



# **PTHA AND COASTAL PROCESSES: Non-stationary statistics and compounding hazards**

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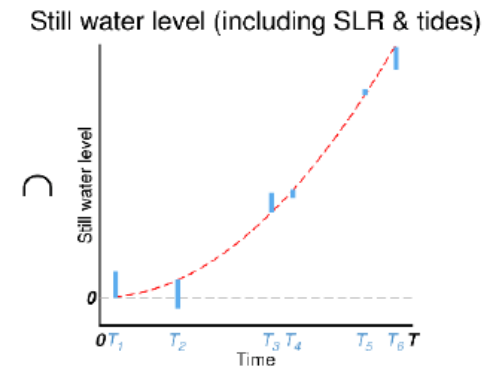
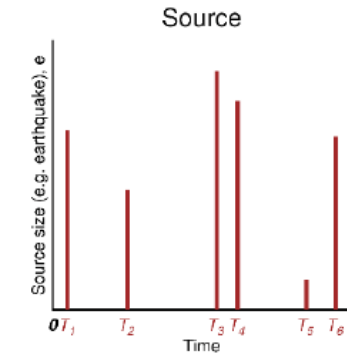
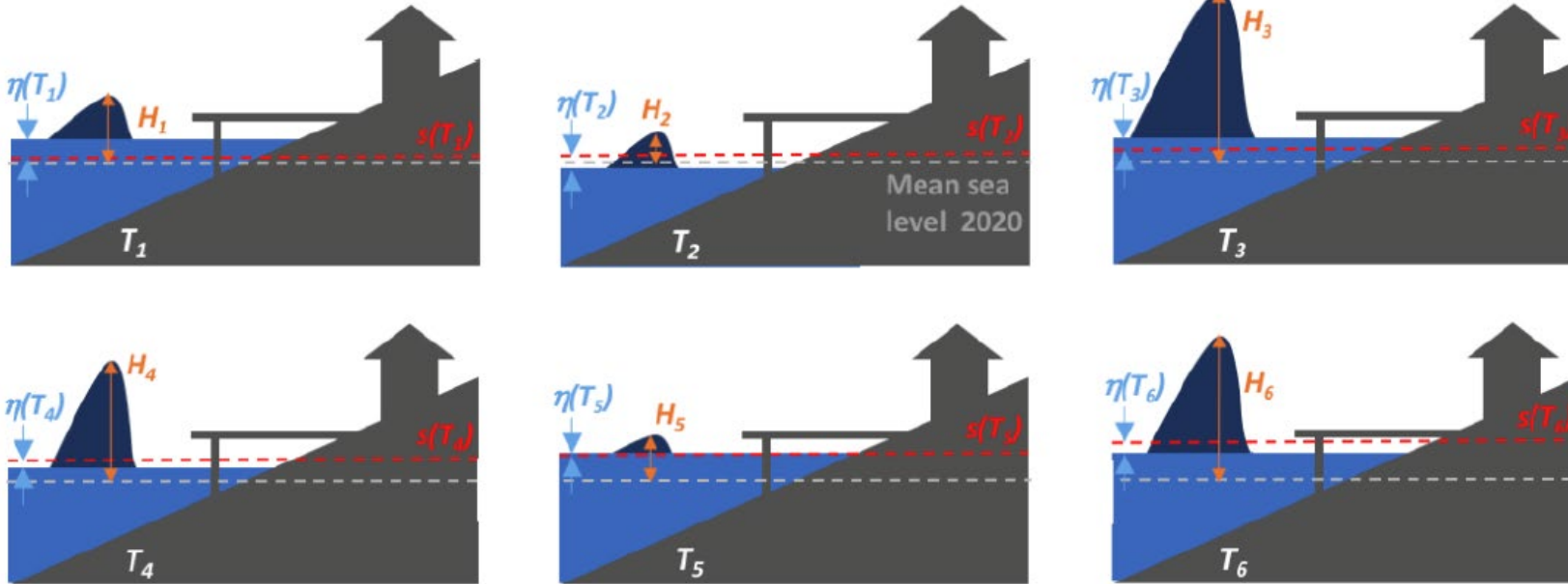
**SDSU**

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Engineering Lab**

# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS

## PTHA AND NON-STATIONARY COASTAL PROPERTIES



A new era in coastal hazards assessments should incorporate **daily-varying** and **non-stationary processes in time scales comparable to coastal project lifetimes**.

For example, tides and the climate-change-driven sea level rise.

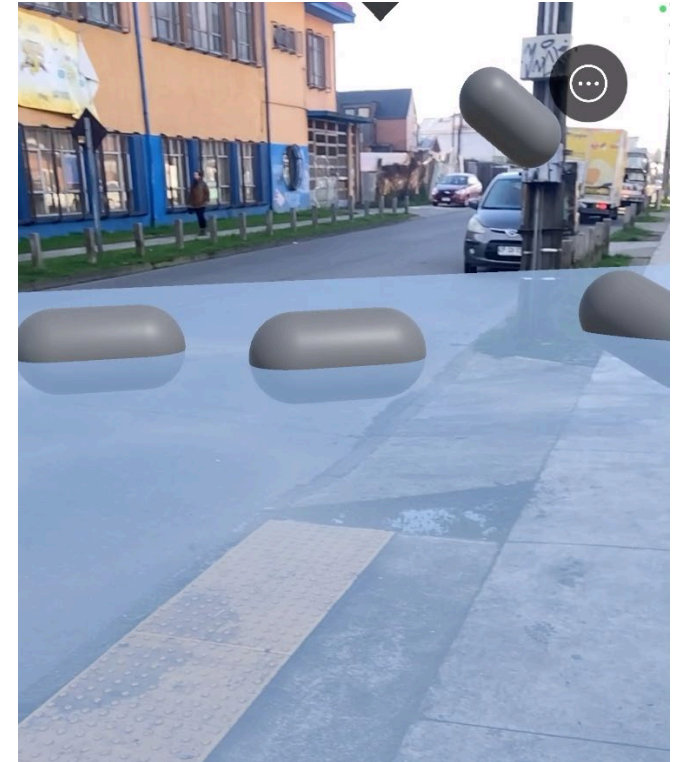


# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS

## PTHA AND NON-STATIONARY COASTAL PROPERTIES



Even, **vulnerabilities** are non-stationary...



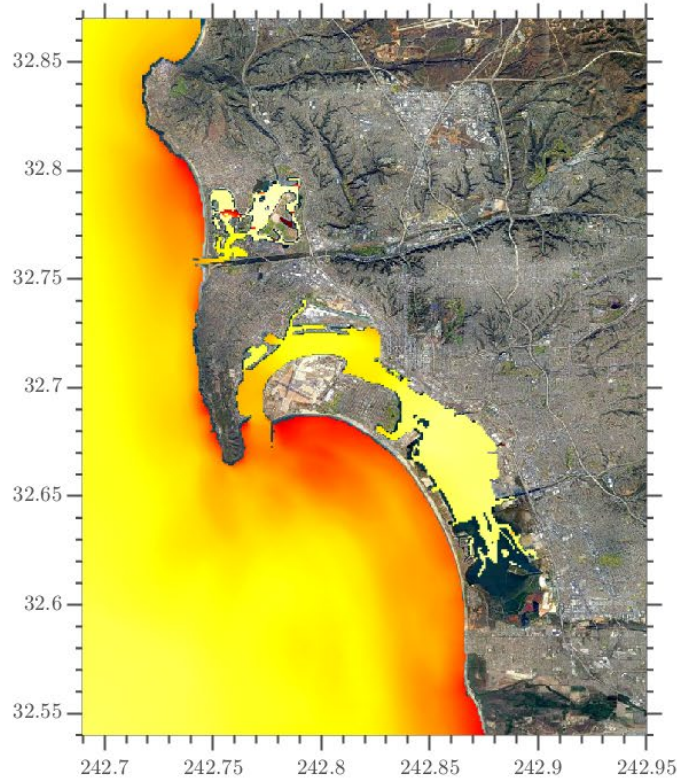
# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS AND COASTAL PROCESSES

## TSUNAMI HAZARDS AFFECTED BY COASTAL TIDES, SLR AND BATHYMETRY

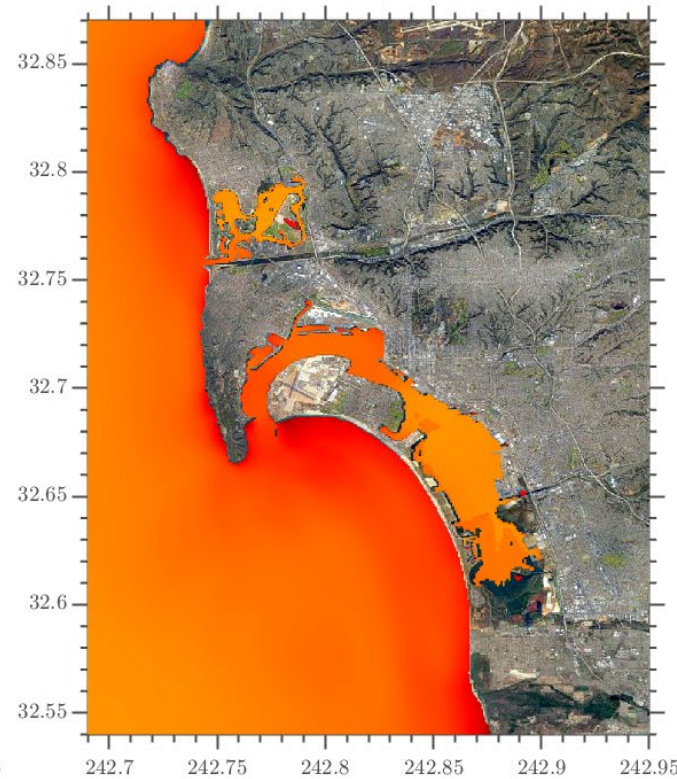


Maximum Considered Tsunami (MCT) in **ASCE 7-16**: Those exceeded with probability of 2% in 50 yrs.

