

California Department of Conservation California Geological Survey Seismic Hazards | Tsunami Unit

Cascadia PTHA Updates

- USGS Powell Center Meeting
 - Subject Matter Experts
- CSZ Logic Tree
- CSZ Megathrust Geometry

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reddit.com/r/gifs/comments/ftlhkz/this_is_a_gif_i_did_by_myself_inspired_by_an_edit/

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USGS Tsunami Sources Powell Center Working Group on Tsunami Sources: Probabilistic Tsunami Hazard Assessment for the Cascadia Subduction Zone



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In May 2022 live and remote participants in the photo gathered to discuss peer reviewed literature that has implications for tsunamigenesis along the Cascadia subduction zone (the CSZ).



The presentations and discussion were used to develop a logic tree that will be used as input for Probabilistic Tsunami Hazard Assessment (PTHA) in Cascadia. One major goal is to keep this PTHA consistent with the USGS NSHM 2023 update.

PTHA and Logic Trees

Probabilistic tsunami hazard assessment (PTHA) is a process used to develop a framework to inform tsunami modeling and hazards analyses.

This process uses a logic tree approach to collate the entire suite of possible and probable tsunamigenic behavior of tsunami sources relevant to the area of interest.



A logic tree is a way to calculate the relative likelihood for each of all possible scenarios for a given phenomenon.



PTHA and Logic Trees

In the logic tree, each possible scenario is organized as a separate branch. Each branch is given a weight, based on expert opinion, that represents the chance that a scenario may happen. Branch weights, for each splay, are additive vertically & sum to 1:

0.025 + 0.025 + 0.16 + 0.53 + 0.26 = 1 (or 100%)

The weights for each splay are multiplied horizontally to calculate the scenario weight:

0.053 * 0.2 = 0.11 (or 11%)



The result is the total weight for each scenario. These total weights also add up to 1, vertically (or 100%).



Logic tree results and how to prepare tsunami hazard curves:

- Logic tree weights are used to constrain slip and tsunami source modeling.
- These tsunami models result in a suite of offshore tsunami heights, representing the percent likelihood (chance) for tsunami sizes for given annual probability of exceedance.
- Using this entire suite of model results, we can calculate quantiles that bracket a range of probabilities.
- From these data we can calculate the tsunami size for tsunami with return periods (such as the "975year tsunami").

Tsunami Model Results

Each line represents an actualization of a logic tree scenario tsunami, a percent chance (likelihood) for tsunami size with an annual probability of occurrence.



Tsunami Model Quantiles

Each line represents a bracketed summary of the data plotted on the left. E.g., the 0.05 and 0.95 lines bracket 90% of the scenario tsunami plotted on the left.



Cascadia Megathrust Geometry:

- Kelin Wang and Matthew Sypus constructed a CSZ megathrust surface deformation model that is being used for the initial conditions for the tsunami wave simulations.
- We have enlisted the cooperation of Suzanne Carbotte and Harold Tobin who are working with their students and collaborators to establish the fault geometry for the CSZ. For the up-dip region of the fault this work is based on new seismic data collected in 2021 during a margin-wide research cruise called CASIE21. The fault geometry in the down-dip region is based on low frequency earthquake analyses from Michael Bostock and their collaborators.

Dislocation Model

The surface deformation is calculated by numerically integrating point-source dislocation solutions of Okada (1992).

(Wang et al., 2003; Wang, 2012)

Slab surface

Fault mesh

Depth (km)

OPEN-FILE REPORT O-24-11

IMPROVED CASCADIA EARTHQUAKE SOURCE MODELS FOR TSUNAMI HAZARD ASSESSMENT

By Matthew Sypus and Kelin Wang¹

https://www.oregon.gov/dogami/pubs/Pages/ofr/p-O-24-11.aspx



Megathrust Contours

Matt constructed two surfaces: (1) top of crust (blue lines on profiles), (2) décollement (orange lines in profiles).



CASIE21: The décollement is in different positions along different parts of the megathrust. Offshore Washington, the décollement is near the top of the crust. Offshore Oregon, the décollement is within the sedimentary section (there is sediment subduction here).



Megathrust Contours

Note the thick sedimentary section offshore of Oregon (difference between orange and blue lines in profiles)



Megathrust Contours

Anne Tréhu provides picks for SCSZ where we lack CASIE21 data.



