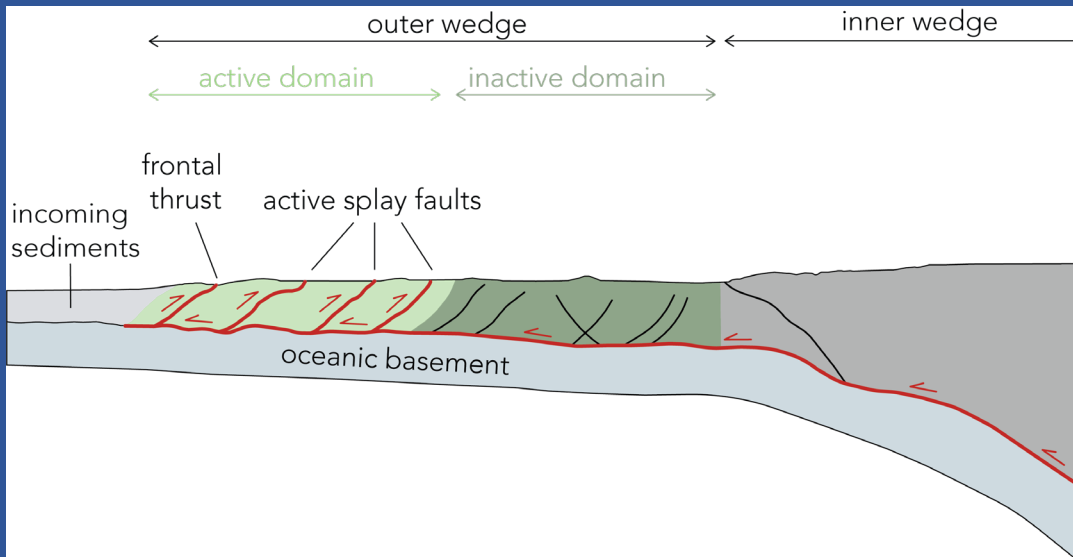
Near-field GNSS and tsunami inundation runup inverted for rupture patch

- Tsunami runups in excess of 16 meters documented
- GNSS sites on the forearc Mentawai Islands
- Seismic reflection line in agreement with model, showing the rupture was confined to the outermost wedge, hypocenter only 6 km depth
- Peak slip in the model of ~10 meters



Take Home Message

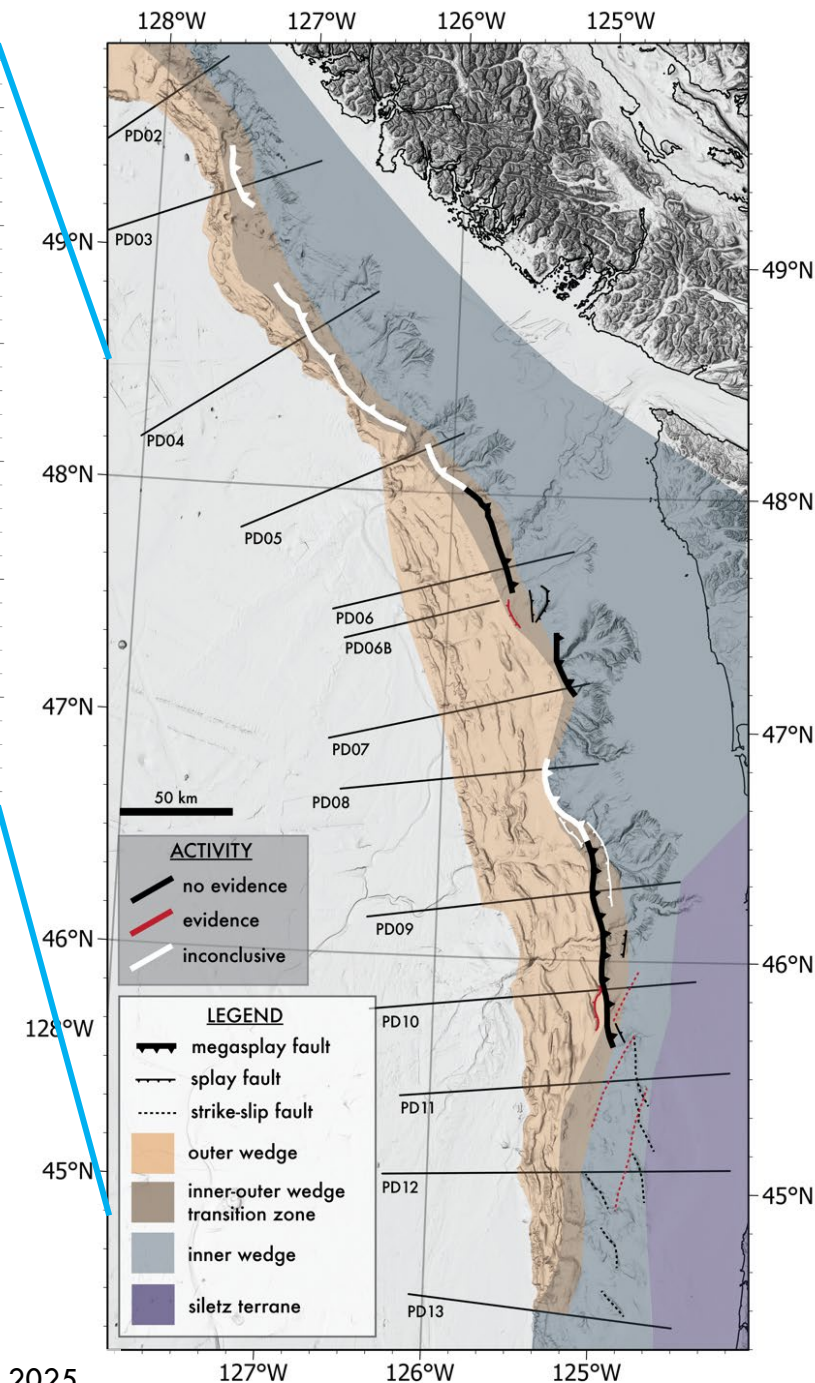
- North-central Cascadia is statically strong but dynamically weak
- Surface displacement may be enhanced by slip up splay faults in the outermost outer wedge
- This locked region is capable of M9 class rupture by itself – full margin rupture is possible but not required.



Ledeczi et al., 2024

Many recently-active faults in outermost region (light green)

“Megasplay” fault is discontinuous to absent



Lucas et al., 2025