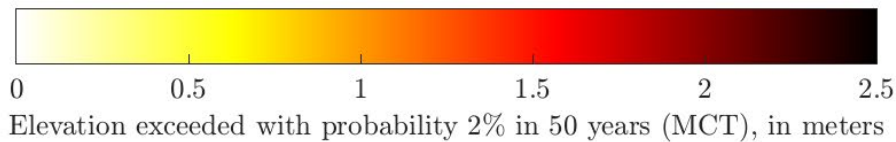
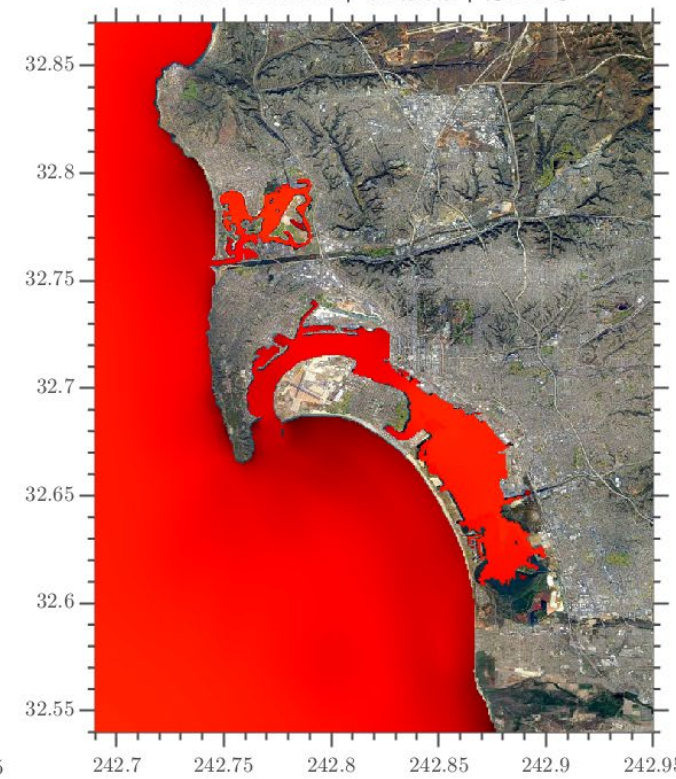
PTHA



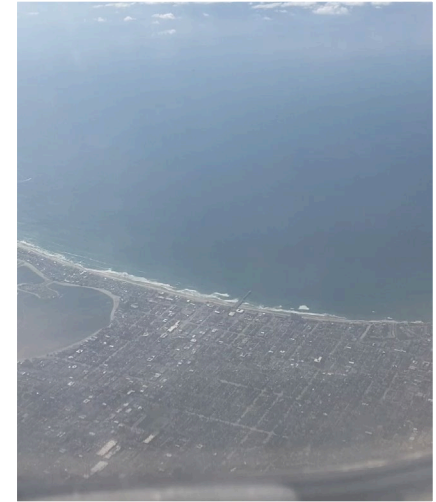
PTHA+tides



nPTHA+tides+SLR



*Sepulveda & Mosqueda (under review)*



Should the “ASCE 7-16 Tsunami Loads and Effects” incorporate coastal influences when developing MCT maps?

# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS WITH SEA LEVEL RISE

## Non Stationary Poisson Process



But my future  
Is non-stationary



i.e.  $p_i(H > h_c | E_i)$  and  
 $\lambda_i$  varying in time

$$P[N_{H>h_c} > 0, T] = 1 - e^{-\sum_{i=1}^n \lambda_i p_i(H > h_c | E_i) T}$$

Let's revisit the fundamentals of the Poisson process...

Consider an infinitesimal partition of T into  $m$  sub-intervals  $dt$ , such that the probability of having 0 or 1 events is nearly 1, and a sea state  $s(t)$  varying in time. **The probability of having**

$$P[N_{H>h_c} > 0, T] = 1 - e^{-\sum_{i=1}^n \int_0^T \lambda_i p_i(H > h_c | E_i, s(t)) dt}$$

$$1 - p_i(H > h_c | E_i, s(t)) \lambda_i dt \approx e^{-\lambda_i p_i(H > h_c | E_i, s(t)) dt}$$

The overall probability in T can be viewed as a sequence of independent Bernoulli trials (i.e. a product of probabilities). This is also known as the non-homogeneous Poisson process:

$$\prod_m e^{-\lambda_i p_i(H > h_c | E_i, s(t)) dt} = e^{-\int_0^T \lambda_i p_i(H > h_c | E_i, s(t)) dt}$$



# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS WITH SEA LEVEL RISE

Non Stationary PTHA



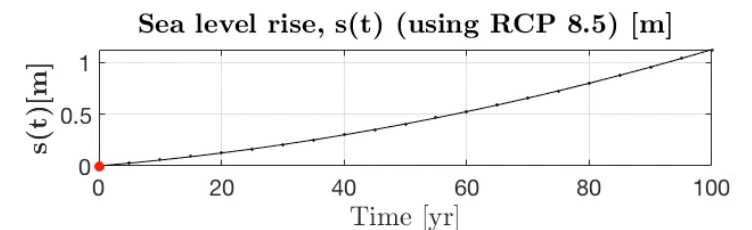
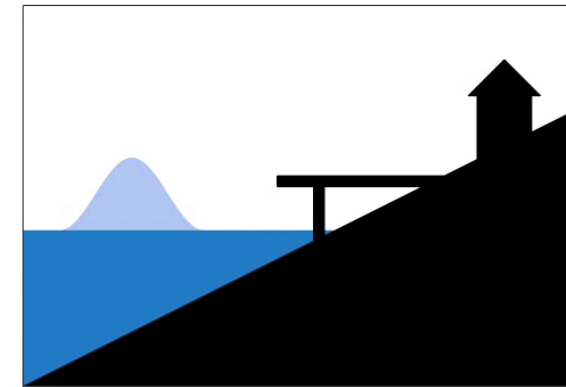
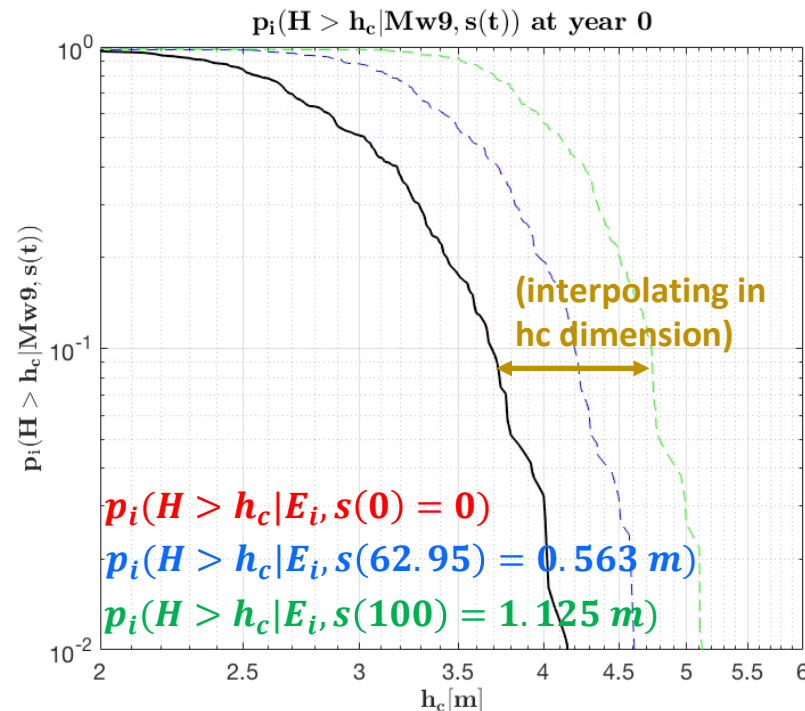
We introduce a non-stationary probabilistic tsunami hazard assessment (nPTHA)

$$P[N_{H>h_c} > 0, T] = 1 - e^{-\sum_{i=1}^n \int_0^T \lambda_i p_i(H>h_c|E_i, s(t)) dt}$$

$p_i$  is now a function of time ( $\lambda_i$  as well if you would like). One example is the effect of sea level rise.

Great, but how do we get  $\int_0^T p_i(H > h_c|E_i, s(t)) dt$  ?...

We apply a surrogate model



# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS WITH TIDES

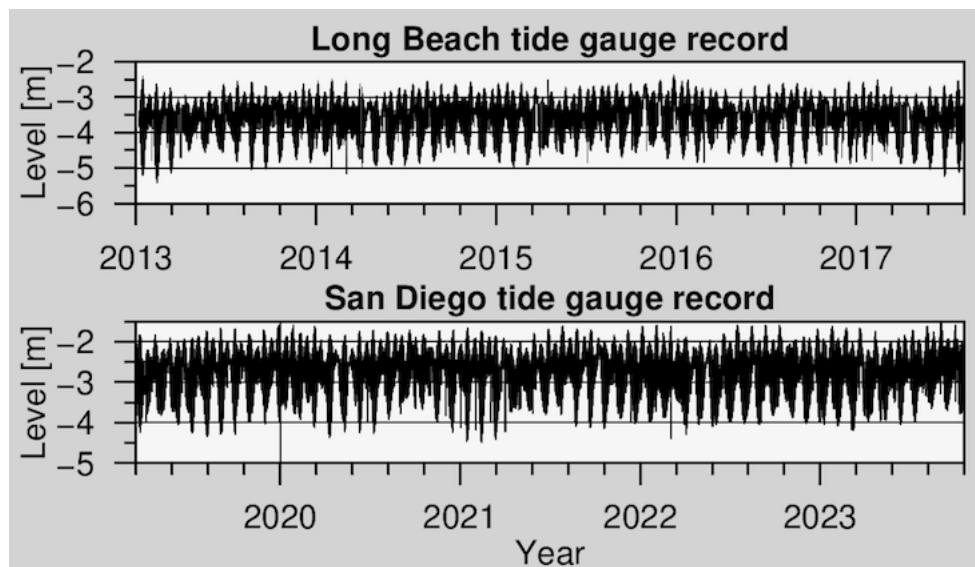
Incorporation of tides in nPTHA



Tidal ranges are comparable to SLR for the next few decades. It is necessary to incorporate them in the nPTHA.