An attempt to sort out the different logic trees, and suggestion for the next version. Feel free to correct, comment, etc.

Only symmetric for splay scenarios















Along Strike Variation in Slip

Map shows the northern and southern limits of the whole-margin (A, A+exp) and segmented ruptures (B–F)



Along Dip Variation in Slip

Map shows three downdip slip termination boundaries as adopted by the USGS NSHM:

- Midpoint between fully locked zone and 1 cm/yr locking contour (red)
 - The 1 cm/yr locking contour (green)
 - The top of the ETS zone (purple)
- Map shows updip boundaries of rupture:
 - deformation front (black)
 - deep buried termination (gray)
 - splay B (blue)

•

splay D (orange)



Slip in the downdip direction

A) Updip skewed (red), symmetric (black), and downdip skewed (blue) bell-shape slip distributions with q= 0.3, 0.5, and 0.7, respectively.

B) For splay-faulting rupture, the bell shape slip profile is cut off at the fault trace.

C) For portions of our trench-breaching rupture, the distribution mirrors the bell shape downdip of the peak slip but tapers updip to the trench to a prescribed percentage of the peak slip.



Splay Faulting in Cascadia

There is a splay fault that has been identified in seismic reflection profiles and has been used by some states as a tsunami source. The participants debated the length and activity of this fault system. There are places where the fault does not leave evidence for activity at the surface (so if it is active, the sedimentation rate exceeds the slip rate).

This splay fault (white line at left) is generally at the boundary between the inner (younger) and outer (older) wedges (green line in map on the right).







Splay Faults

⁻⁴⁸ Traces of the two assumed splay faults used for modelling.
 ⁻⁴⁸ Splay B (or "extended L1") is shown in blue and the new
 ⁻⁴⁸ Splay D trace is in green. South of 46.5°N, the two traces
 ⁻⁴⁶ fully overlap. Yellow squares are the interpreted splay ⁻⁴⁶ fault locations from the CASIE21 profiles.

Vergence Mapping from Chris Goldfinger (triangles) and Janet Watt (hexagons)





Fault Vergence

- Fault vergence in the outer wedge can also have influence on fault slip and tsunamigenesis.
- Watt & Brothers (2021) remind us that landward vergent regions may be more favorable to strain accumulation and more prone to trench breaking rupture (e.g., Han et al., 2017; Beeson et al., 2017).
- Watt and Brothers (2021) present an overview of fault vergence
- Following the workshop, we compared their observations with those of others, like from Chris Goldfinger and results from CASIE21.

Watt and Brothers, 2021

Vergence Mapping from Chris Goldfinger (triangles) and Janet Watt (hexagons)





46°N

45°N



Frontal Thrust Geometry

The top left: hypothetical seaward and landward vergence faulting in cross section view.

- A) Vergence directions in the frontal accretionary prism compiled by CSWG members Janet Watt and Chris Goldfinger.
- B) Vergence directions compiled by CASIE21 member Shuoshuo Han.
- C) Simplification for modeling trenchbreaching rupture scenarios.

Where there is any evidence for seaward-vergent thrust faulting (yellow), a single frontal thrust is used. Where there is no such evidence (green), highly up-skewed buried slip is assumed.





Examples for Whole Margin (A) Deep and (B) Shallow Buried Rupture

Distance from deformation front (km)

Profile '

Uplift (m) 20 ip (m) Deformation front Deformation front Deformation front th (km) Megathrust 20 40 50 150 50 200 50 0 100 0 150Λ 100

Profile 2

Profile 3

Distance from deformation front (km)

Maps

Distance from deformation front (km)

- Slip Distribution
- Vertical Deformation Along Dip Profiles
 - Vertical Deformation
- Slip Distribution
- Megathrust Geometry





Examples for Whole Margin (A) up-skew (B) symmetric (C) down skew buried rupture Maps

- **Slip Distribution**
- **Vertical Deformation Along Dip Profiles**

Profile 3

Megathr

50

- **Vertical Deformation**
- **Slip Distribution**
- **Megathrust Geometry**

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Examples for Whole Margin (A) Midpoint
 locked & 1 cm locked, (B) Locked zone and (C)
 ETS downdip limits

- Slip Distribution
- Vertical Deformation Along Dip Profiles
 - Vertical Deformation
- Slip Distribution
- Megathrust Geometry