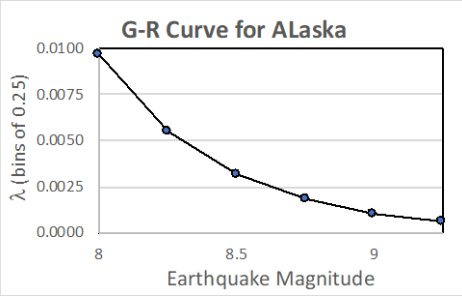
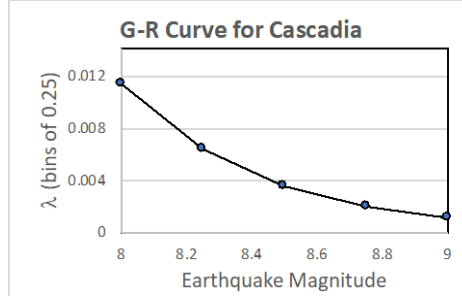
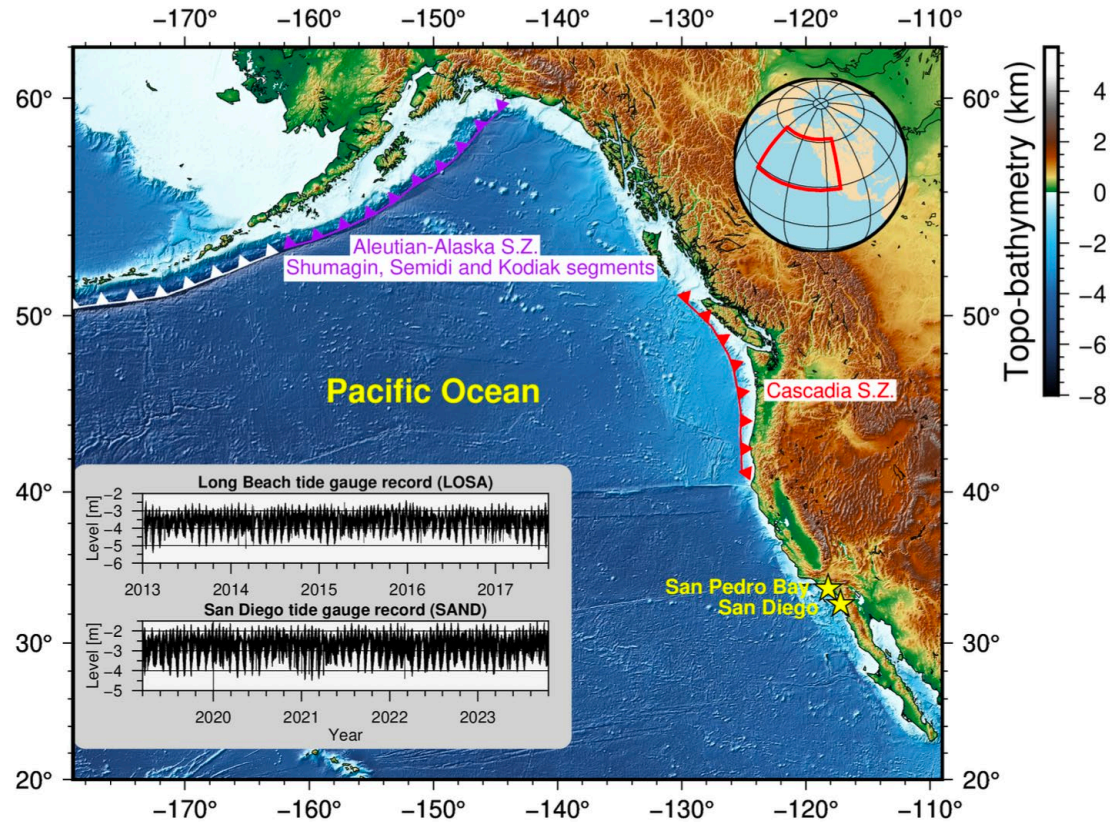
$$P(N_{h_c}(T) = 0) = e^{-\sum_i \lambda_{E_i} \int_0^T \int_{\eta_{min}}^{\eta_{max}} f_{\eta}(\eta) p_i(H > h_c | E_i, s(t), \eta) d\eta dt}$$

nPTHA with tides

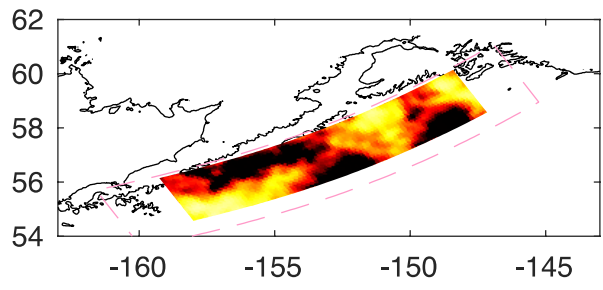


# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS

## The classic PTHA in Southern California

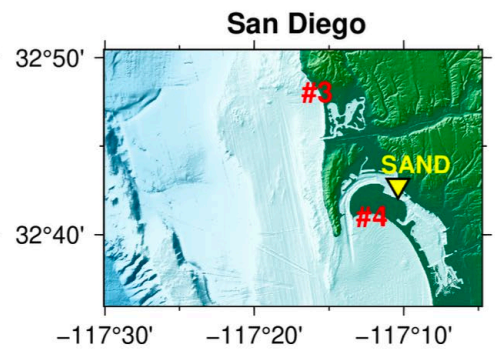
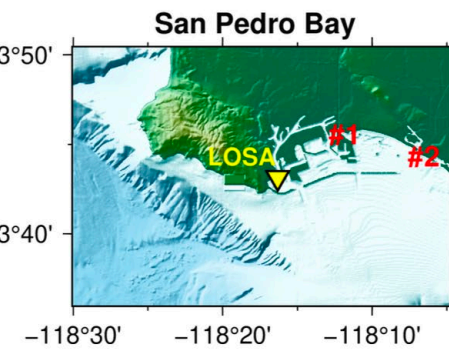


1) G-R curves:  
Alternative values of  $\lambda_i$  are determined.



2) Earthquakes with random slip and location are modeled (Sepulveda et al, 2017).

$p_i(H > h_c | i)$  are calculated for  $i=[1,11]$  ensembles of earthquakes and using a tsunami model.



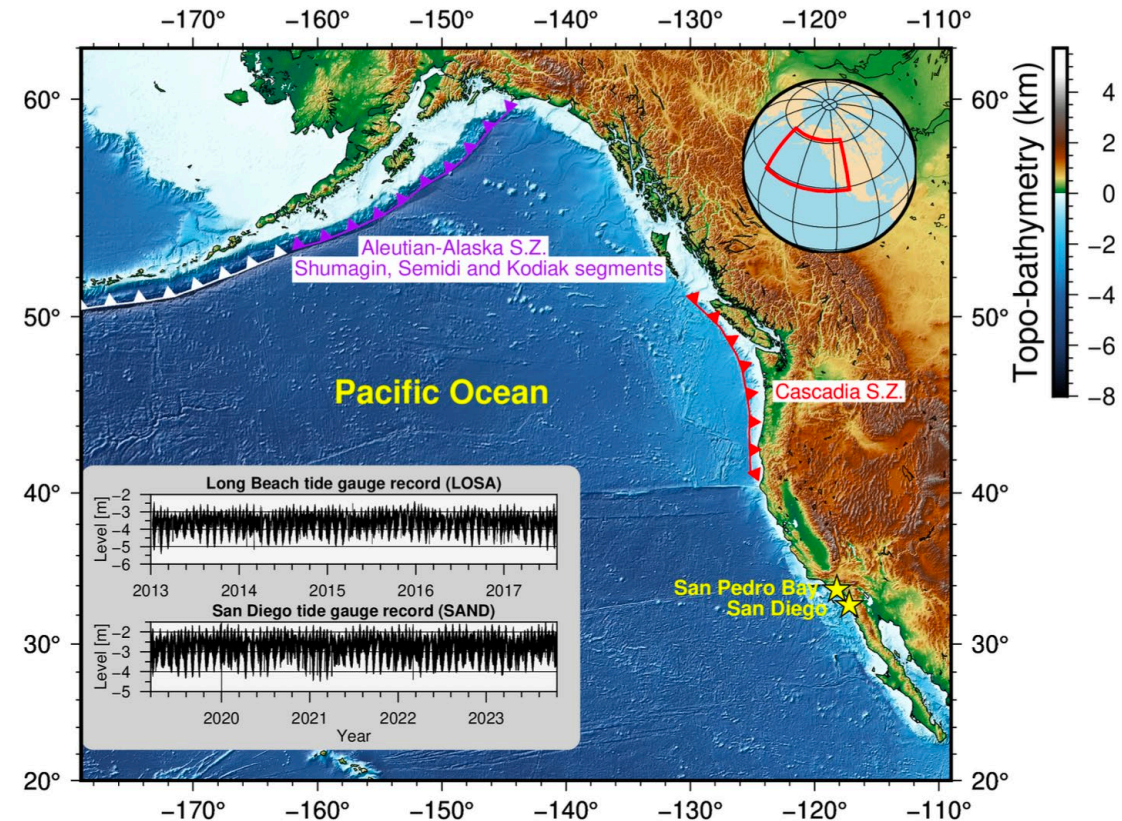
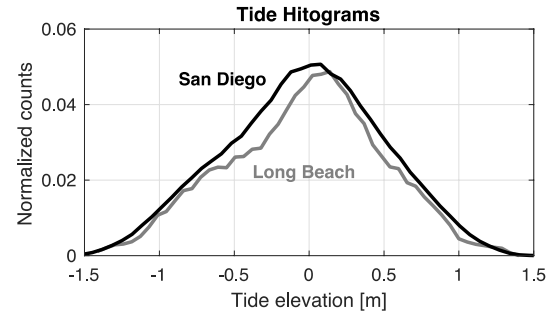
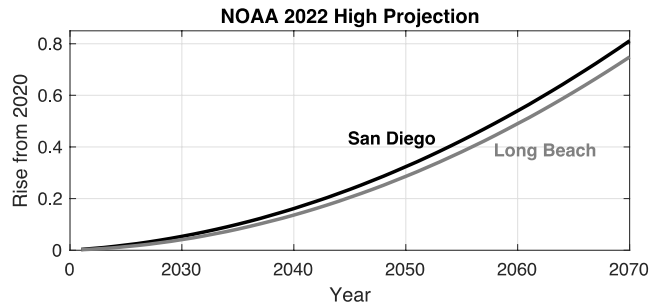


# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS

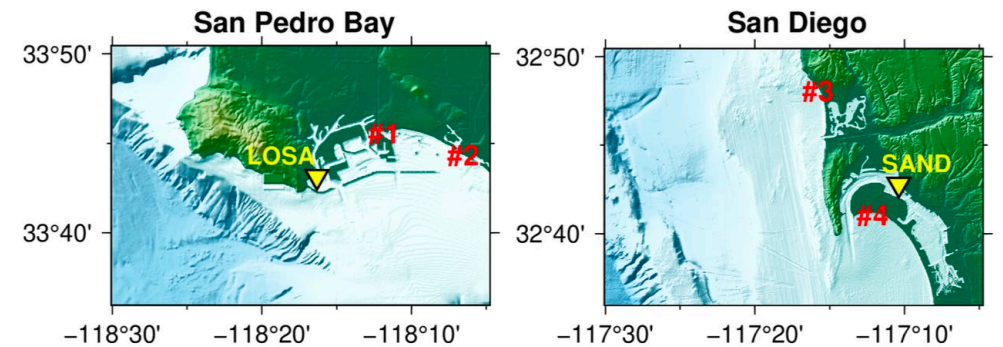
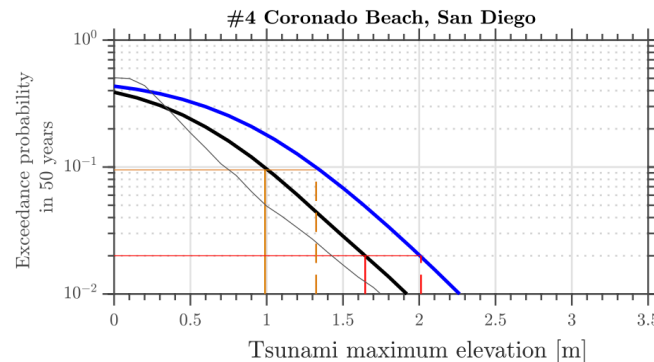
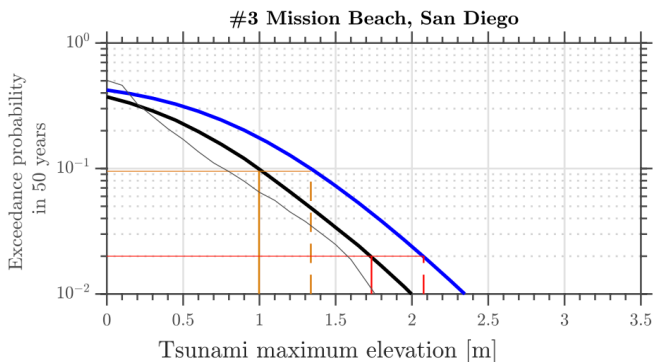
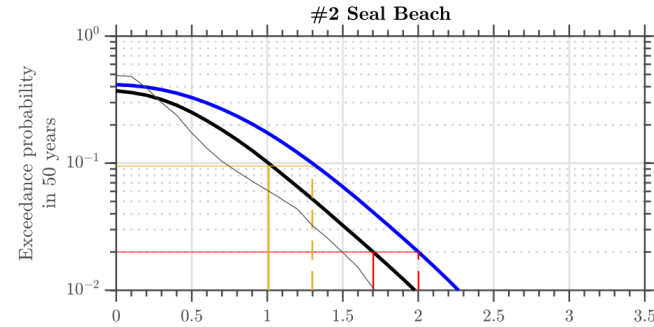
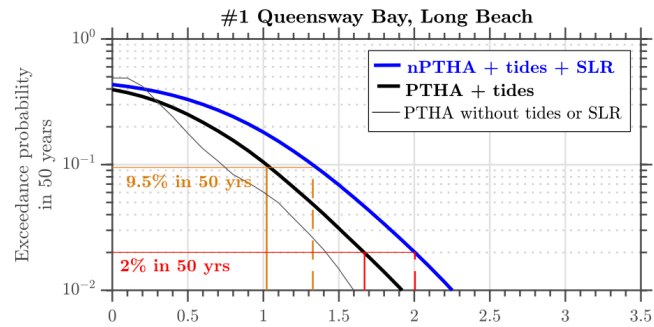
## nPTHA IN Southern California incorporating tides and SLR



Let's apply the nPTHA!! We obtain the sea level rise projections and tides of Long Beach and San Diego. We use the new nPTHA.



Exceedance probabilities in 50 years can be obtained.