128'W



126 W

124 W

128 W

126 W



Examples for Whole Margin for (A) Splay B and (B) Splay D

Distance from deformation front (km)

Profile 1

Profile 2 Profile 3 Jplift (m) 20 lip (m) 10 Deformation front ↓ D,B Deformation front Deformation front Megathrus[,] 20 40 0 50 100 150 0 50 150 200 0 50 100 100

Maps

Distance from deformation front (km)

- Slip Distribution
- Vertical Deformation Along Dip Profiles

Distance from deformation front (km)

- Vertical Deformation
- Slip Distribution
- Megathrust Geometry









Examples for Whole Margin Trench Breaching (A) symmetric (B) up-skew w/lower downdip rupture extent. Equivalent 100% peak slip breaches for (C) symmetric (D)up-skew

- Maps
- Slip Distribution
- Vertical Deformation Along Dip Profiles
- Vertical Deformation
- Slip Distribution
 - Megathrust Geometry



Califo

Examples for DOGAMI Whole Margin asperity models (A) northern (B) center (C) southern asperities. These use symmetric shallow buried rupture.



Examples for Whole Margin asperity models. These use symmetric shallow buried rupture with 1 cm/yr locking downdip limit.





Definition of the five-level cluster ruptures based on the cluster setup in the USGS

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122°W

Five Level Cluster Ruptures



4a-4 Mw8.712

4b-4 Mw8.777

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128°W 126°W 124°W 5-4 Mw8.713 48°N 46°N 126°W 124°W 128°W 5-3 Mw8.509 46°N 44°N 124°W 128°W 126°W 5-2 Mw8.418 44°N 42°N 126°W 124°W 128°W 5-1 Mw8.373 42°N 40°N 128°W 126°W 124°W

5-5 Mw8.628

50°N



Floating rupture boundaries

The use of red and blue colors is only for display
 clarity with no other significance.
 All the ruptures are elliptical and of the same size in map view.





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Thank You!

Are there any questions?

Clear your mind of questions.

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giphy.com/gifs/starwars-3ornk03njkdi5mNKJG





The full rupture branch presumes Return & a margin-wide Periods 。 earthquake (e.g., for Cascadia Earthquakes Goldfinger et al., 2012, 2017). The partial rupture branch represents earthquakes that do not span the 100 km VGS 1984 UTM Zone 1 entire margin.

ears (110-1150)

years

years (120-720)

320

years

220

years (40-580)

years?

Event-04-1-1-2-04

Event_04_1_1_3_

Event-04-1-1-1-04

(40-720)

(120-720)

B

The two partial branches are segmented and floating: (1) based on the Goldfinger et al. (2012, 2017) paleoseismic scenarios (2) based on floating ruptures





PTHA and Logic Trees

In the logic tree, the order of the branches does not matter.

Example on the left has the deep/shallow branch before buried/splay branch. The example on the right has the buried/splay branch before the deep/shallow branch. Note how the probabilities for each option sum to 1.00.



Rupture Size – Magnitude Scaling Relations

These empirical scaling relations allow us to take rupture dimensions and estimate a magnitude for that sized rupture.
 Example scaling relations from Strasser et al., 2010 show magnitude vs. rupture length, rupture width, and rupture area.





Wirth and Frankel, 2019

Down-Dip Limit of Rupture

The slip distribution at the down dip limit of rupture may have little influence on tsunamigenesis, unlike for ground motions used in the USGS NSHM. This highlights some ways in which these logic trees must be slightly different.

(a) the Japan view



(b) a North America view



(c) a hybrid view



Wang and Trehu, 2014



Long Term Slip Rate: Coupling Ratio

Participants agreed that long term slip needs to be balanced by plate motion rate in some way.

- The coupling ratio is amount of plate motion rate that is accumulated as tectonic strain (i.e., the ratio of seismic vs. aseismic slip).
- If the coupling ratio is less than 1, then the entire plate rate that contributes to long term slip is less than the plate convergence rate.













 $\begin{array}{c}
\uparrow u' = u \cos \alpha \approx u \\
\Rightarrow \alpha + \beta \\
\uparrow u \\
\hline \end{array}$

B)

 D_w : water depth at actual trench D_s : sediment depth at actual trench u: real rise of seafloor u': modelled rise of seafloor α : seafloor slope angle β : near trench fault dip s: fault slip





















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Thank you

Any Questions?

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