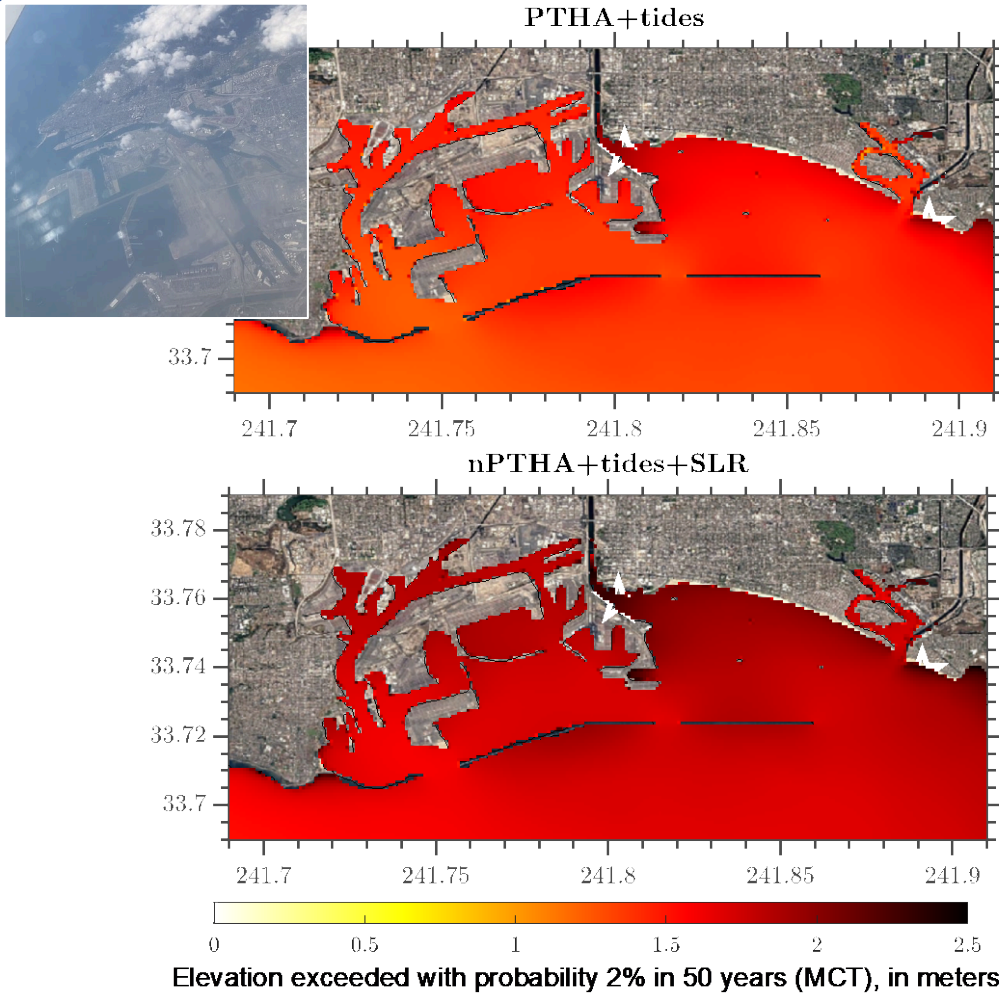


# PROBABILISTIC TSUNAMI HAZARD ASSESSMENTS

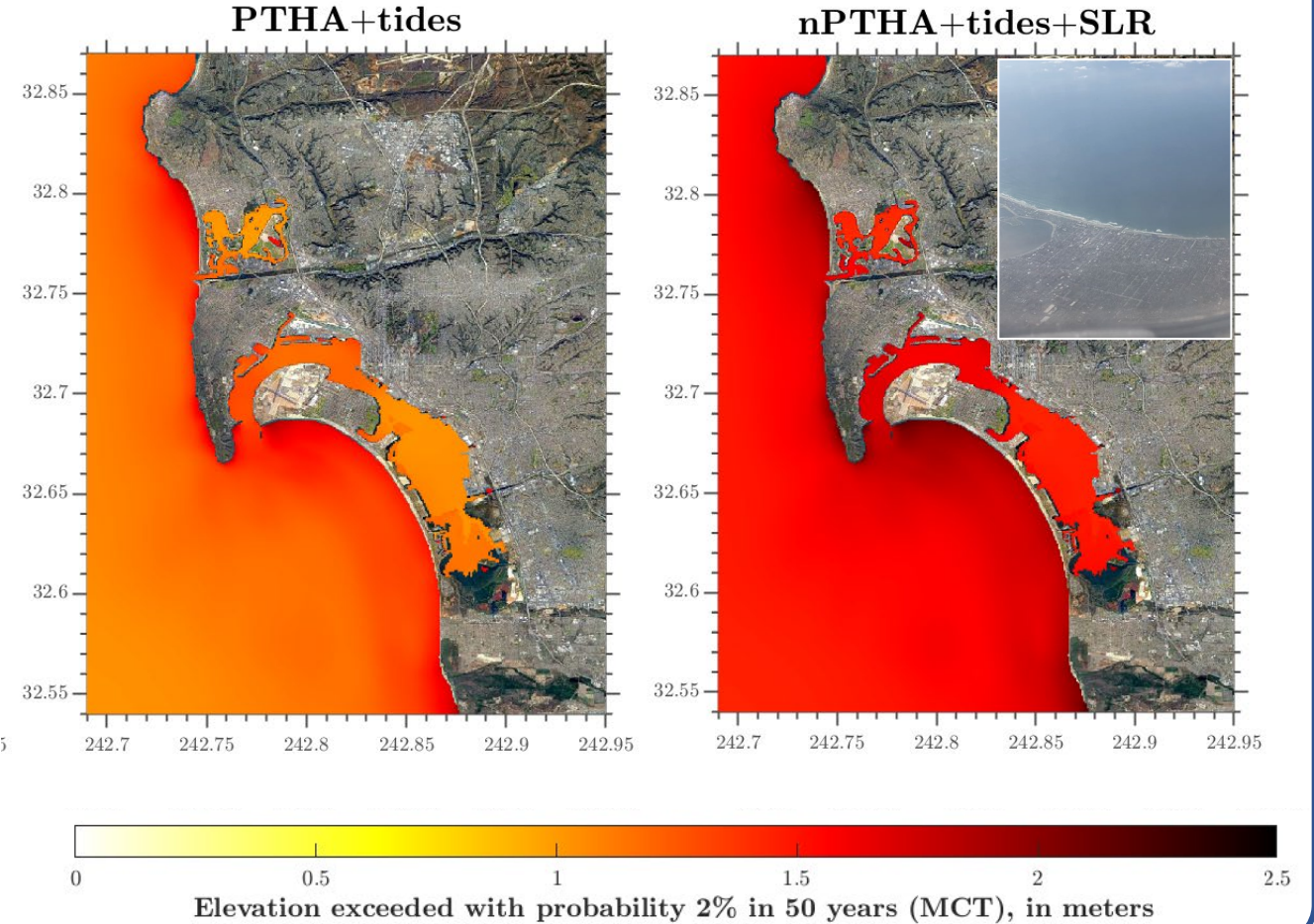
## nPTHA in San Pedro Bay and San Diego (effects of tides and SLR)



Maximum Considered Tsunami (MCT) in ASCE 7-16 Tsunami Loads and Effects: Those exceeded with probability of 2% in 50 yrs.



*Sepulveda & Mosqueda (under review)*

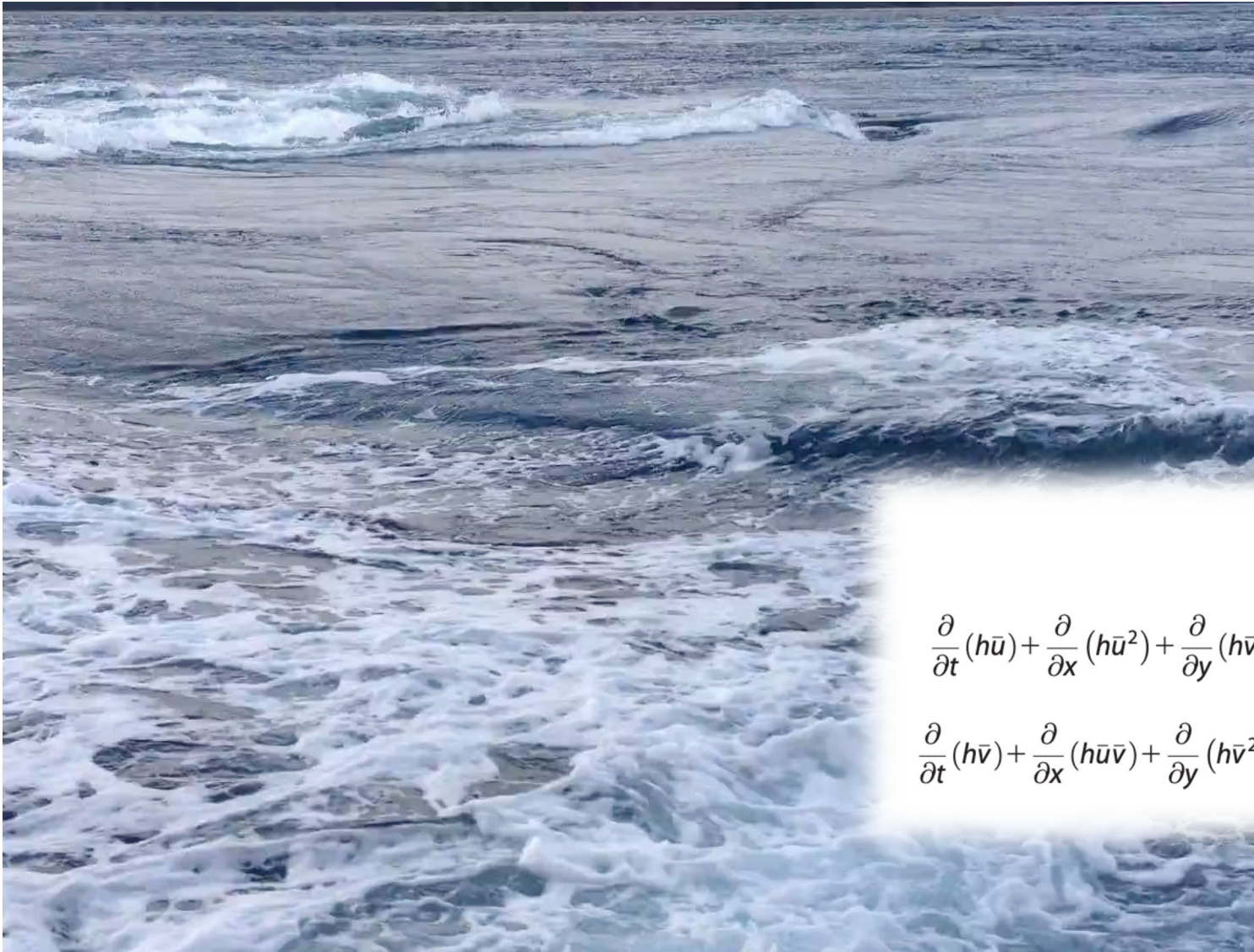


Should the "ASCE 7-16 Tsunami Loads and Effects" incorporate climate change and tides in MCT?

# TSUNAMI-TIDE INTERACTION



## TSUNAMI-TIDE INTERACTION: THE DETERMINISTIC FULLY-COUPLED MODEL



A fully-coupled tsunami-tide models can be built to evaluate the interaction of these processes.

The computational demand for a deterministic scenario, though, is high. Demand will be much higher for nPTHA

$$\frac{\partial h}{\partial t} + \frac{\partial h\bar{u}}{\partial x} + \frac{\partial h\bar{v}}{\partial y} = 0, \quad (1)$$

$$\frac{\partial}{\partial t}(h\bar{u}) + \frac{\partial}{\partial x}(h\bar{u}^2) + \frac{\partial}{\partial y}(h\bar{v}\bar{u}) = fh\bar{v} - gh\frac{\partial\eta}{\partial x} - \frac{n^2g\bar{u}\sqrt{\bar{u}^2 + \bar{v}^2}}{h^{\frac{1}{3}}} + \frac{\partial}{\partial x}(hT_{xx}) + \frac{\partial}{\partial y}(hT_{xy}) - h\frac{\partial\Omega}{\partial x}, \quad (2)$$

$$\frac{\partial}{\partial t}(h\bar{v}) + \frac{\partial}{\partial x}(h\bar{u}\bar{v}) + \frac{\partial}{\partial y}(h\bar{v}^2) = -fh\bar{u} - gh\frac{\partial\eta}{\partial y} - \frac{n^2g\bar{v}\sqrt{\bar{u}^2 + \bar{v}^2}}{h^{\frac{1}{3}}} + \frac{\partial}{\partial x}(hT_{yx}) + \frac{\partial}{\partial y}(hT_{yy}) - h\frac{\partial\Omega}{\partial y}, \quad (3)$$

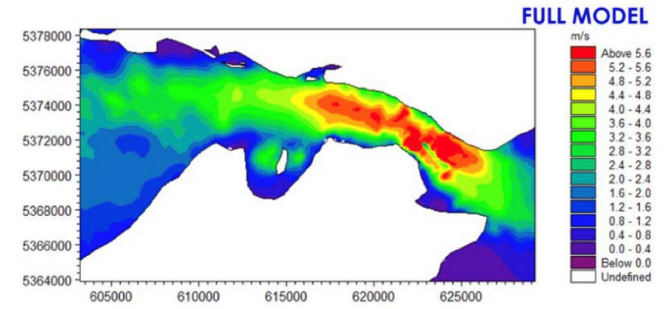
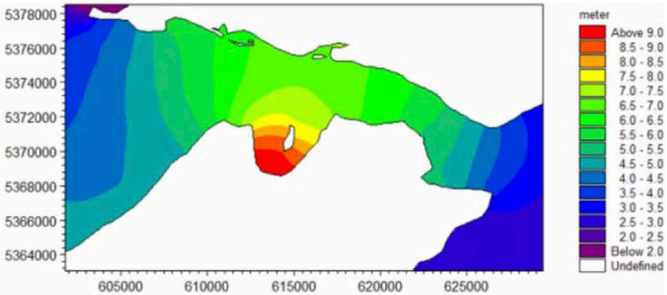
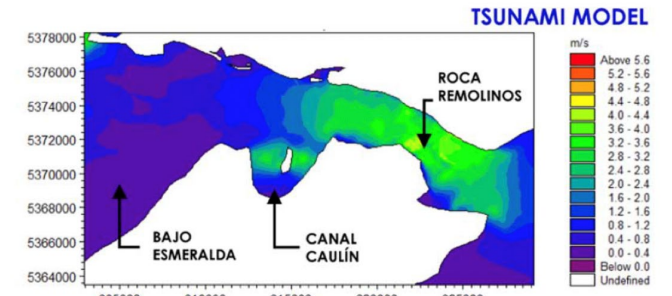
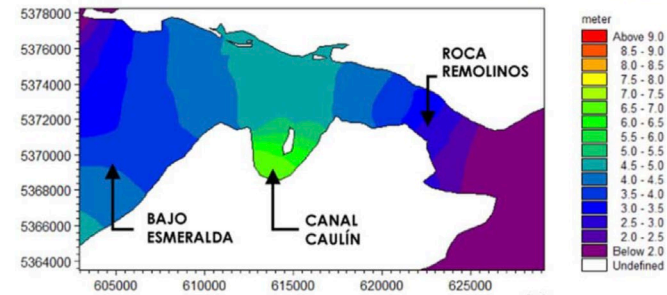
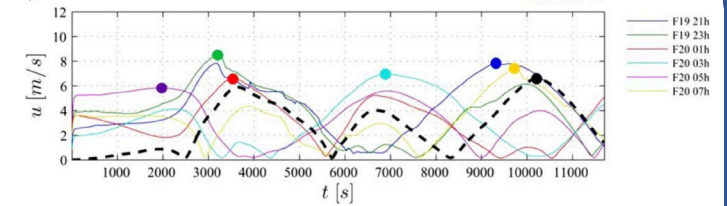
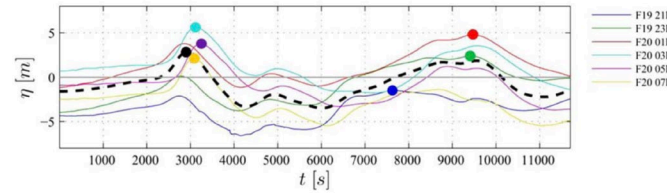
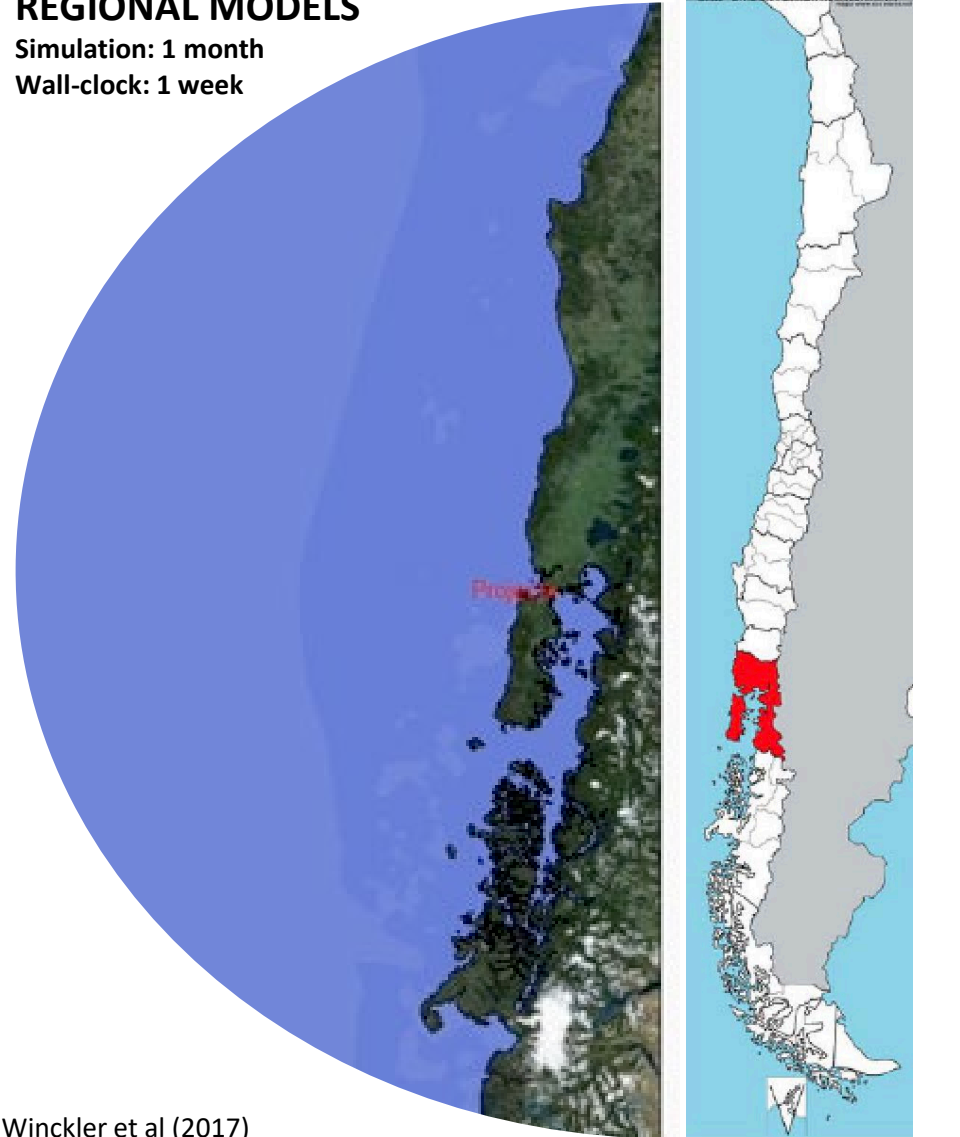
# TSUNAMI-TIDE INTERACTION



## TSUNAMI-TIDE INTERACTION: THE DETERMINISTIC FULLY-COUPLED MODEL

### REGIONAL MODELS

Simulation: 1 month  
Wall-clock: 1 week



TSUNAMI MODEL

FULL MODEL

# TSUNAMI-TIDE INTERACTION



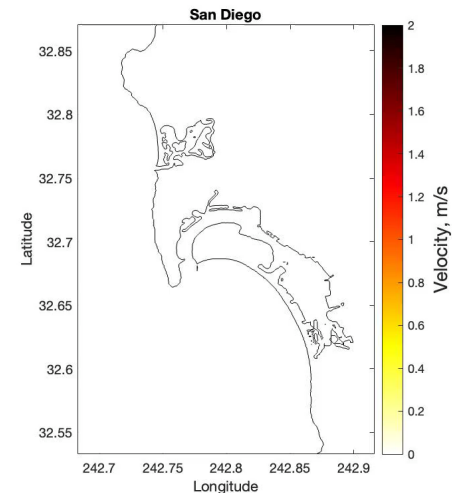
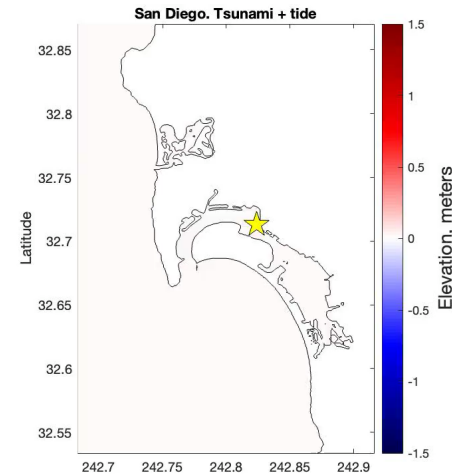
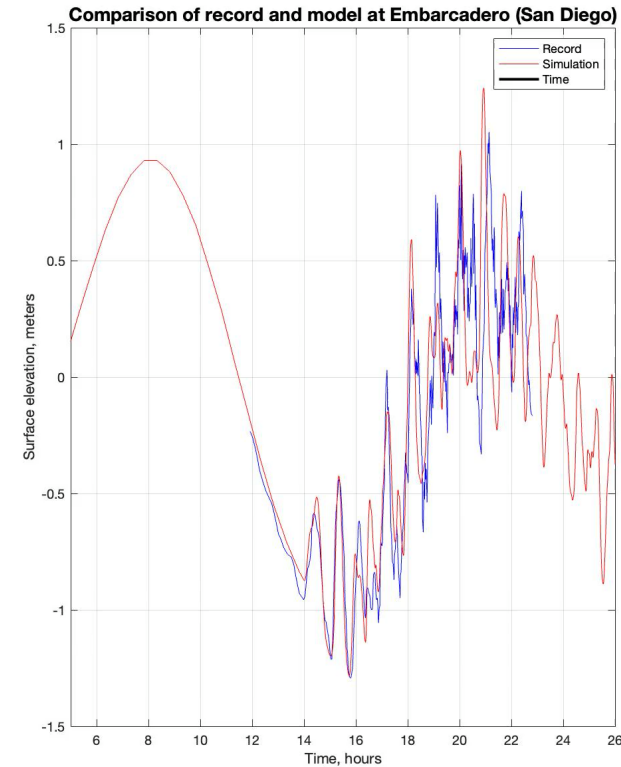
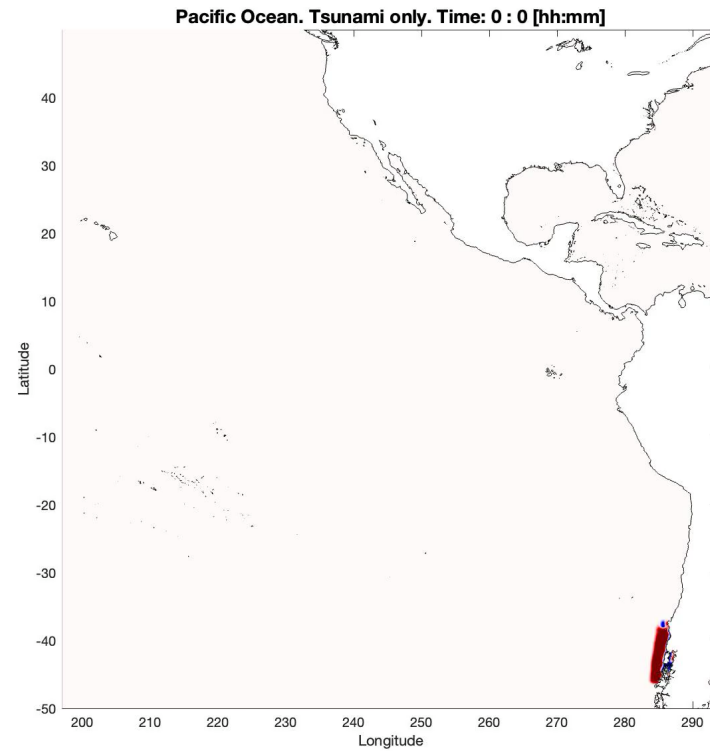
## TSUNAMI-TIDE INTERACTION: THE DETERMINISTIC ONE-WAY-COUPLED MODEL

We may simplify the problem depending on the case of study.

For example, in San Diego tide temporal variation are relevant but tidal currents are not high.

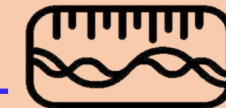
We set a one-way coupling tsunami-tide model. Tides are specified from input files...

Tsunami in California caused by 1960 Chile Earthquake  
Coastal Engineering Lab. San Diego State University



```
*****  
*                               *  
*          COMCOT TIM           *  
*                               *  
*  Cornell Multi-grid Coupled Tsunami Tide model  *  
*                               *  
*  Version Sepu 1.8xx          *  
*  updated on 2024-10-25      *  
*                               *  
*  TIM developed by           *  
*                               *  
*  Dr Ignacio Sepulveda       *  
*  Email: isepulveda@sdsu.edu *  
*                               *  
*  COMCOT version developed by *  
*                               *  
*  Dr Xiaoming Wang           *  
*  Email: x.wang@gns.cri.nz   *  
*                               *  
*****
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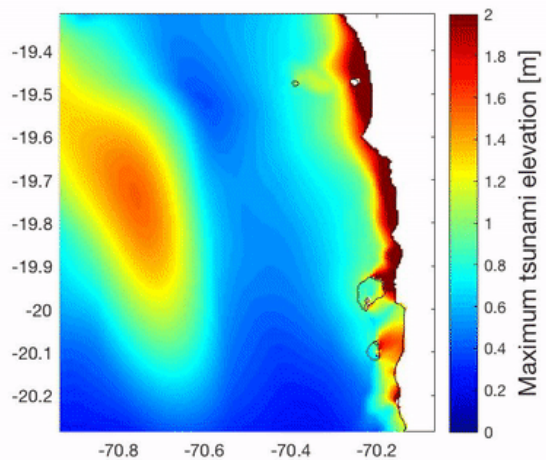
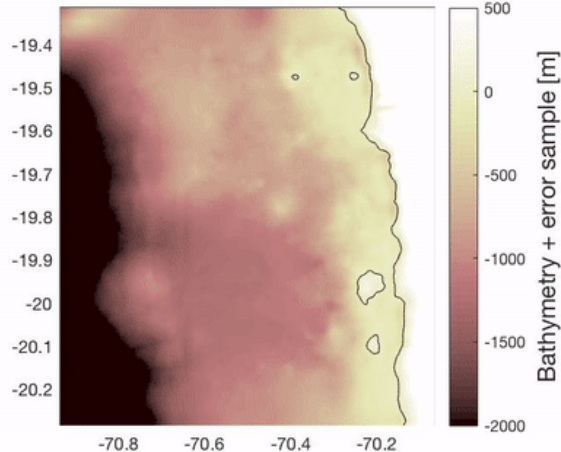
# TSUNAMI MODELING WITH UNCERTAIN BATHYMETRY



SHOULD BE WORRIED?

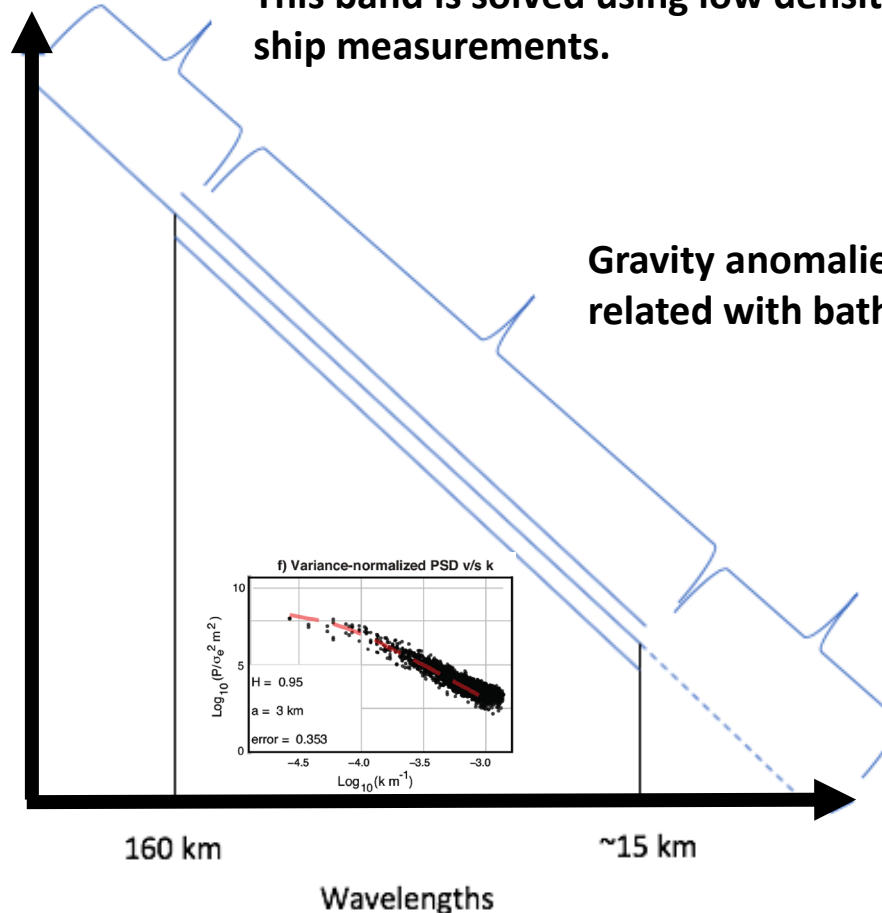
## Satellite bathymetry from a spectral perspective

The bathymetry exhibits a fractal behavior. It follows a power law.



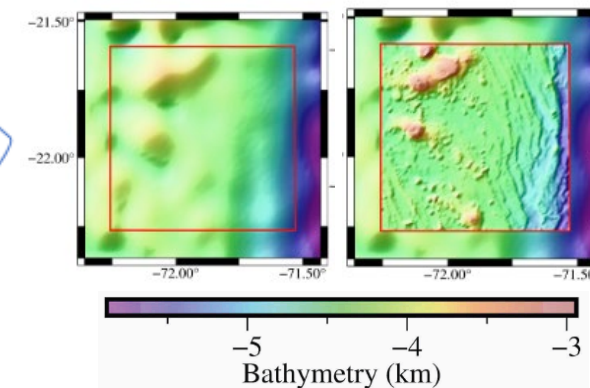
This band is solved using low density ship measurements.

Power spectral density of estimated bathymetry



Gravity anomalies,  $g$ , is linearly related with bathymetry,  $h$ .

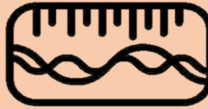
High wavenumber content not recovered. **The errors are related with this content.**



DEM partially uses satellite bathymetry or sparse ship-board soundings.

What is the effect of errors?

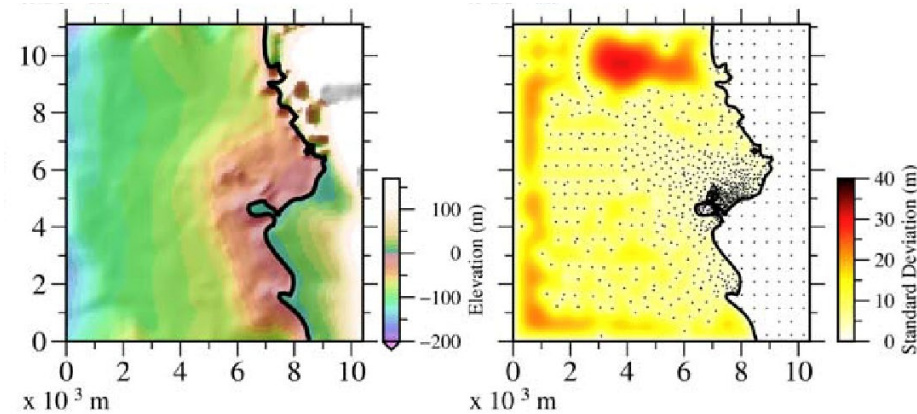
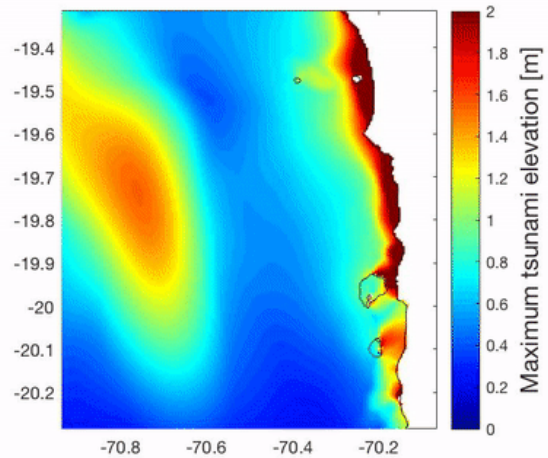
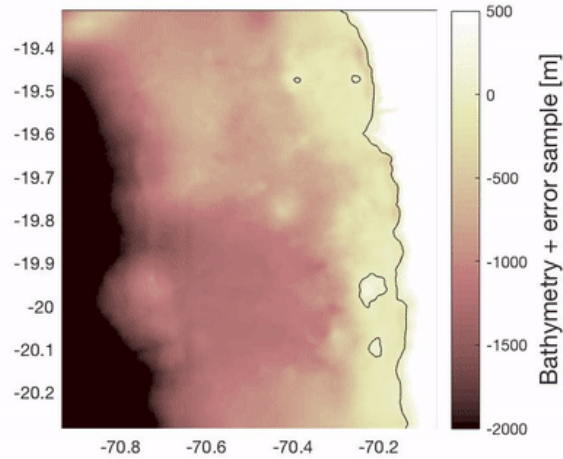
# TSUNAMI MODELING WITH UNCERTAIN BATHYMETRY



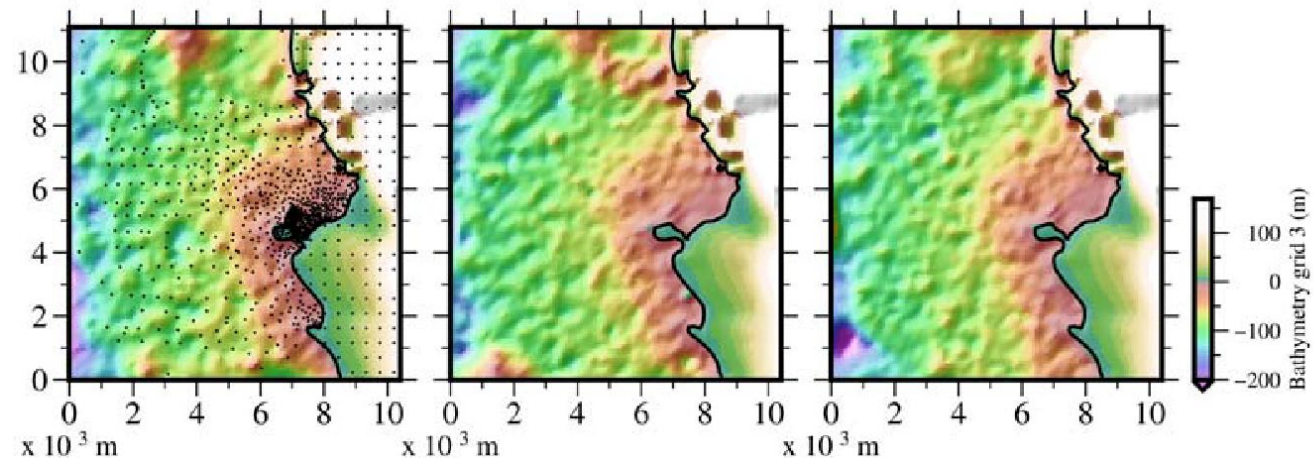
SHOULD BE WORRIED?

## Interpolation of bathymetry from a spectral perspective

We use a conditional random field to model bathymetry uncertainties. This allow to design a random interpolator of bathymetry.



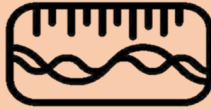
Samples of interpolated bathymetry plus synthetic high-wavenumber content



DEM partially uses satellite bathymetry or sparse ship-board soundings.

What is the effect of errors?

# TSUNAMI MODELING WITH UNCERTAIN BATHYMETRY

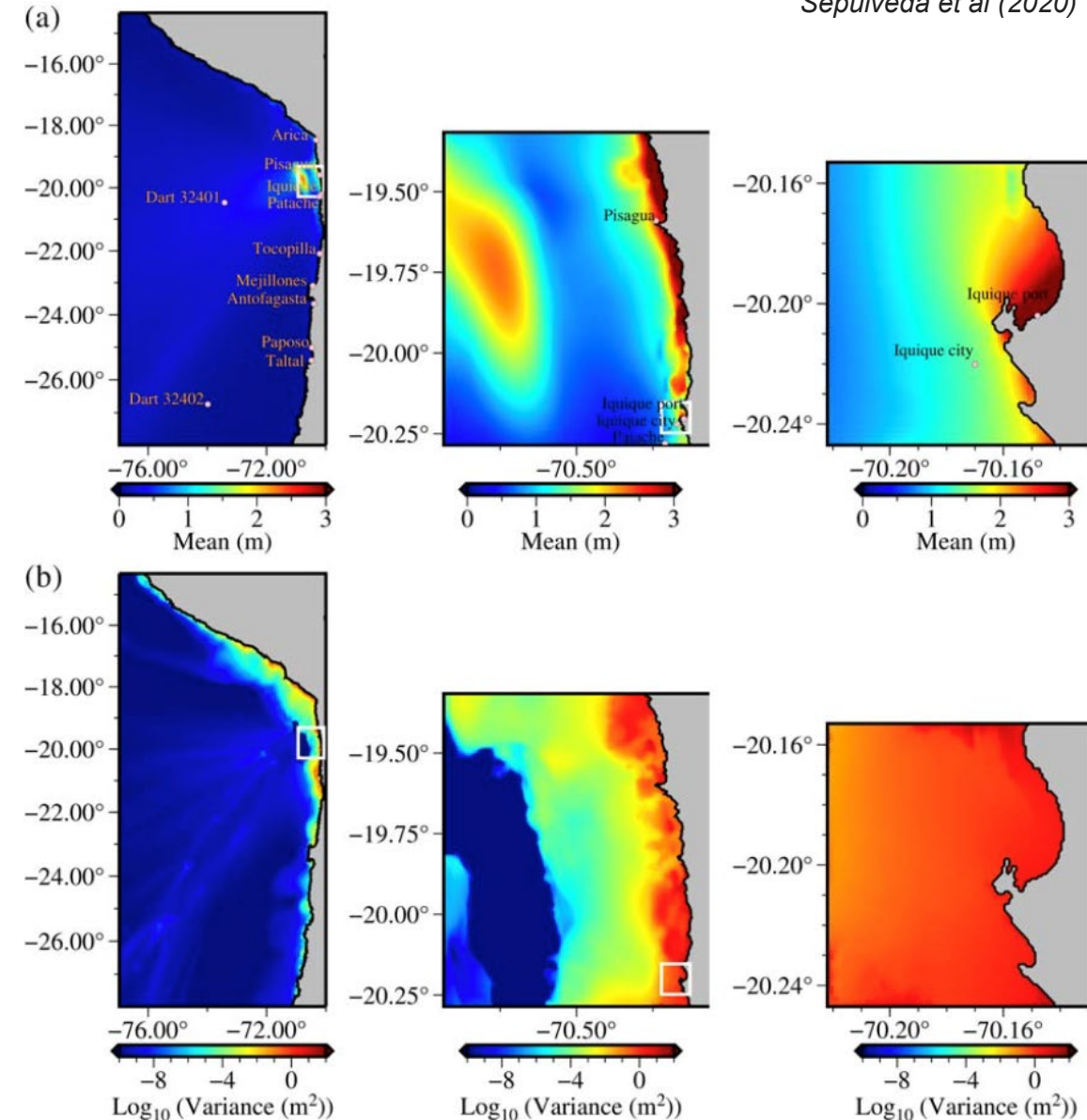
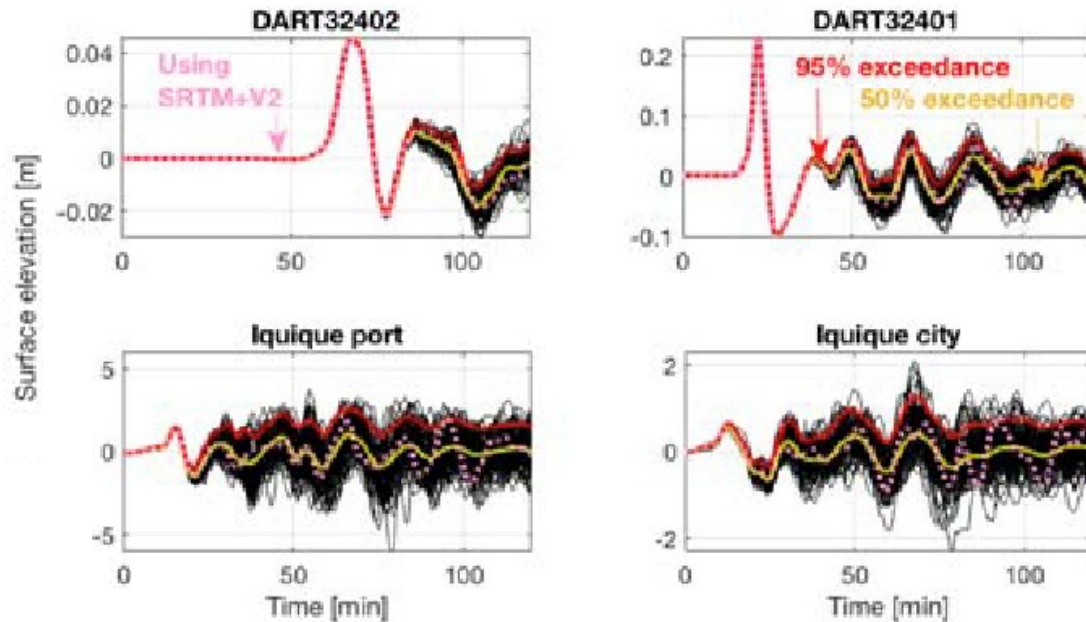


SHOULD BE WORRIED?

*Sepulveda et al (2020)*

1) Trailing waves (not leading wave) are sensitive to bathymetry errors, even in shallow depths. High sensitivity may be explained by shorter tsunami wavelengths and longer time exposed to large depth relative errors.

2) Leading waves not propagating through shallow depths are not sensitive







# SDSU

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Coastal  
Engineering Lab



“... is uncorrupted relationship, taken from  
the truth, cut to size...”

“La Araucana” by Alonso de Ercilla (1569-89)